

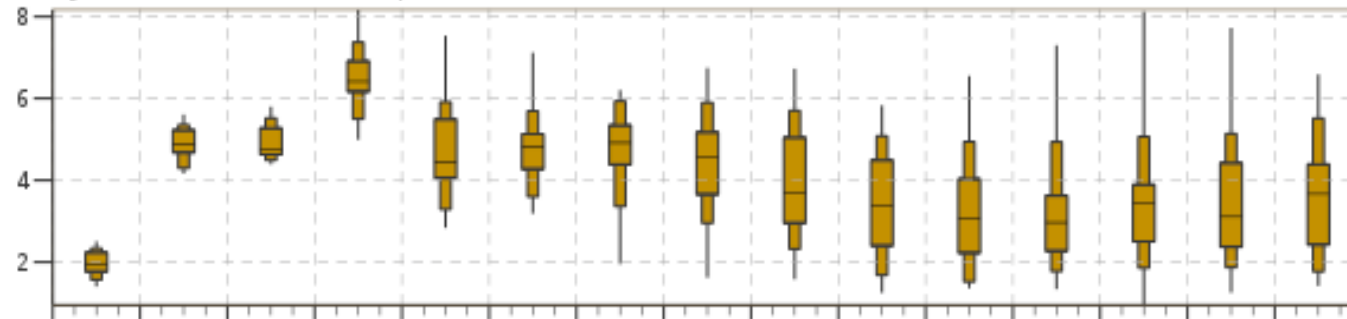
Introduction to intrinsically stochastic parameterization

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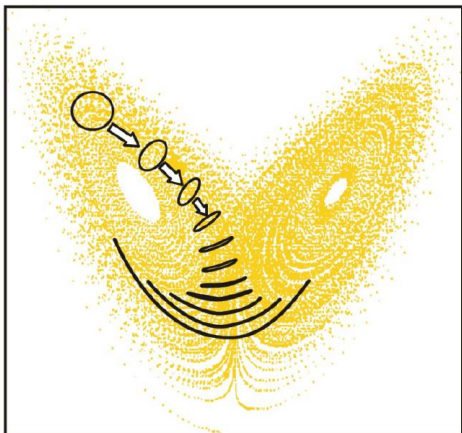


Chaos, probability and model uncertainty

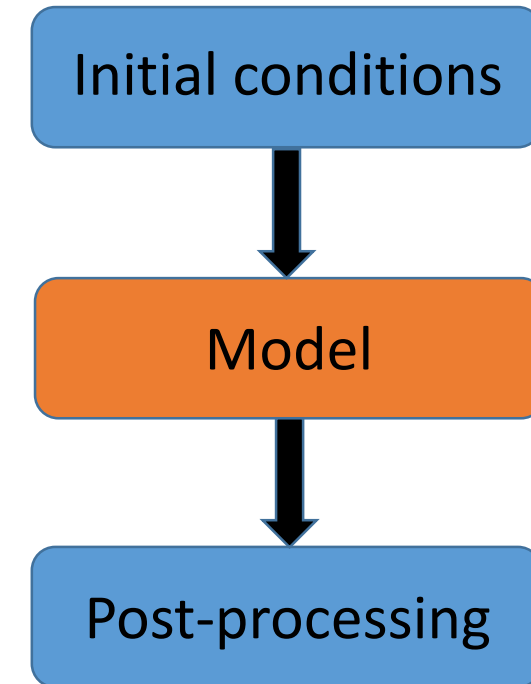
Daily mean of 10m Wind Speed (m/s)



2. Forecasts need to show uncertainty



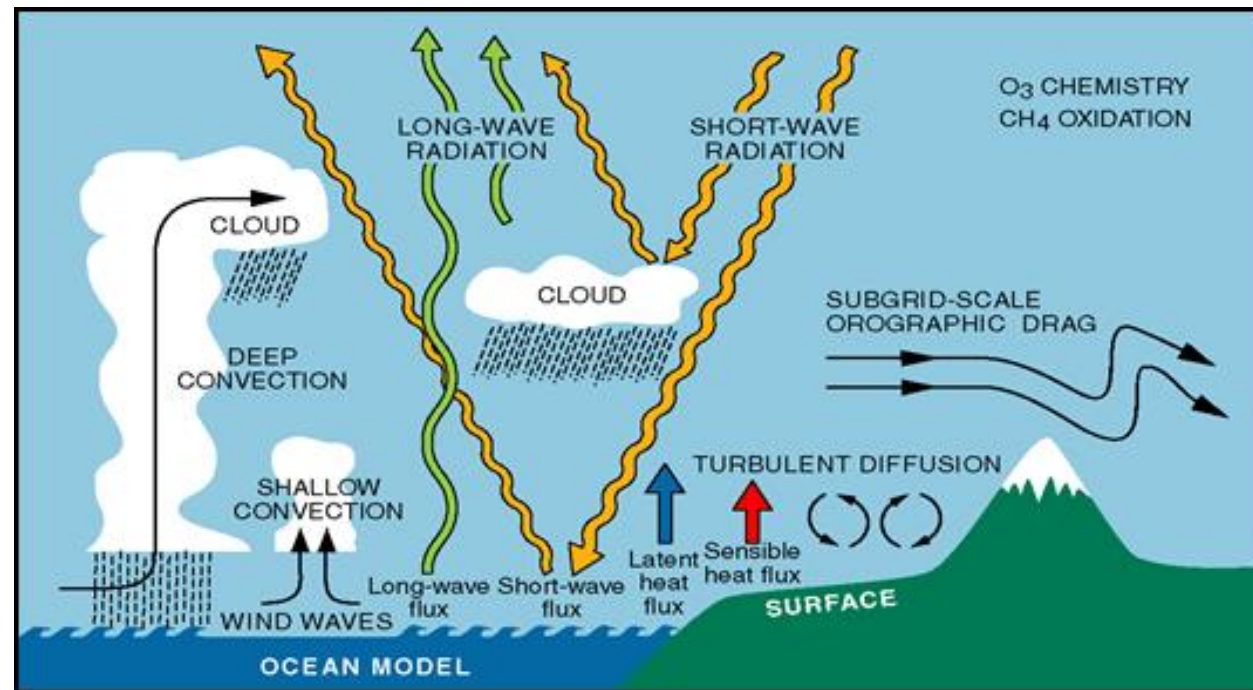
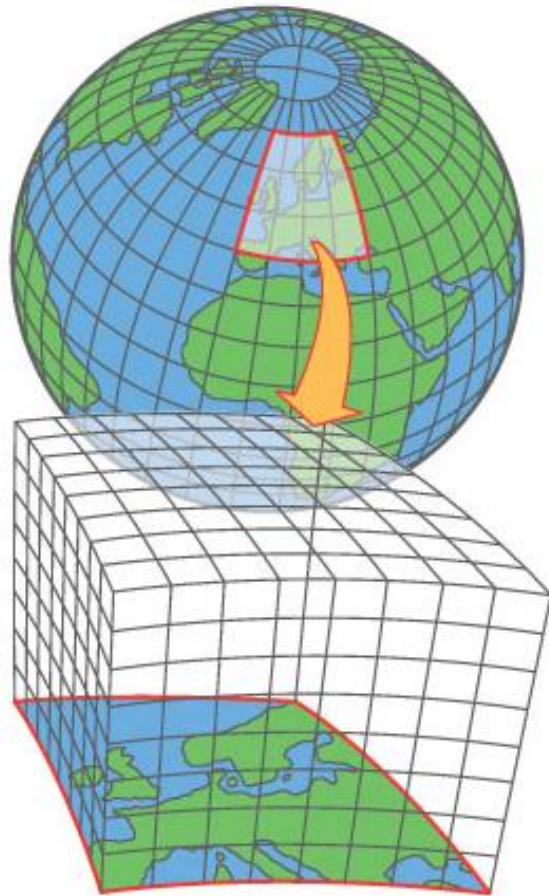
1. Predictability is intrinsically limited



3. Each part of the NWP chain contributes uncertainty

Model uncertainties

- Finite resolution – unresolved motions
- Physical processes – not part of fluid equations



Types of model error

Systematic errors

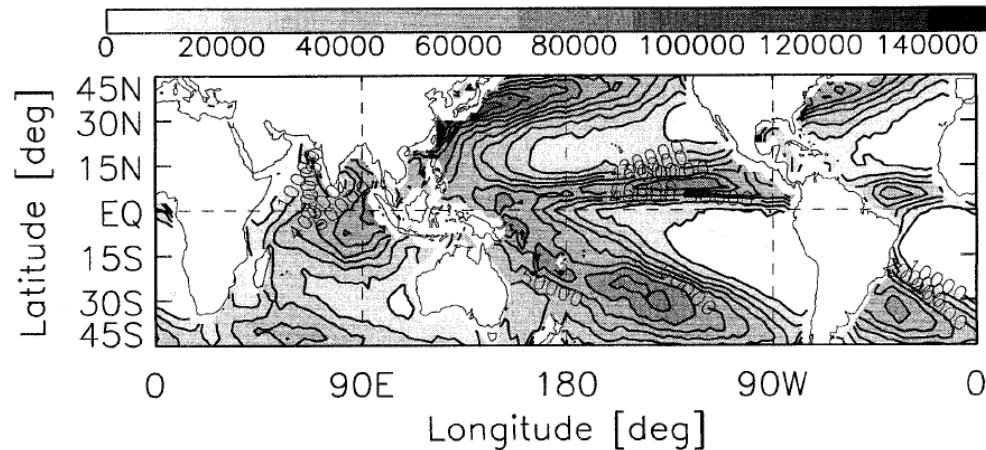
- e.g. dry bias in boundary layer leads to lack of convective initiation
- Need multiparameter or multimodel ensemble

Random variability

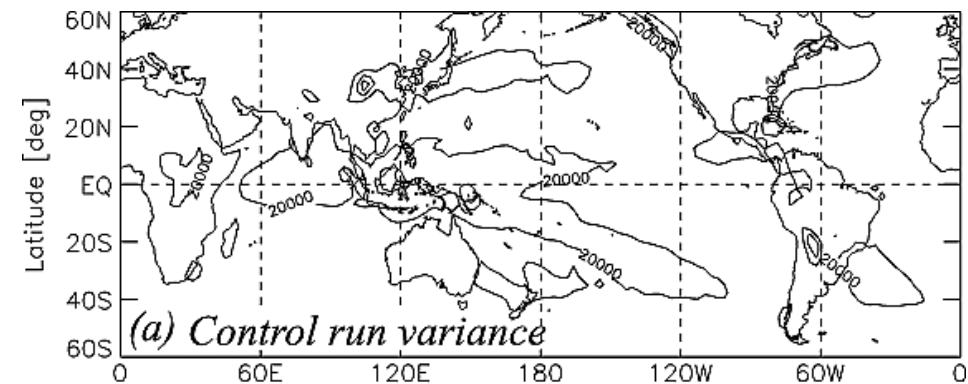
- e.g. unresolved motions trigger convection in wrong place
- Need stochastic variability among ensemble members

So add the missing variability, right?

Observed precipitation variance



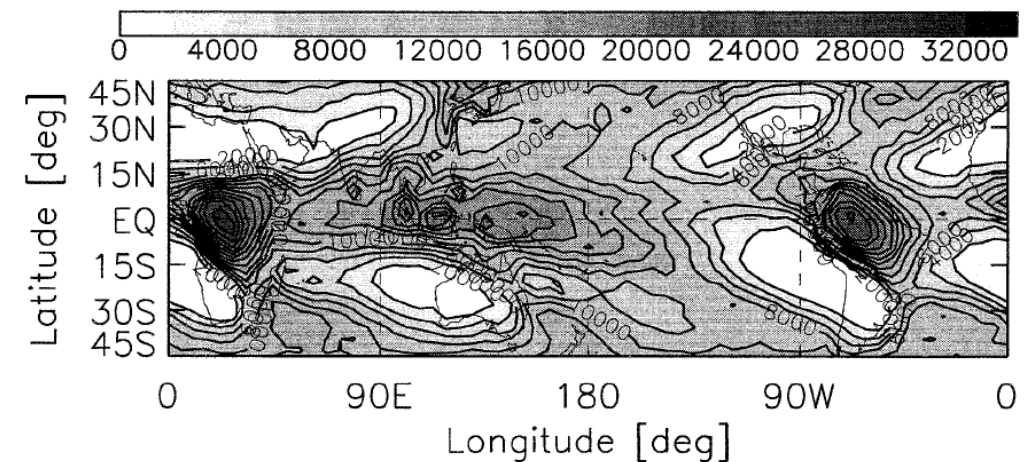
Default model has almost no variance



Response to forcing is nonlinear!

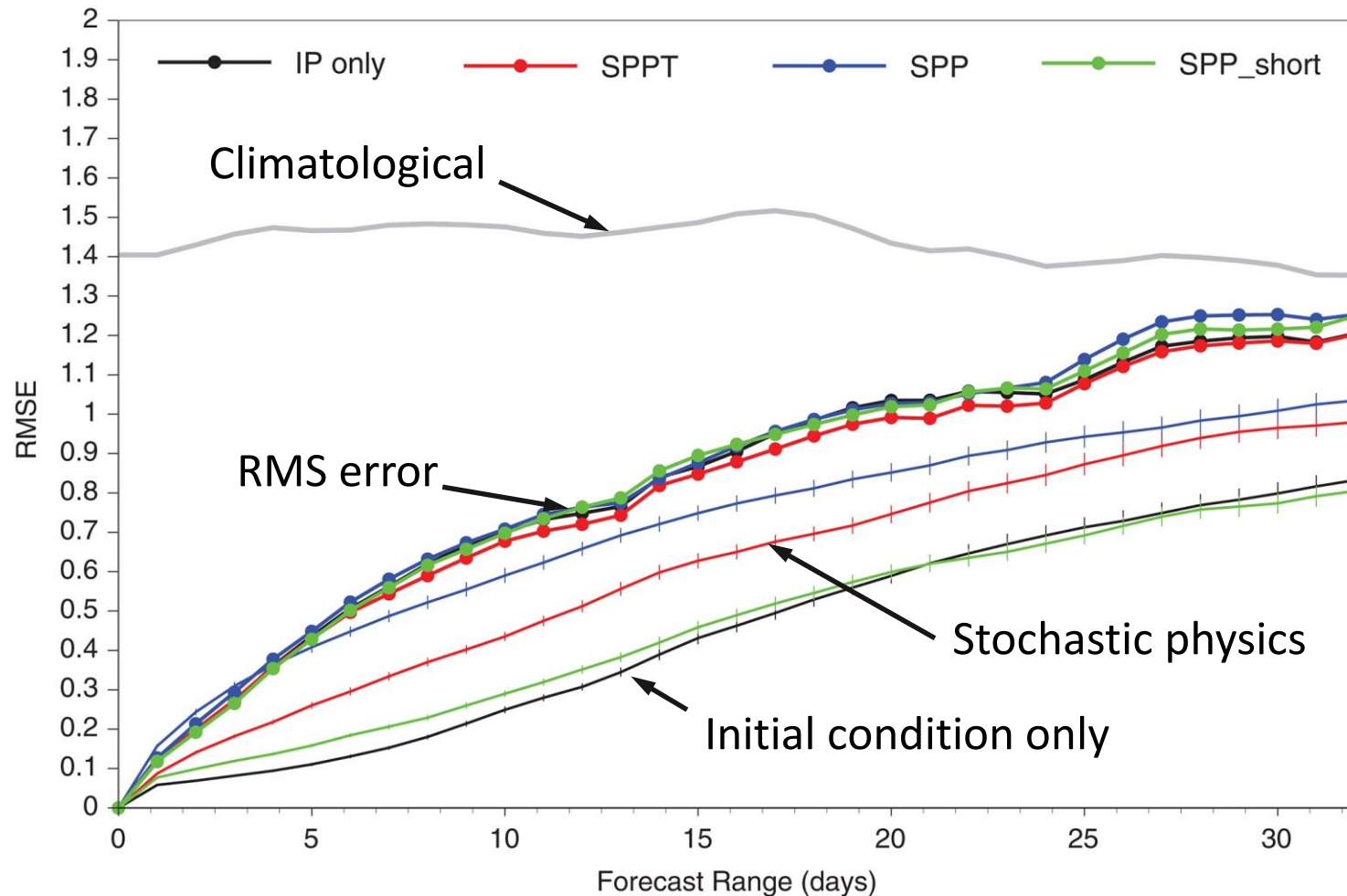
- Lin and Neelin (JAS 2002) added stochastic forcing to GCM based on observed precipitation variance
- Model variance increased in response to forcing but with different properties

Model variance with stochastic forcing



Lin and Neelin, JAS 2002

Pragmatic solution – look at response

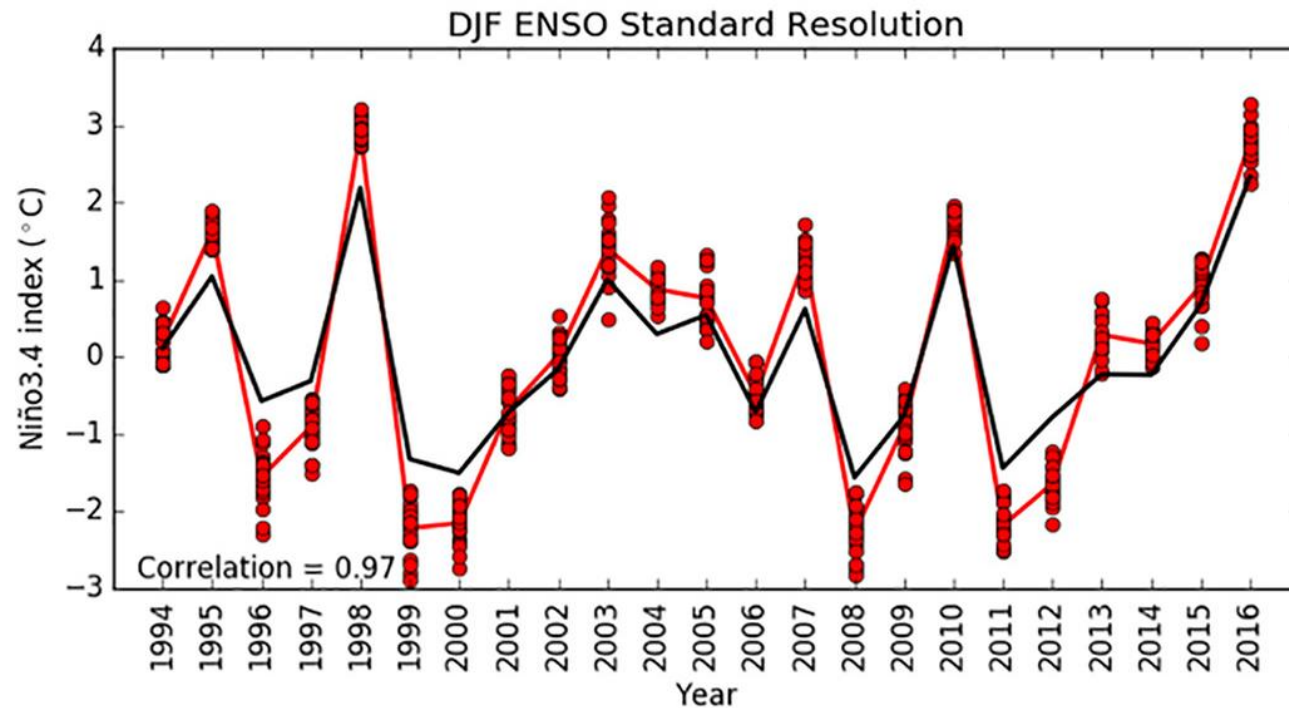


- Ensemble is over-confident – spread does not capture uncertainty
- Calibrate ensemble by comparing spread and skill (eg. ECMWF well-calibrated)
- Reality should resemble random ensemble member

But:

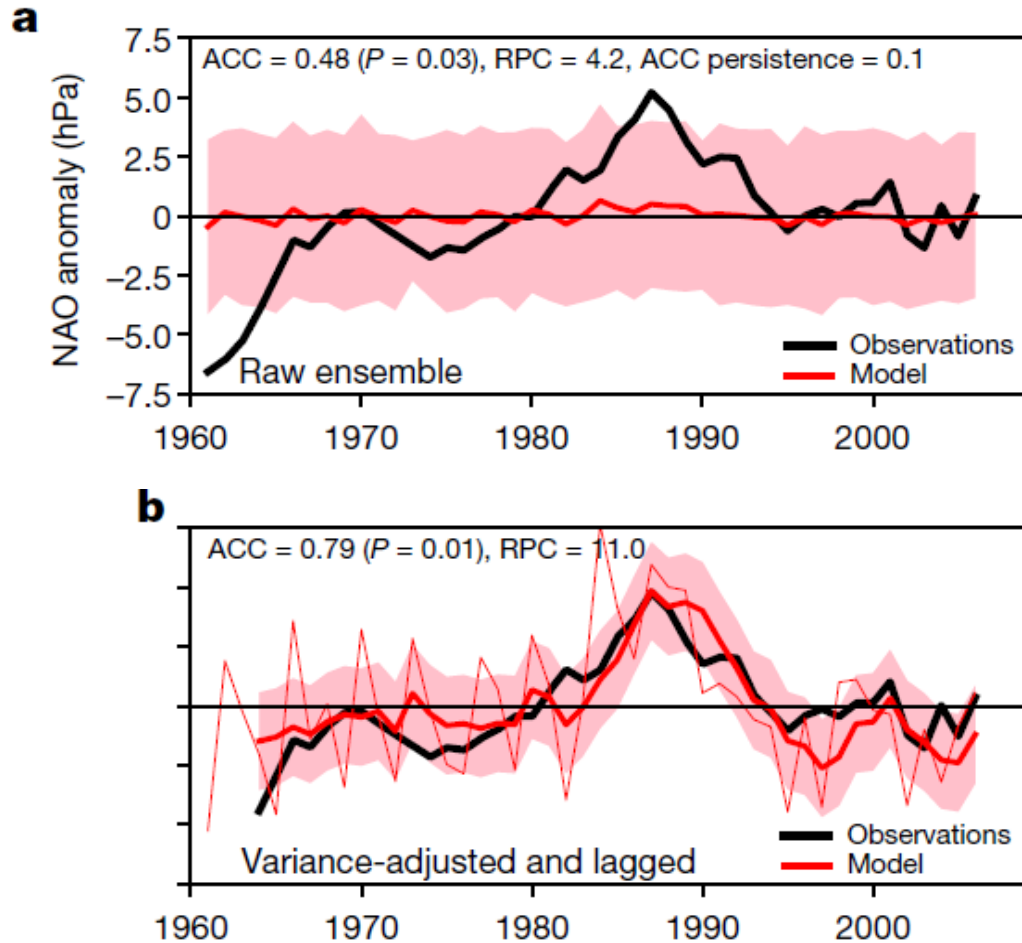
- Calibration expensive
- Incomplete description of variability

Variability on different scales I



- Hadley Centre ENSO forecast over-confident
- Problem is ENSO signal too strong, not lack of variability
- More stochastic perturbations could degrade forecast

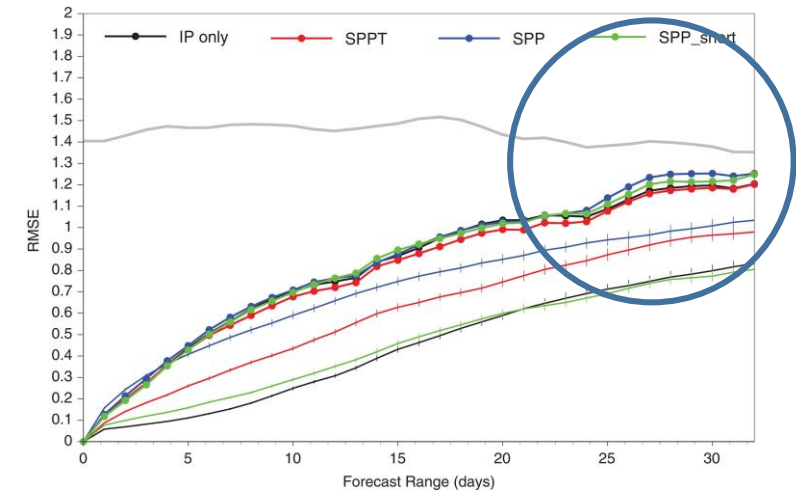
Variability on different scales II



- Hadley Centre decadal NAO forecast has correct variance, but on short timescales
- Mean of large ensemble can average out variance and correlates with observations
- Reality very different from ensemble members (S/N ratio paradox)

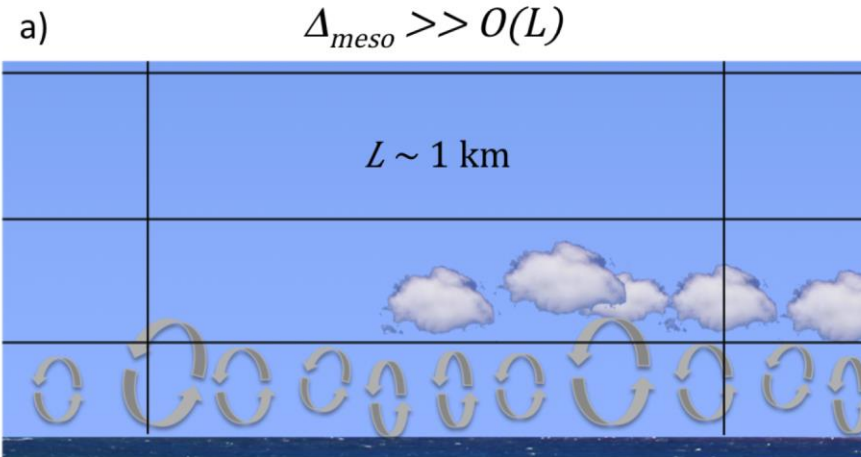
Variance on different scales

- Random model variability (stochastic physics) mainly “small-scale”
 - Also have larger scale “climatological” variability
 - Not enough to consider spread-skill calibration
 - *Analogy to convective vs synoptic scale?*
-
- Need to consider physical processes behind variability
 - intrinsically stochastic parameterization

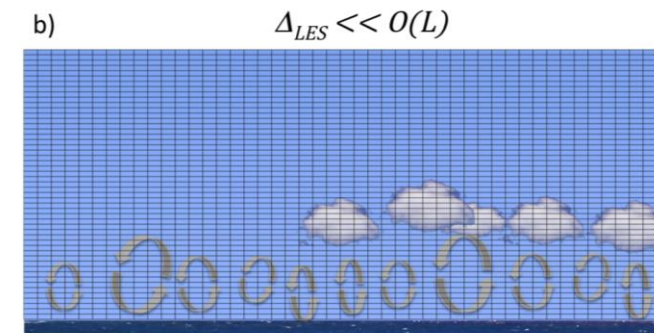
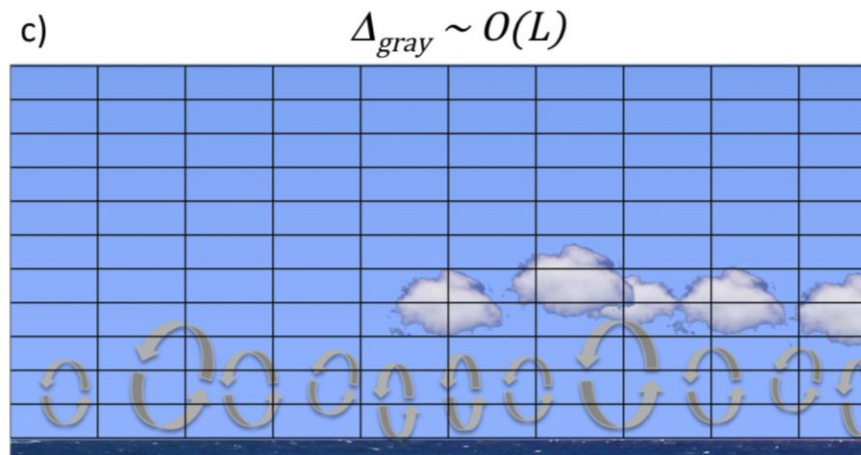


Parameterizations are intrinsically stochastic

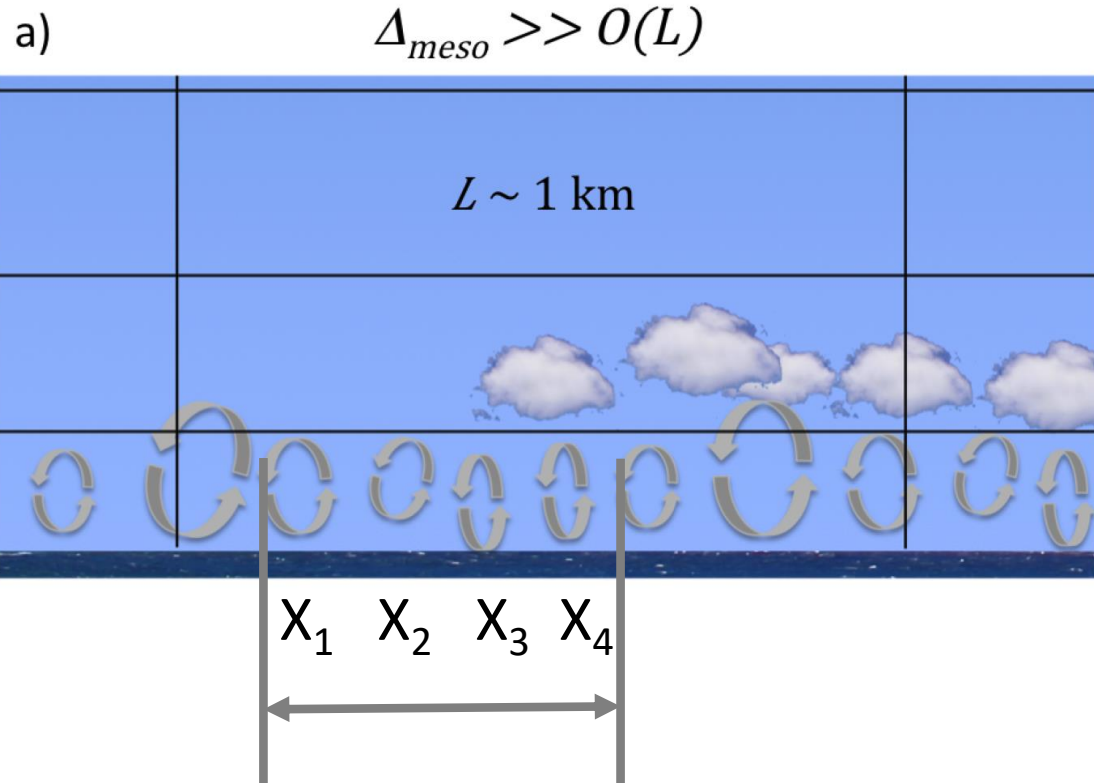
- Parameterization gives net effect of many small elements (size L)
- Scale separation, grid length $\Delta \gg L$, ensures well-defined result
- Gray zone, $\Delta \sim L$, range of possible results on grid scale – **stochastic**



...or resolve elements $\Delta \ll L$, LES



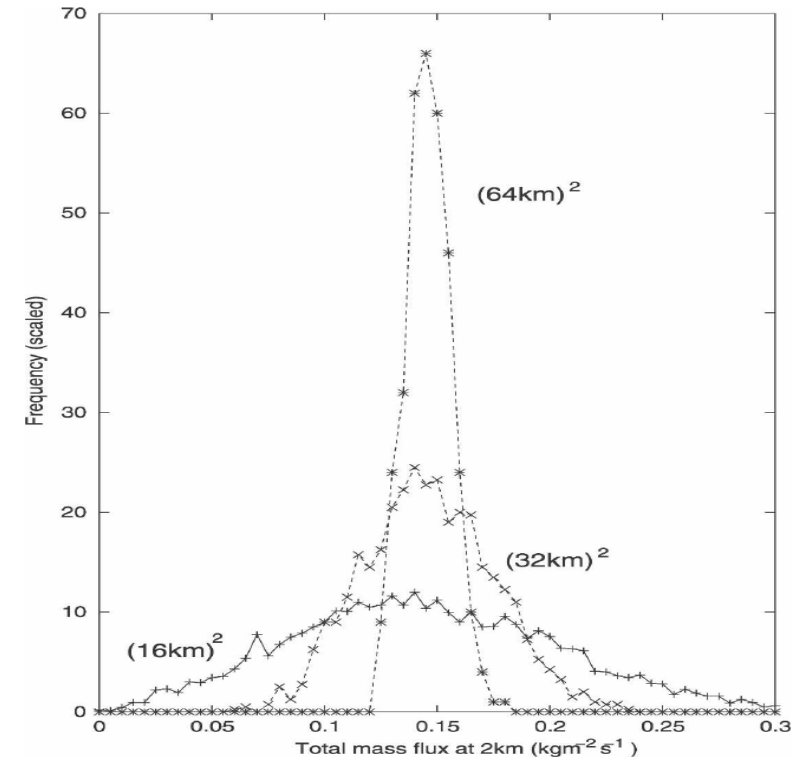
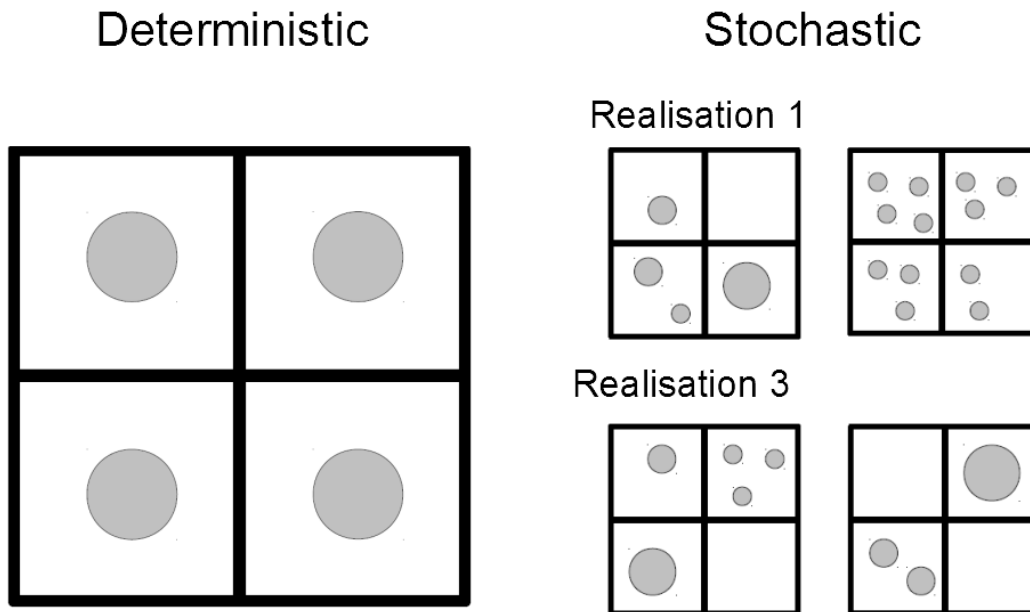
Theory for fluctuations (Einstein 1905)



- Want net effect of (small) number N of subgrid features
- Sum of N random variables $X_1 + X_2 + \dots$
- Central limit theorem says if N not too small, sum follows Gaussian distribution
- Standard deviation $\sigma = C/\text{sqrt}(N)$
- C depends on environment (but not on N)

Stochastic convection (Plant and C. 2008)

- Average number of clouds: $N = M / m$
 = total mass flux (closure) / mass flux per cloud (constant)
- Analytic solution is Poisson-like process for M
 with $\sigma_M = 2M/\text{sqrt}(N)$



Plant and Craig JAS 2008

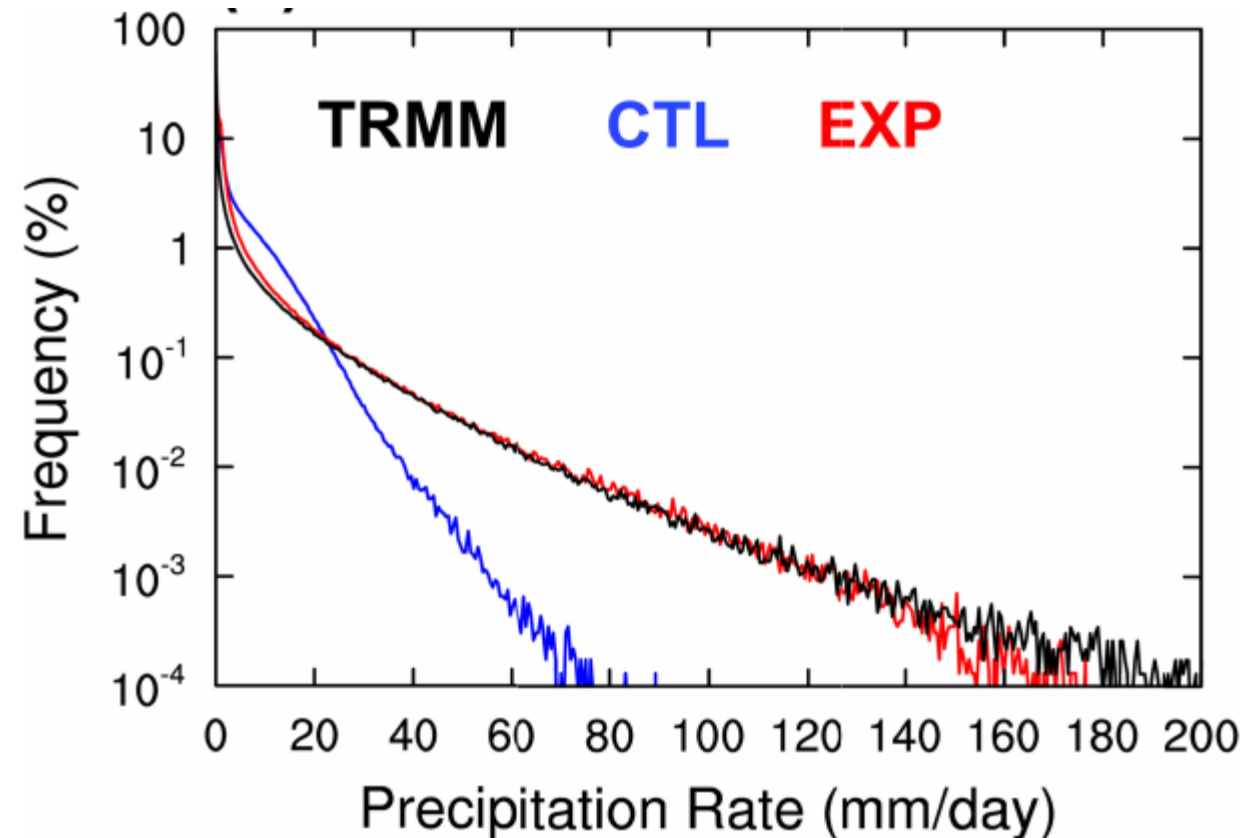
Impact on precipitation rates

- NCAR CAM5
- 10 year climate simulation
- Frequency of precipitation rates in tropics (20S – 20N)

Observations

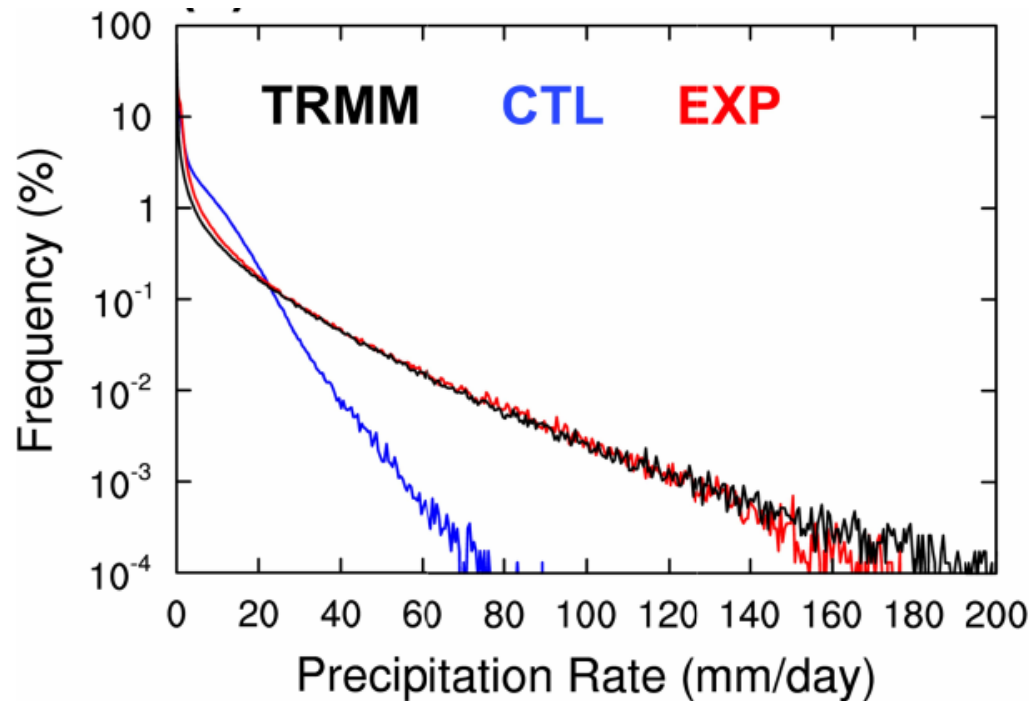
Default Zhang-McFarlane convection scheme

Stochastic ZMPC

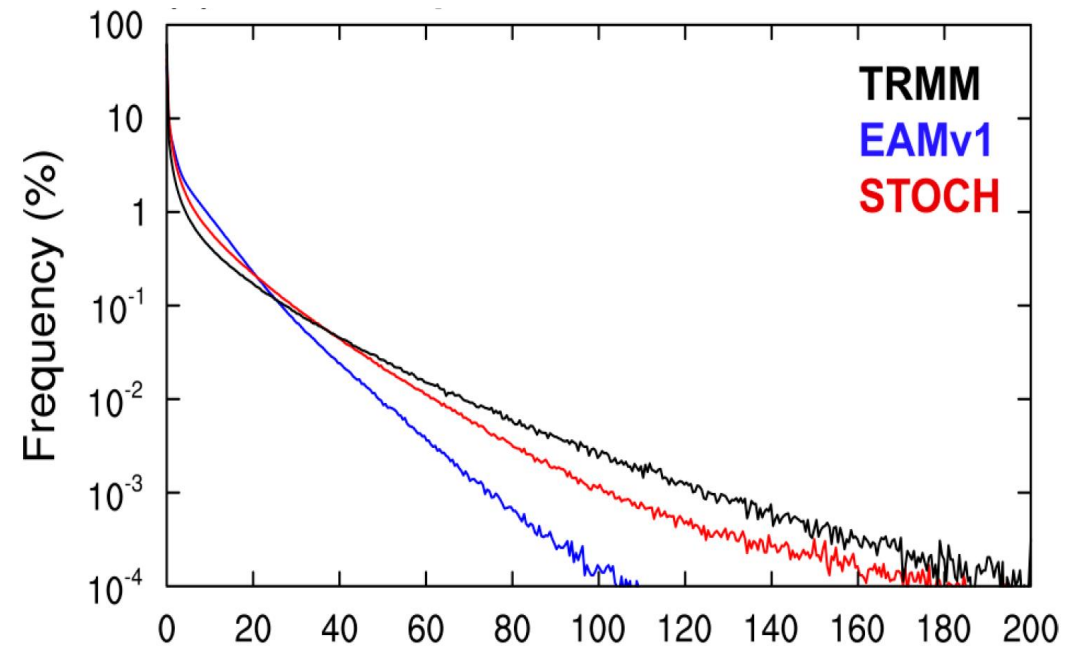


Model dependence

NCAR CAM5

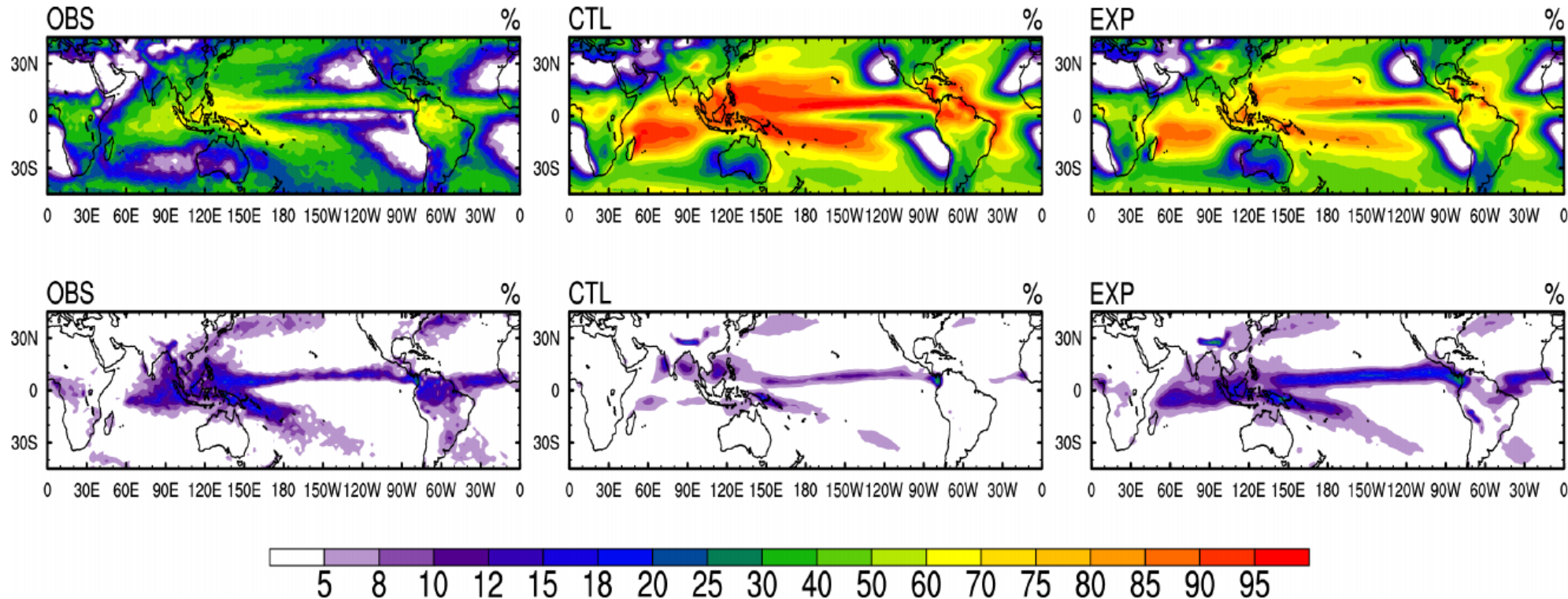


DOE E3SMv1



Sensitivity to vertical resolution

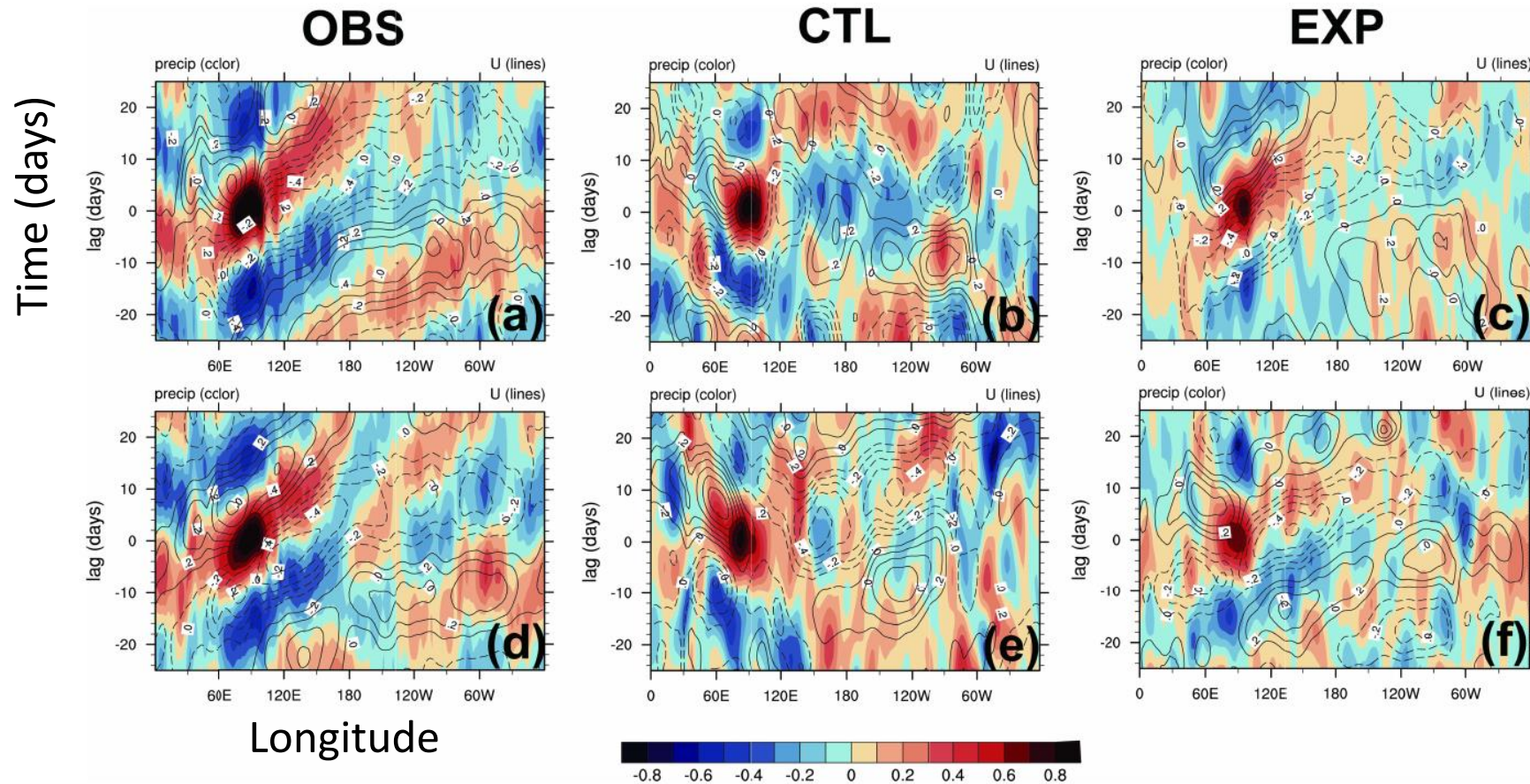
Distribution of Light/Heavy precipitation



Frequency of precipitation > 1 mm/day (top row)

> 10 mm/day (bottom row)

MJO variability



20-100 day bandpass filtered precipitation and U850 (20S-20N)

May-Oct (top, Nov-Apr (bottom))

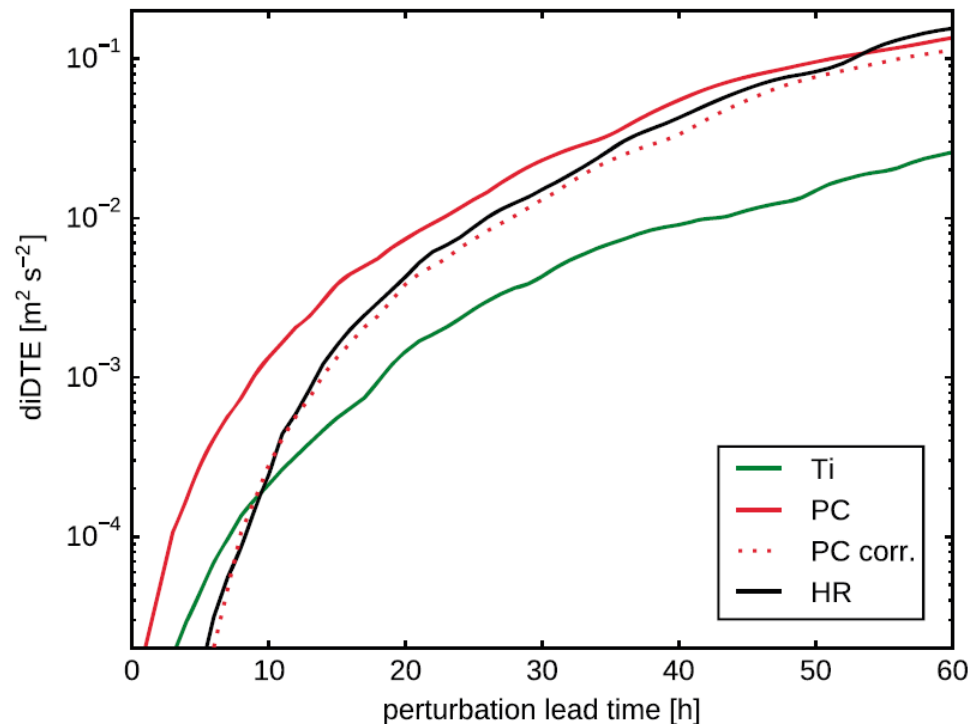
Perturbation growth experiments

Difference Total Energy (DTE):

$$DTE := \frac{1}{2} \left(\Delta u^2 + \Delta v^2 + \frac{c_p}{T_0} \Delta T^2 \right)$$

- COSMO domain covering Europe and western Atlantic, 4 start times
- Perturbed by small-scale noise in initial conditions and stochastic convection (PC)

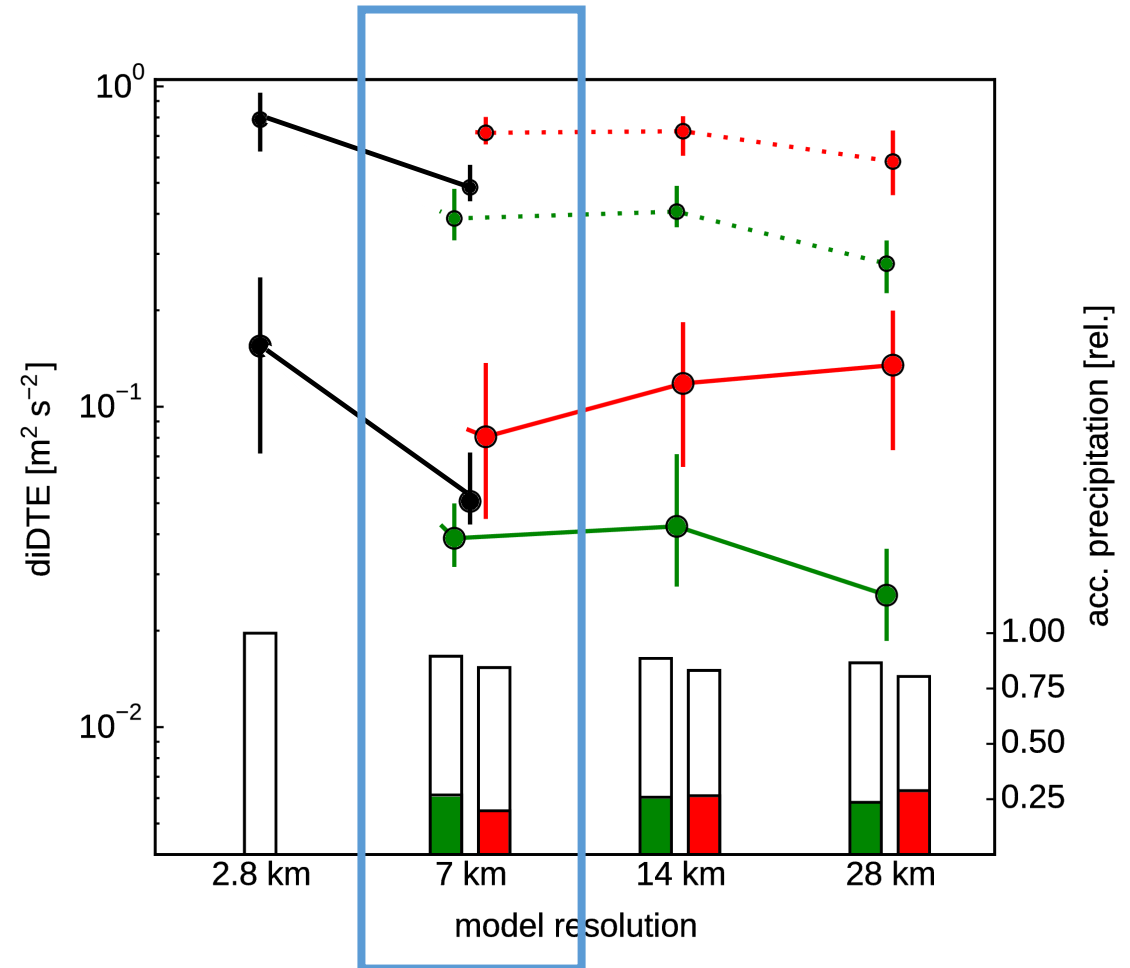
- HR: 2.8 km resolution, explicit convection
- Ti: 28 km resolution, default Tiedtke scheme
- PC: 28 km resolution, PC stochastic scheme
- PC corr: like PC but shifted 5 h (spin-up time for convective variability)



Scale invariance

- DTE on medium (dashed) and large (solid) scales after 60 h perturbation growth
- **No parameterization** – growth damped at low resolution
- **Default Tiedtke scheme** – too little growth
- **Plant-Craig stochastic** – realistic growth independent of resolution

Resolution 7 km
– gray zone



Criteria for a stochastic parameterisation

1. Is the the scheme stable and well-behaved in the full model?
(e.g. resolution dependence)
2. Is the variability contributed by the scheme significant?
(compared to initial condition uncertainty, etc.)
3. Is the forecast skill superior to that obtained with a deterministic scheme? (on some score!)
4. Are there nontrivial interactions with the resolved flow?
(Could the same skill be obtained by postprocessing output of model with deterministic scheme?)
5. Could the same skill be achieved with an inexpensive ad hoc scheme?

Conclusions

Summary

- Parameterizations well-posed if scale separation
- Otherwise intrinsically stochastic (Schär – multiple gray zones)
 - Cumulus convection (~10 km)
 - Convective boundary layer (~1 km)
- Generic theory for variability as sum of random processes
- e.g. Stochastic convection scheme
 - Improved precipitation variability
 - Improved upscale error growth

Issues

- Interaction with “climatological variance” (convection like S2S?)
- Interaction with numerical artifacts “gray zone”