



SAINT PT: status, plan and discussion

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Proposed COSMO Priority Task (PT) Project – SAINT

- **Phase I:** validation of current multi-layer scheme – update as needed
- **Phase II:** implementation – adjustment of currently implemented parametrizations ; radiation (albedo), turbulence , tile approach ...
- **Phase III:** validation of implementation especially diagnostic parameters (e.g. T_2m)
- **Phase IV:** documentation (paper, technical report etc.)
- **Duration:** 2 years (50%), Start July 2017, End June 2019

Proposed COSMO Priority Task (PT) Project – SAINT



A MULTI-LAYER SNOW COVER MODEL FOR NUMERICAL WEATHER PREDICTION AND CLIMATE MODELS

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ABSTRACT: The seasonal snow cover strongly influences boundary layer processes such as turbulence and radiation. Therefore, knowledge of the current state of the snow cover on the ground is of paramount importance for numerical weather prediction (NWP) and climate models. This is especially true since the horizontal resolution (up to 1 km) of NWP models strongly increased in recent years. Currently, most NWP models use simplified – typically one snow layer – snow cover schemes, which are in general not capable to simulate snow cover formation, evolution and melt with adequate accuracy. Within the framework of the COSMO consortium (Consortium for Small-scale Modelling) we are developing a new multi-layer snow cover module for the regional weather forecasting and climate model COSMO and the global model ICON (Icosahedral Nonhydrostatic). The snow cover model consists of six snow layers with increasing but constant thickness for the upper part (top 29 cm) of the snow cover and a variable amount of snow layers (minimum = 2) in the lower part of the snow cover. Preliminary results indicated an improvement of the snow surface temperature with only minor effects on the near surface air temperatures, when compared to the single layer counterpart at the location of a network of high-alpine weather stations.

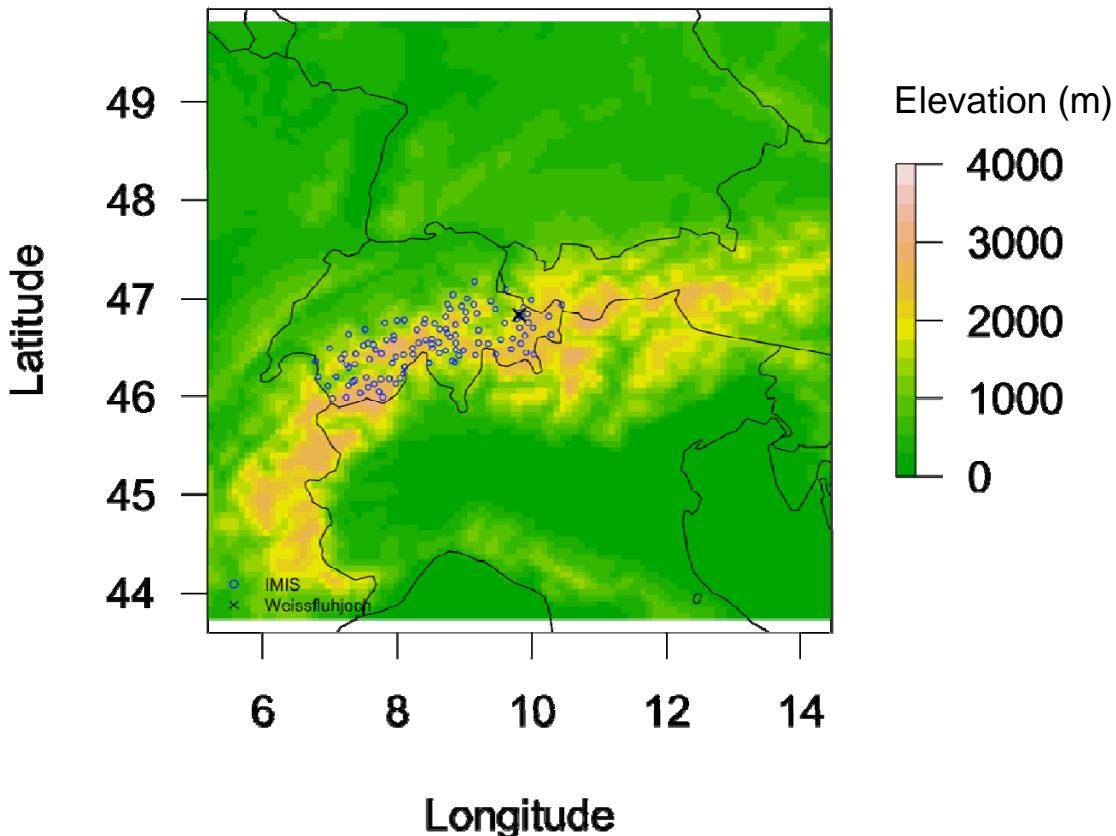
KEYWORDS: snow cover modelling, numerical weather prediction, climate modelling

1. INTRODUCTION

The seasonal mountain snow cover can lead to hazardous situations (e.g. avalanches, flooding), provides water resources during the melt phase and constitutes an important factor for global and regional weather and climate. Forecasting the snow cover evolution in complex terrain provides

heat conduction into the snow cover. Heat conduction from the surface, through the snow cover, to the snow/soil interface depends on the surface temperature, the snow cover layering, i.e. density, as well as the current temperature distribution as well as phase changes happening in the snow. Therefore, it has been shown that single layer models are basically not capable of

Phase II: Implementation/Validation – COSMO-7 Setup



- ... ~ 700 km x 700 km domain centered around Davos ...
- ... covering most of the Alpine ridge ...
- ... computationally inexpensive ...
- .. boundary conditions from COSMO-7 analysis ...
- ... 72-hour forecast/hindcast 16 February 2017 00UTC

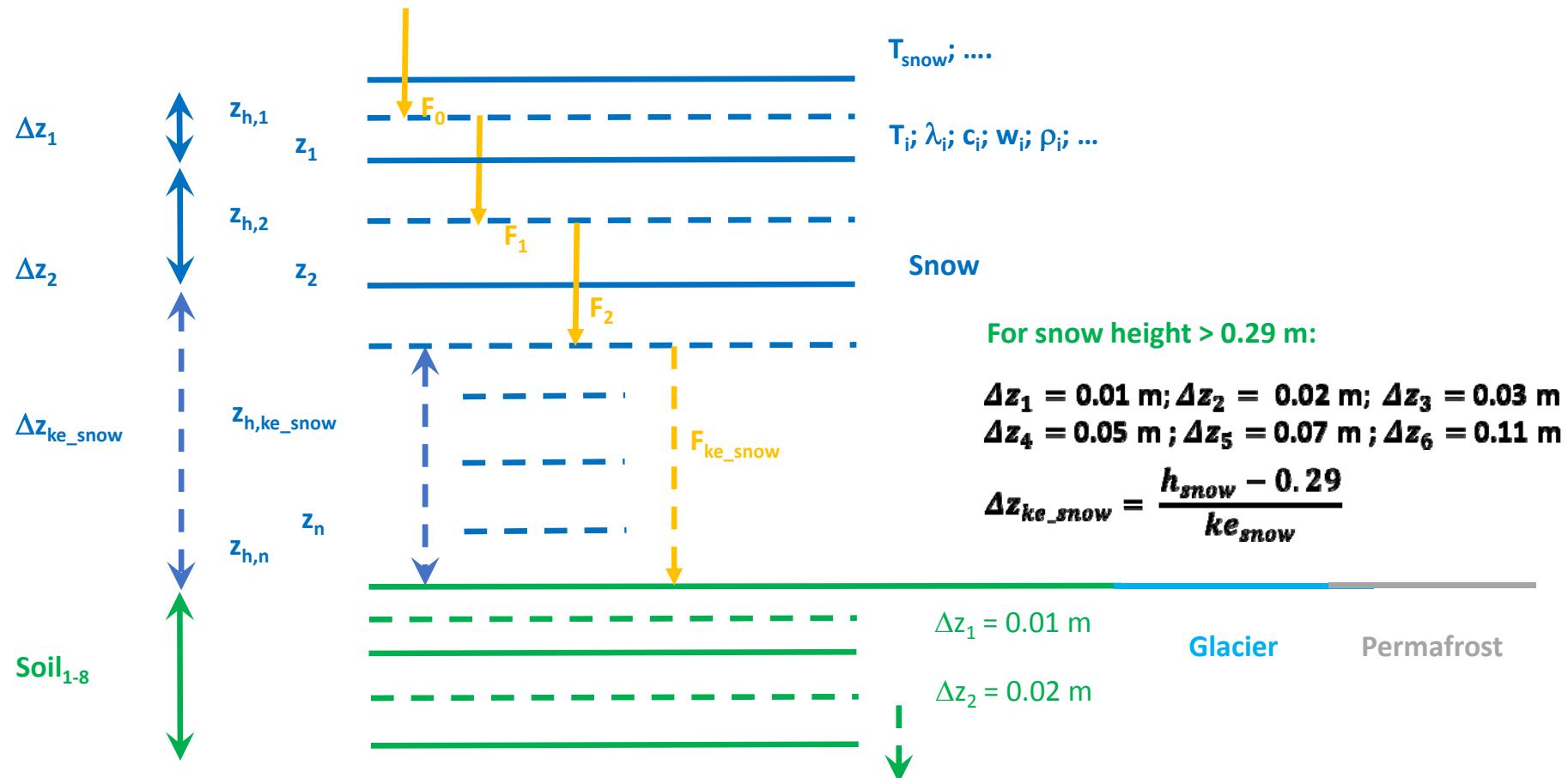
Intercantonal Measurement and Information System (IMIS)



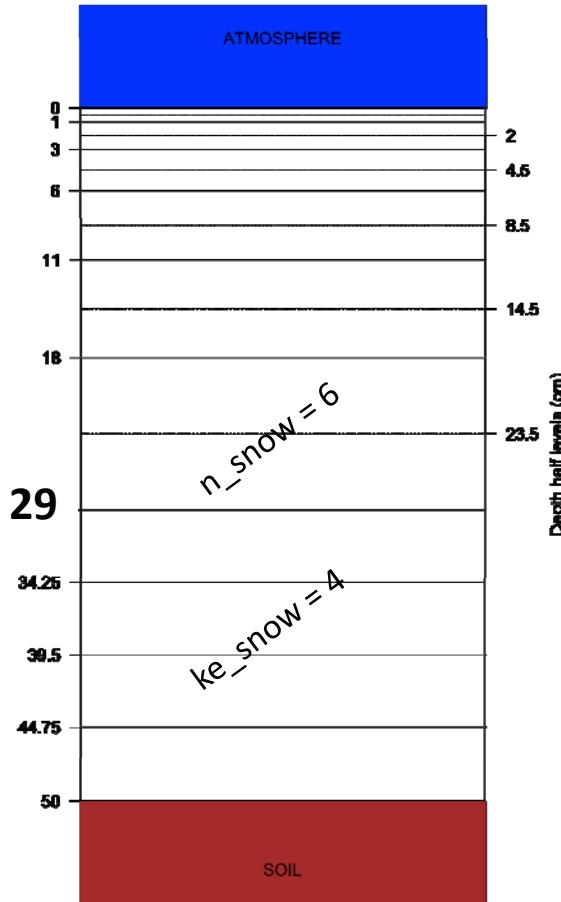
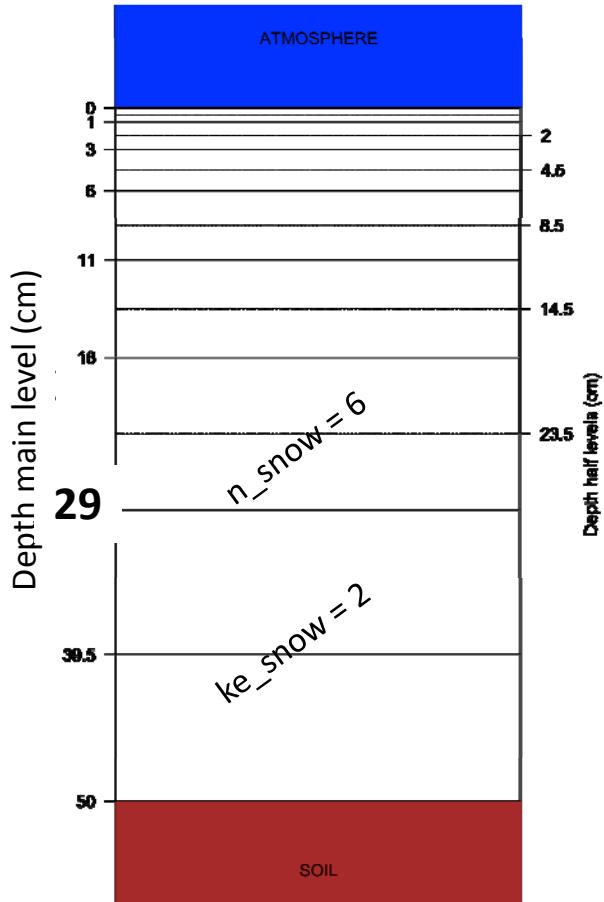
Experimental Site Weissfluhjoch, Davos, Switzerland



Phase II: Snow cover scheme (MLS) – Schematic



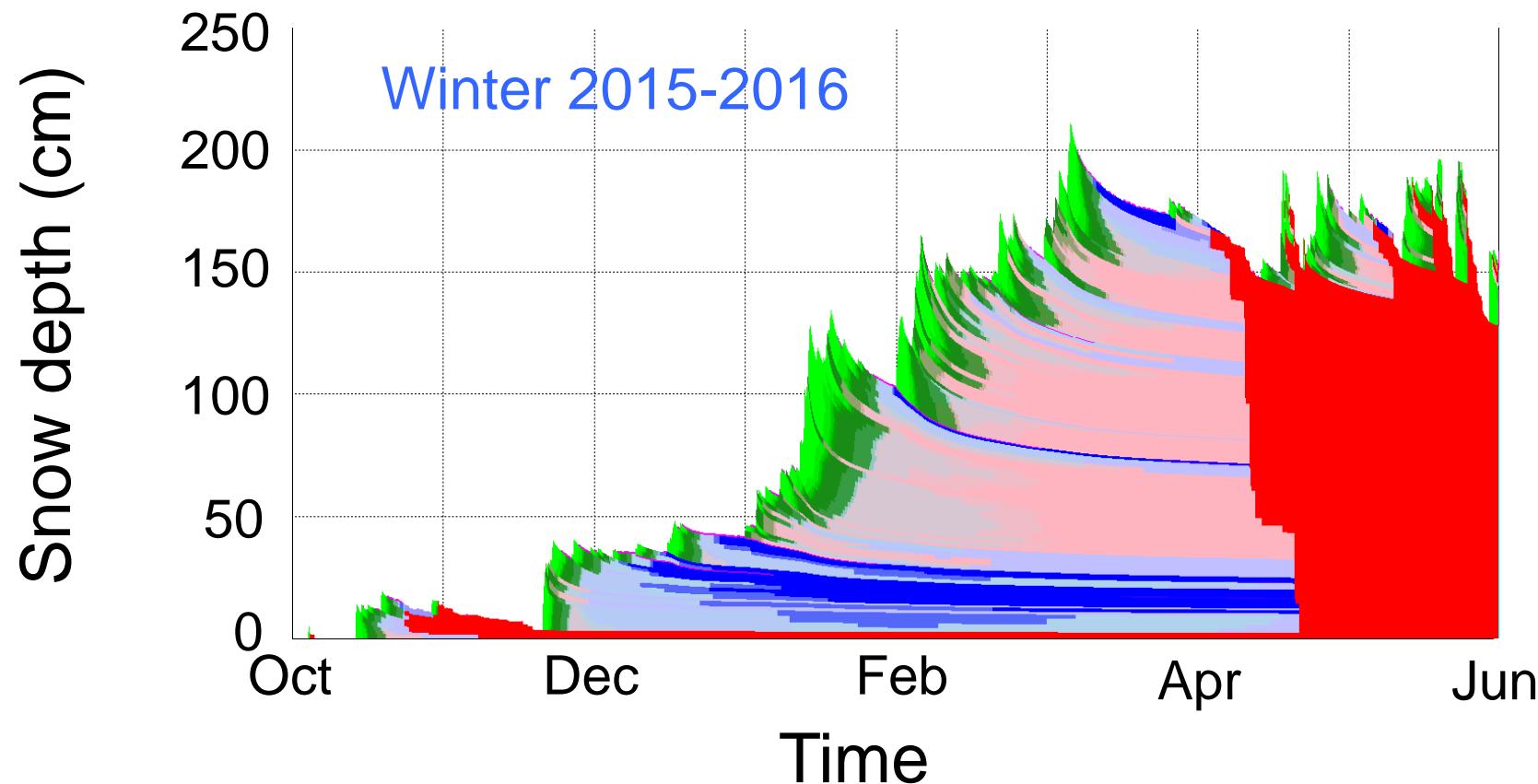
Phase II: Snow cover scheme schematic (MLS) – SAINT



Phase II: Snow cover scheme (MLS) – Implementation

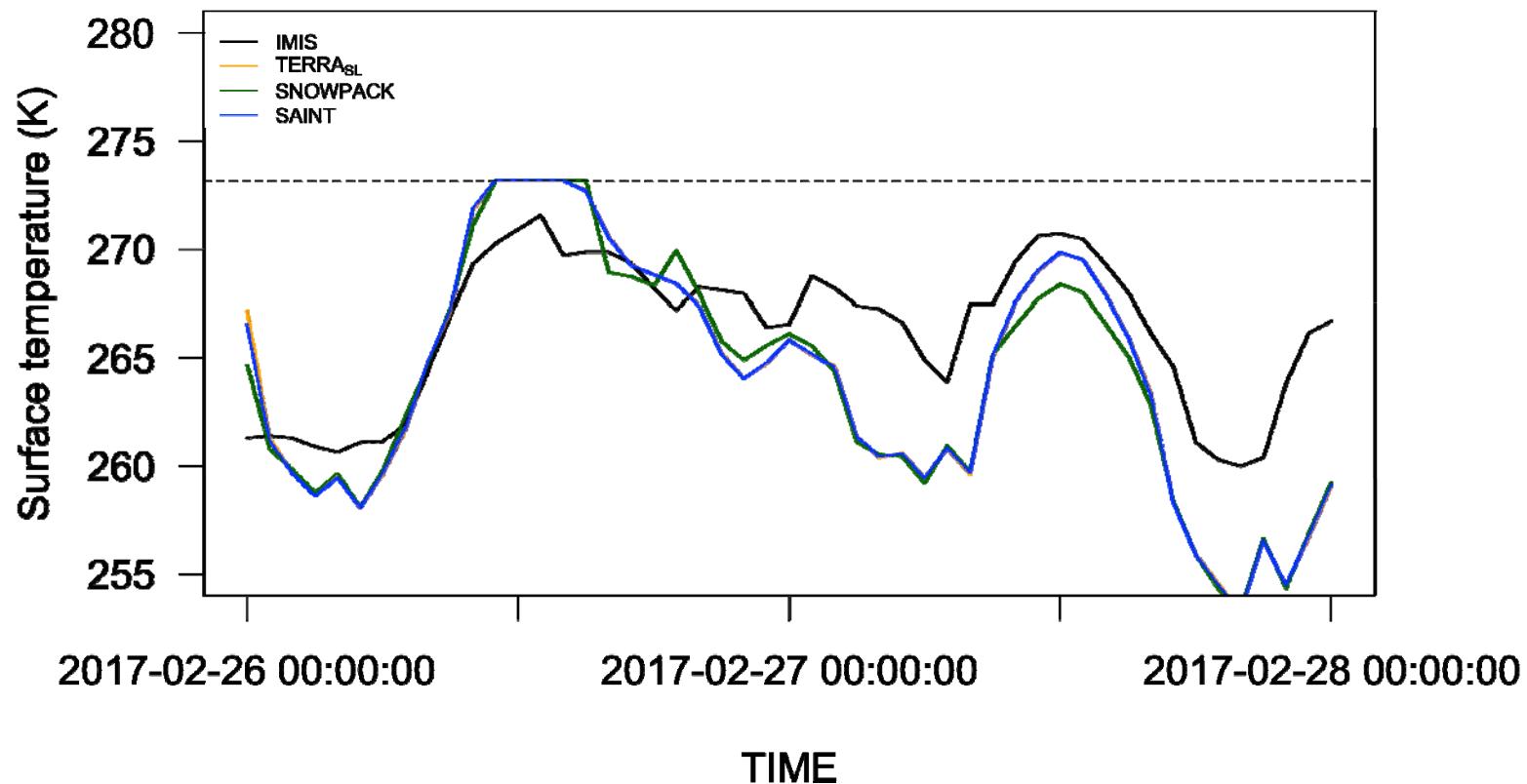
```
39 !!=====
40
41 MODULE sfc_terra
42
43 !=====
44 !SB <
45 USE sfc_snow_
46 !SB >
47
48 !=====
49 !SB <
50 ! Call subr
51 !SB >
52 !=====
53 !S
5430 ! Call subr
5431 CALL snow_or
5432 !-
5433
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5429
5430 ! Call subr
5431 CALL snow_or
5432 !-
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561 ! -----
562 ! Setup tridiagonal matrix for set of linear equations for each layer ...
563 ! -----
564
565 ! ... TOP LAYER
566 dz_low = zm_sn_so(2) - zm_sn_so(1)
567
568 a(1) = 0.0_wp
569 b(1) = 1 + (1 - cn) * alpha(1) * hcon_sn_so(1)/dz_low - alpha(1)*dlw_u_sn
570 c(1) = - (1 - cn) * alpha(1) * hcon_sn_so(1)/dz_low
571
572 d(1) = t_sn_so(1) + alpha(1) * (for_sn - dlw_u_sn*t_sn_so(1) + cn*hdif_sn_so(1))
573
574 ! ... INNER LAYERS
575 DO i = 2, n-1, 1
576
577 dz_up = zm_sn_so(i) - zm_sn_so(i-1)
578 dz_low = zm_sn_so(i+1) - zm_sn_so(i)
579 a(i) = - (1 - cn) * alpha(i)
580 b(i) = 1 + (1 - cn) * alpha(i)
581 c(i) = - (1 - cn) * alpha(i)
582
583 d(i) = t_sn_so(i) + cn*alpha(i)
584
585 ENDDO
586
587 ! ... BOTTOM LAYER
588
589 dz_up = zm_sn_so(n) - zm_sn_so(n-1)
590
591 a(n) = - (1 - cn) * alpha(n) * h
592 b(n) = 1 + (1 - cn) * alpha(n) * h
593 c(n) = 0.0_wp
594
595 d(n) = t_sn_so(n) - cn*alpha(n-1)
596
597 ! -----
598 ! Solve the system - Thomas Algorithm
599 ! -----
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663 ! -----
664 ! -----
665 ! -----
666 ! Section VIII: Updating
667 ! -----
668 ! -----
669 ! -----
670 DO ksn = 1, n_snow+ke_snow
671
672 IF(ksn == 1) THEN
673
674 t_sn_now(ksn) = MAX( MIN(e(ksn),t0_melt), 240.0_wp)
675
676 ELSE
677
678 t_sn_now(ksn) = e(ksn)
679
680 ENDIF
681
682 END DO
683
684
685 h_snow = h_snow
686
687 t_sn_sfc_now = MAX( MIN(e(1),t0_melt), 240.0_wp)
688
```

Phase II: Benchmark simulations – SNOWPACK vs. MSL



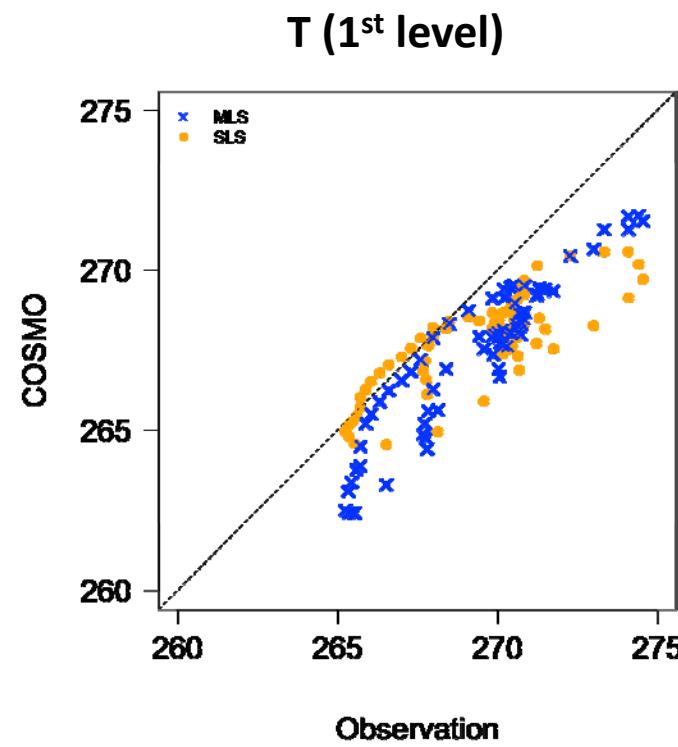
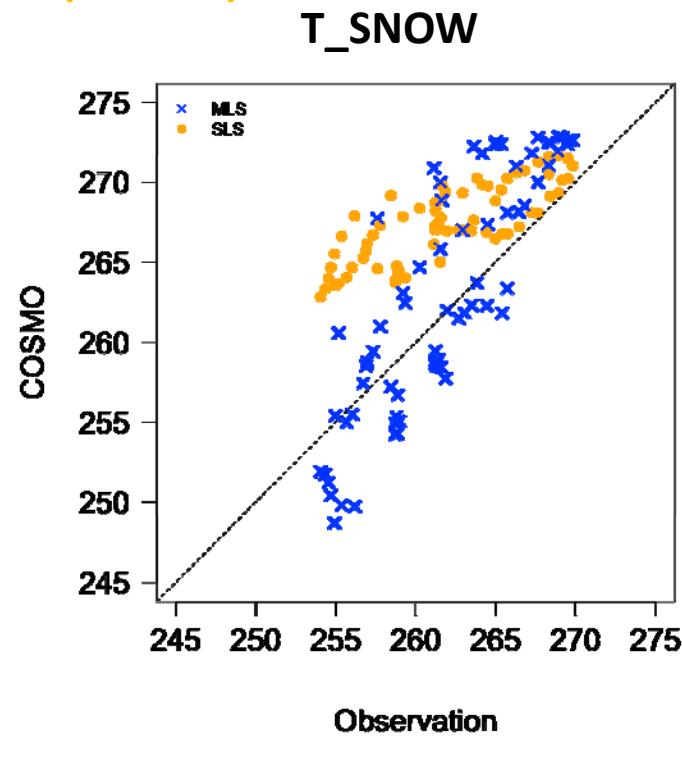
Phase II: Benchmark simulations – SNOWPACK vs. MSL

Weissfluhjoch



Phase II: Snow cover scheme (MLS) – Initial Results

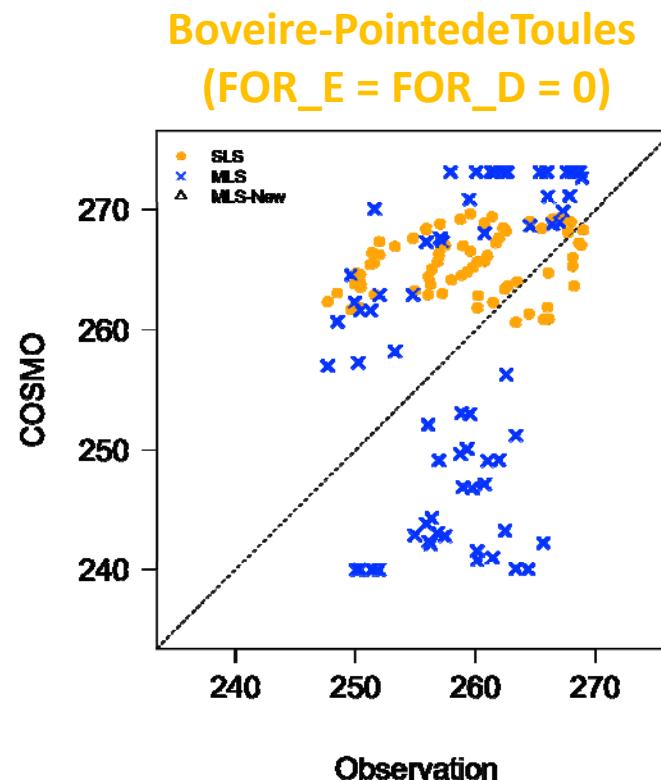
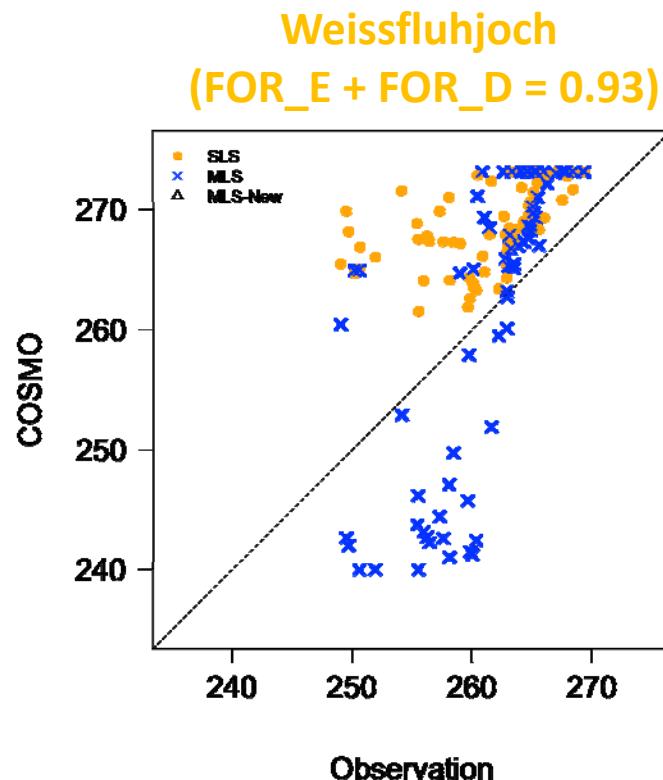
All IMIS (N = 112)



● = Single layer snow cover scheme (SLS)

✖ = Multi-layer snow cover scheme (MLS)

Phase II: Snow cover scheme (MLS) – Initial Results



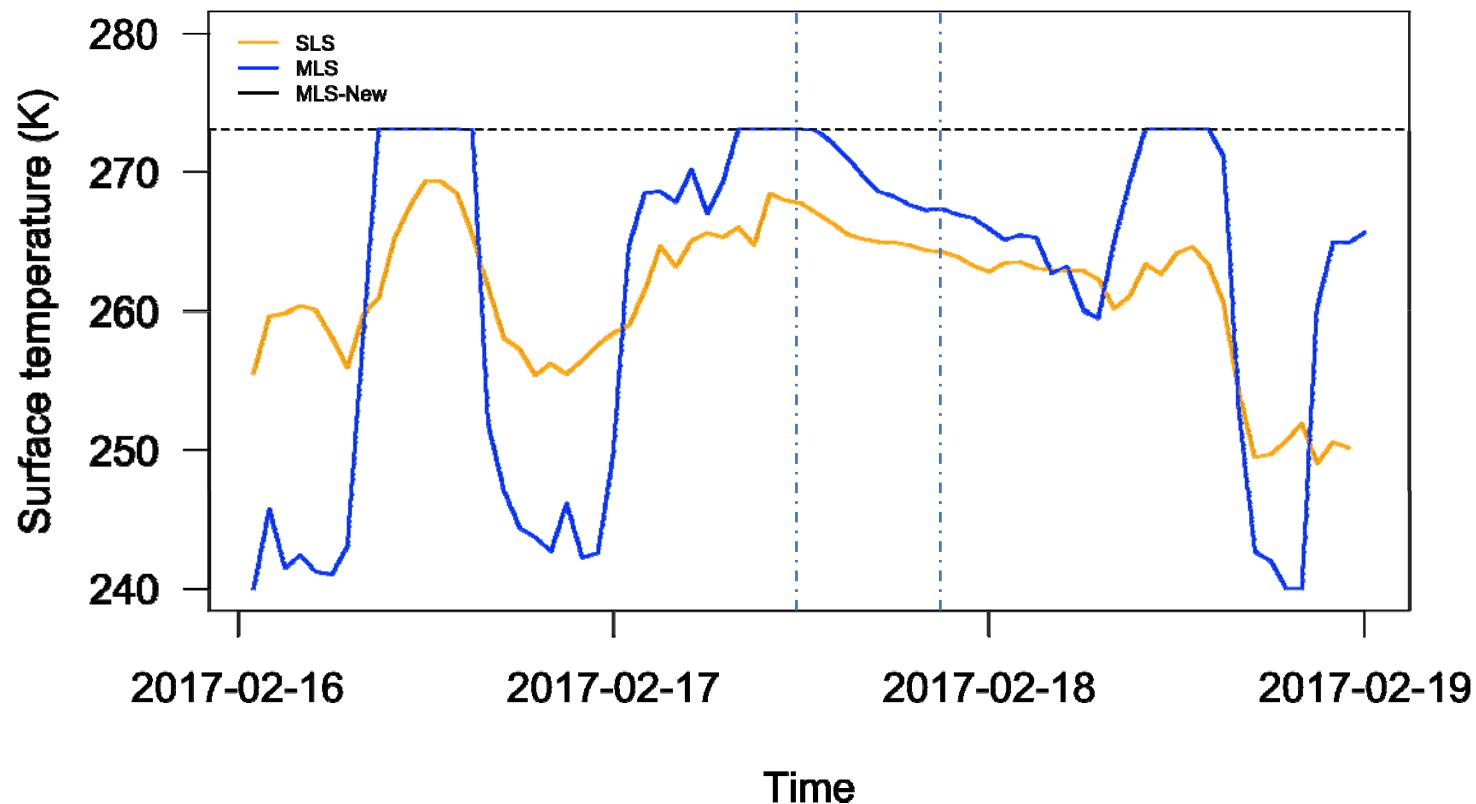
● = Single layer snow cover scheme (SLS)

✖ = Multi-layer snow cover scheme (MLS)

Phase II: Snow cover scheme - Meteorology

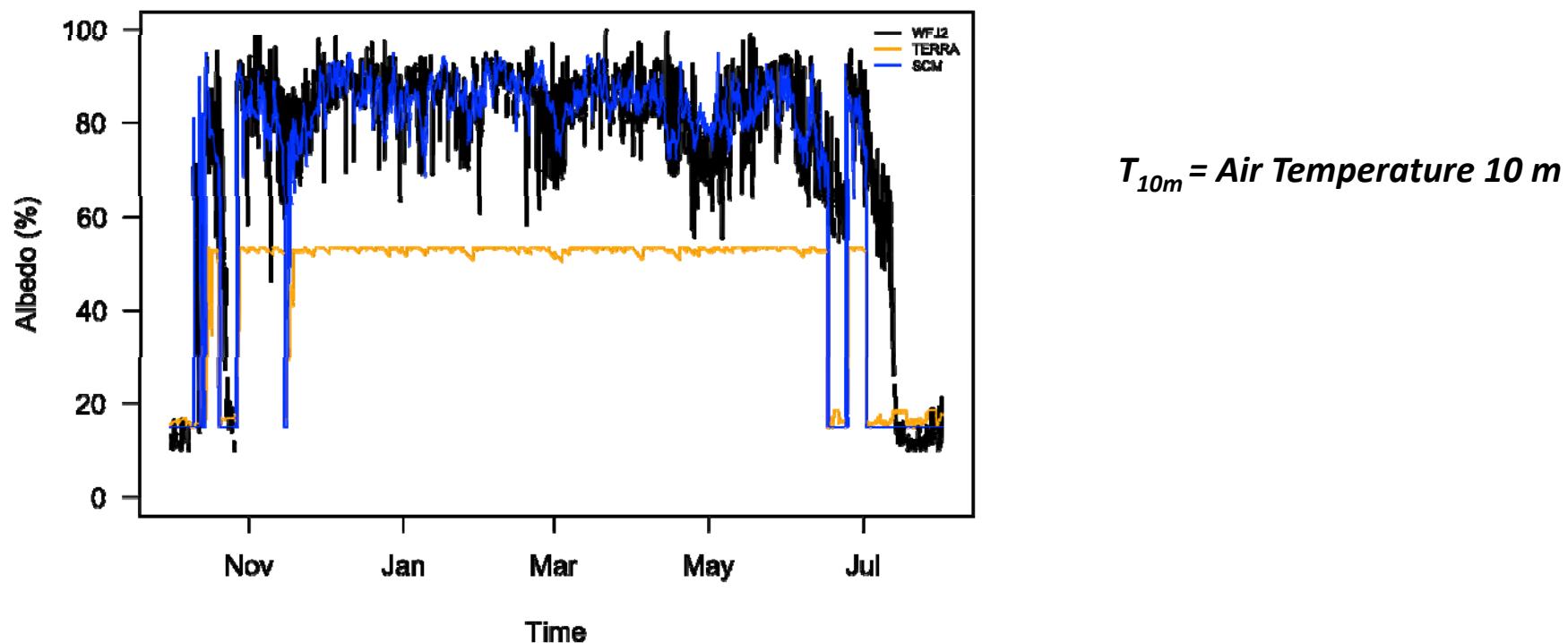
Weissfluhjoch

Snow



Phase II: Snow cover scheme (MLS) – Albedo

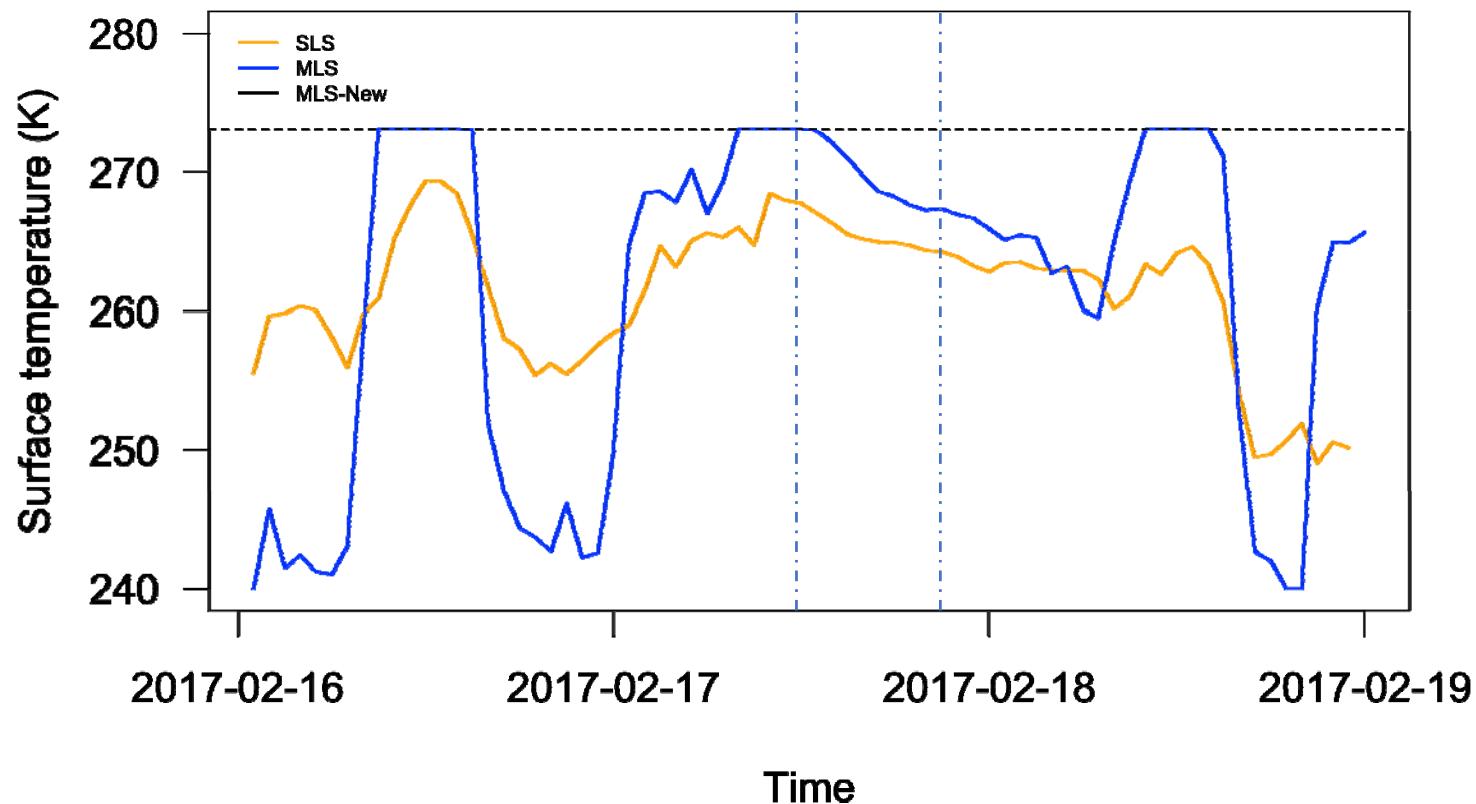
$$\alpha_{SCM} = a + b \times P_{rate} + c \times T_{SFC} - d \times T_{10m}$$



Phase II: Snow cover scheme - Meteorology

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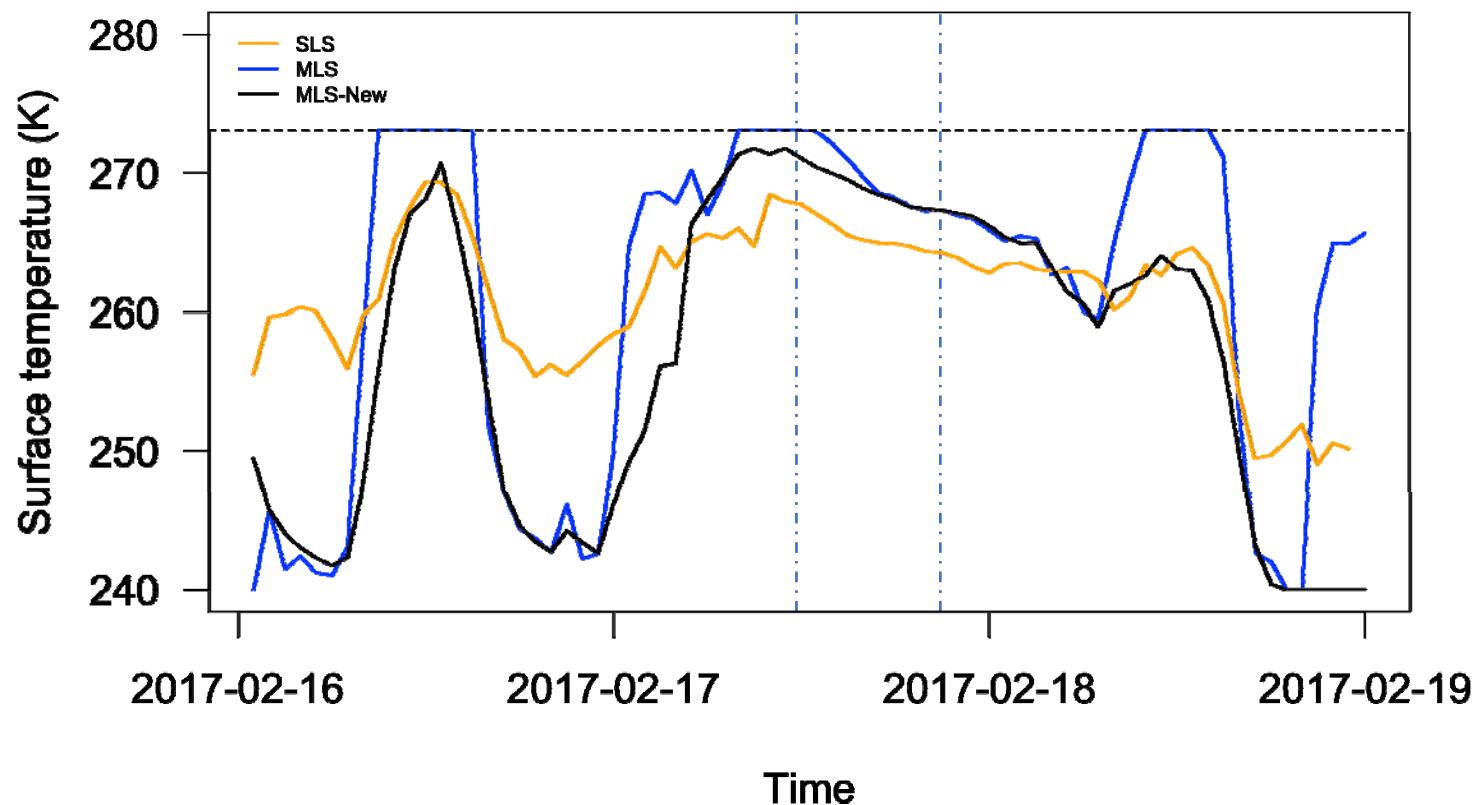
Snow



Phase II: Snow cover scheme - Meteorology

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Snow



Phase II: Snow cover scheme (MLS) – TCH

Boundary-Layer Meteorology
October 2017, Volume 165, Issue 1, pp 161–180 | Cite as

How do Stability Corrections Perform in the Stable Boundary Layer Over Snow?

Authors	Authors and affiliations
Sebastian Schlögl  , Michael Lehning, Kouichi Nishimura, Hendrik Huwald, Nicolas J. Cullen, Rebecca Mott	

Stability Corrections:

$$\psi_m(T, T_{sn}, \bar{U}) = a_1 B + b_1 S,$$

$$\psi_s(T, T_{sn}, \bar{U}) = a_2 B + b_2 S,$$

$$B = \Delta T / \bar{T}$$

$$S = z_{\text{ref}} g / \bar{U}^2$$

test site	a_1	b_1	test site	a_2	b_2
WFJ07 (3 m)	3.227	0.0043	WFJ07 (3 m)	-982.90	-0.0005
WFJ07 (5 m)	-4.441	0.0025	WFJ07 (5 m)	-642.51	0.0009
WFJ11	-30.74	0.0008	WFJ11	-1135.4	-0.0015
PM07 NWW	-191.93	0.0008	PM07 NWW	-751.73	-0.0005
PM07 SEE	-29.55	0.0090	PM07 SEE	-692.74	-0.0123
GRoo (1 m)	-145.41	-0.0914	GRoo (1 m)	-378.92	-2.0489
GRoo (2 m)	-179.56	-0.0369	GRoo (2 m)	-243.93	-0.7448
Universal	-65.35	0.0017	Universal	-813.21	-0.0014

Sensible heat flux:

$$H = \rho c_p C_H \bar{U} \Delta \theta,$$

Transfer Coefficient:

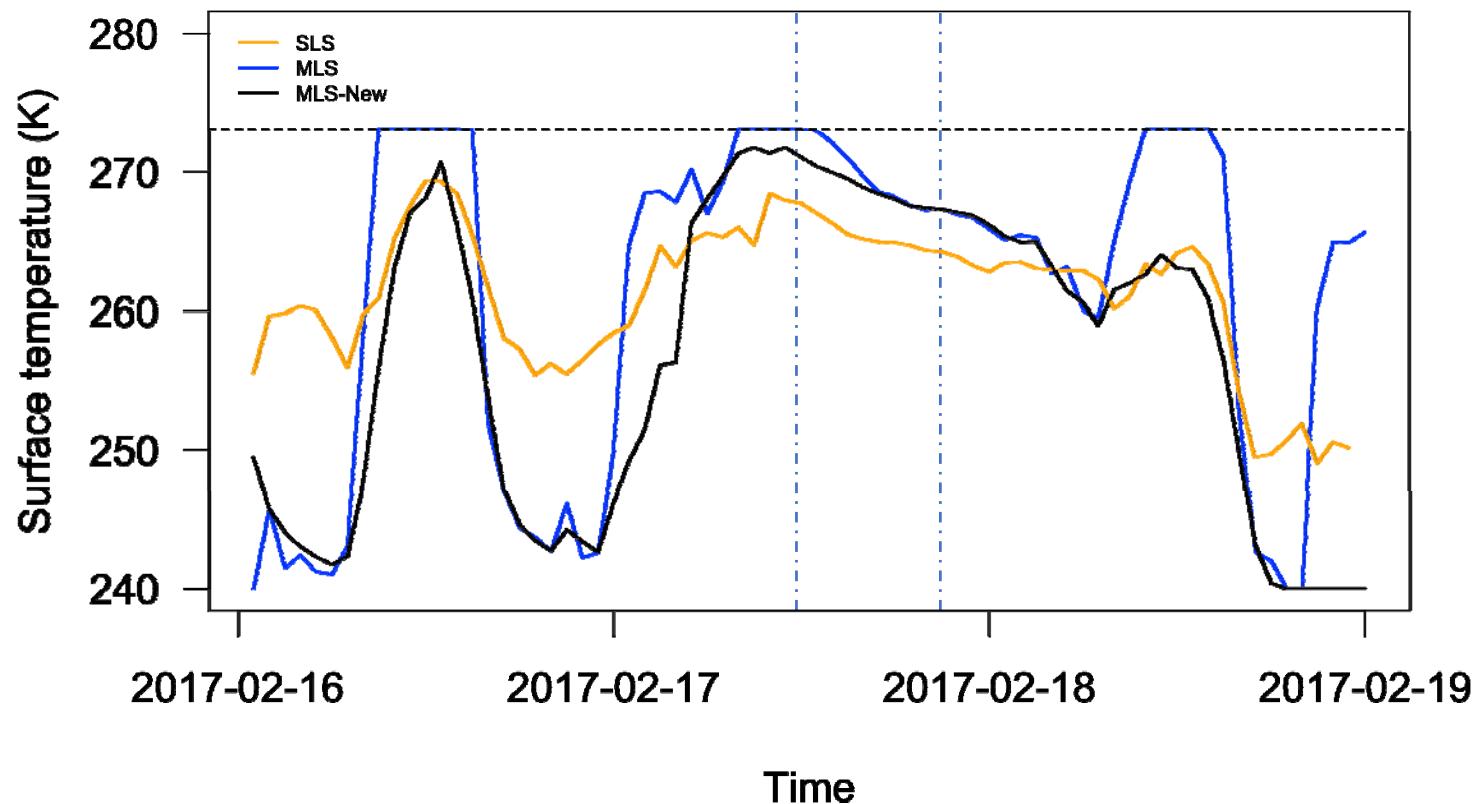
$$C_H = \frac{k^2}{\left[\ln\left(\frac{z_{\text{ref}}}{z_{0M}}\right) - \psi_m(\zeta) \right] \left[\ln\left(\frac{z_{\text{ref}}}{z_{0M}}\right) - \psi_s(\zeta) \right]},$$

where $k = 0.4$ is the von Kármán constant, $\zeta = (-k z_{\text{ref}} g T_*) / (\theta_s u_*^2)$ is the modelled stability parameter (stability parameter henceforth), $u_* = k \bar{U} (\ln(z_{\text{ref}}/z_{0M}) - \psi_m)^{-1}$ is the modelled friction velocity, $T_* = k (\theta_s - \theta_{z_{\text{ref}}}) (\ln(z_{\text{ref}}/z_{0M}) - \psi_s)^{-1}$ is the modelled temperature scale, z_{0M} is the aerodynamic roughness length and ψ_m and ψ_s are the stability corrections for momentum and scalars. In our analysis, we used the simple approach that the roughness

Phase II: Snow cover scheme - Meteorology

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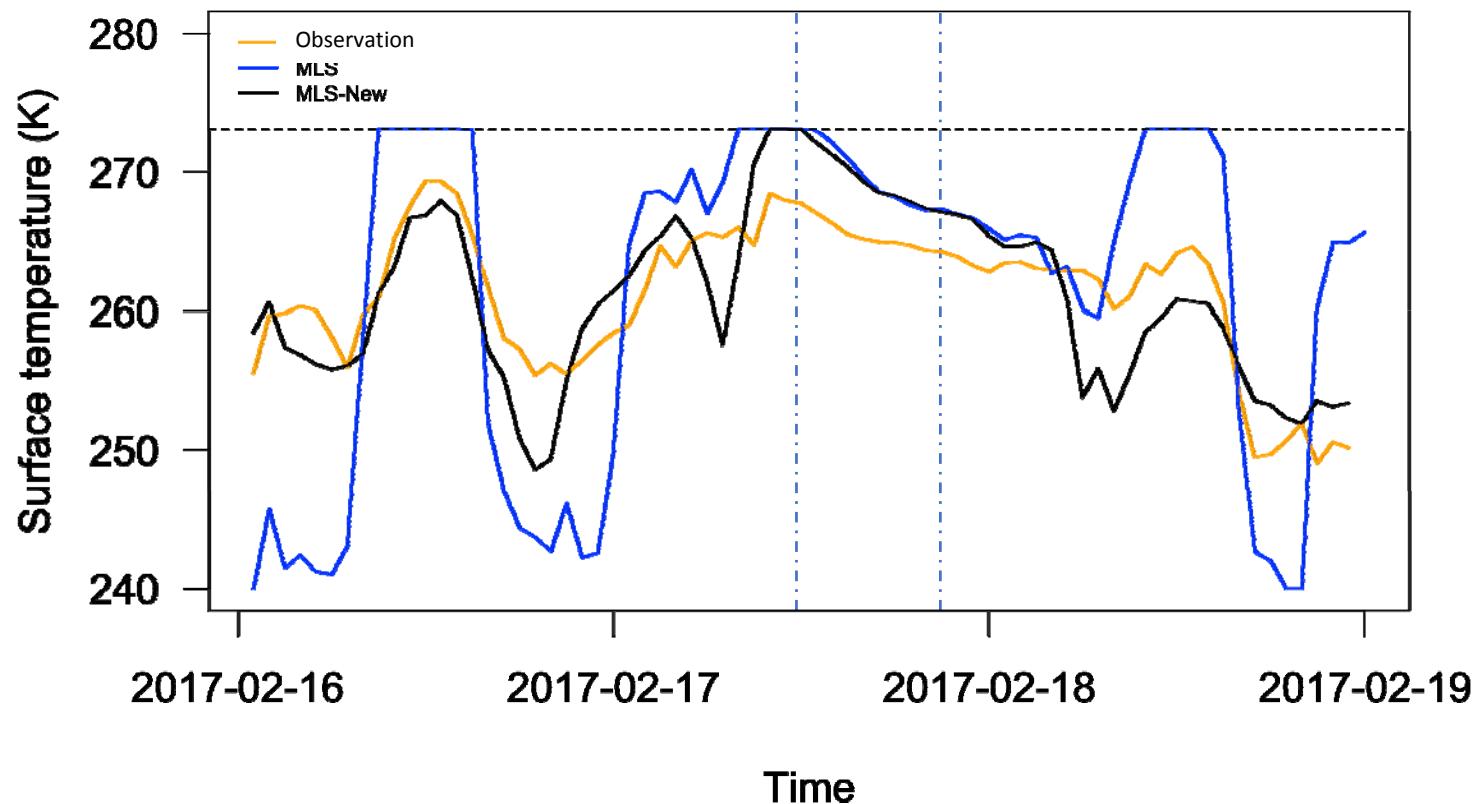
Snow



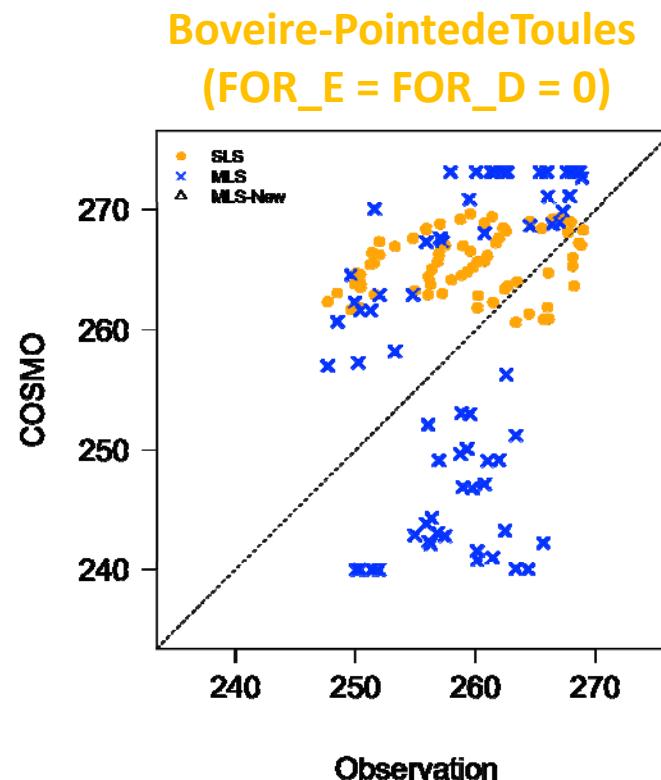
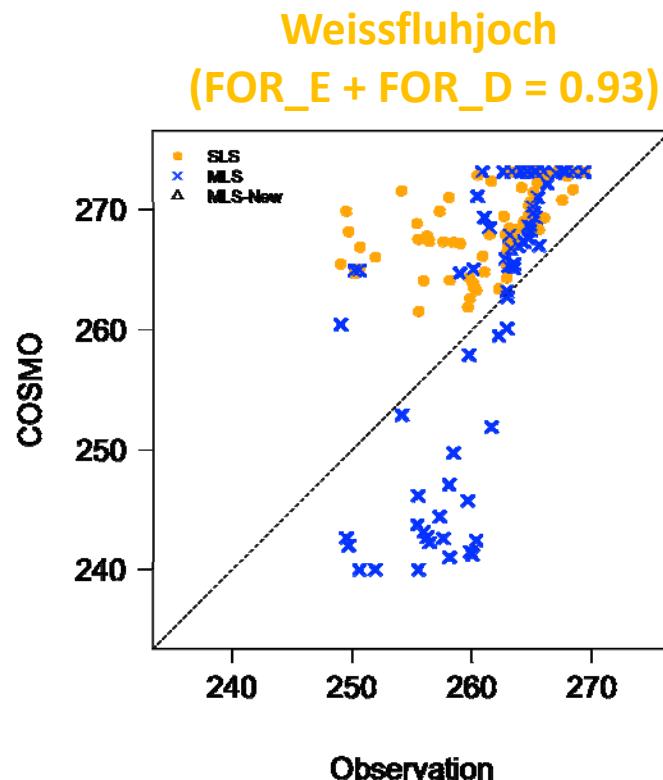
Phase II: Snow cover scheme - Meteorology

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Snow



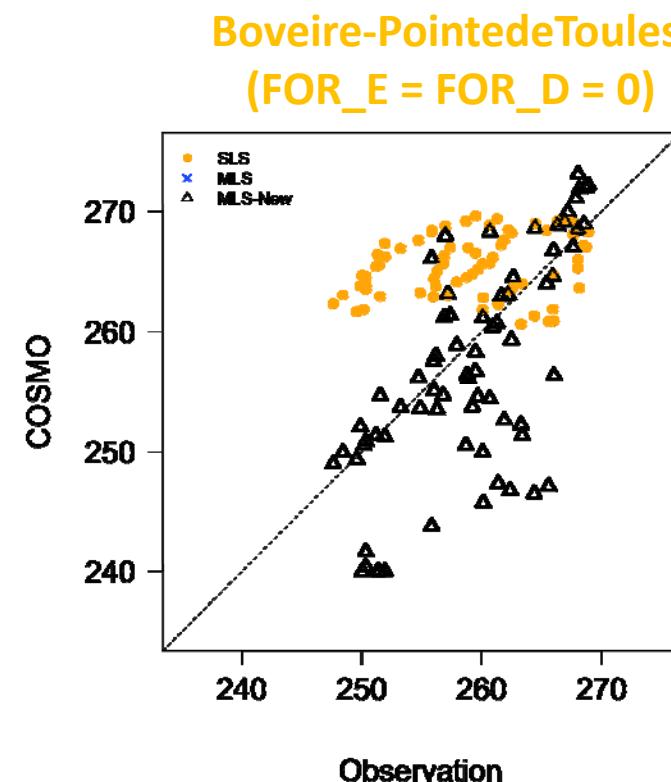
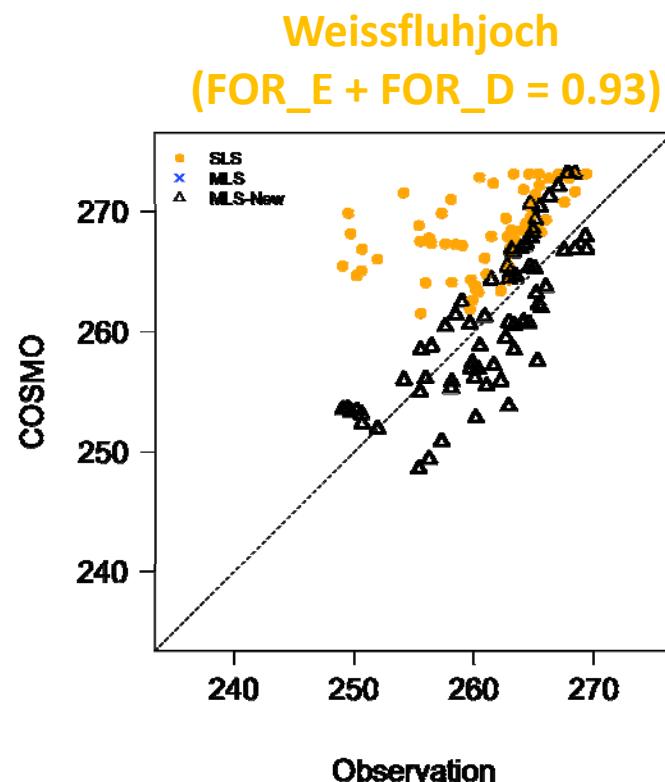
Phase II: Snow cover scheme (MLS) – Initial Results



● = Single layer snow cover scheme (SLS)

✖ = Multi-layer snow cover scheme (MLS)

Phase II: Snow cover scheme (MLS) – Initial Results



Δ = Multi-layer snow cover scheme (MLS) with new flux parameterizations

Phase II: Snow cover scheme (MLS) - Status

- **Code development based on COSMO version 5.04h including 'latest' developments by Matthias R.**
- **Call of a subroutine (*snow_on_soil*) in TERRA, which ...**
 - ... **uses a fixed number of snow layers (*n_snow* = 6) and an additional number of layers set by namelist (default *ke_snow* = 2)**
 - ... **solves the heat equation for the whole column (snow + soil = 16 layers)**
 - ... **calculates a snow specific atmospheric forcing (new albedo parameterization & transfer coefficients required).**
 - ... water transport (bucket) through snow column only (INTENT(OUT) > soil, runoff, storage(?))
 - ... settling, absorption solar energy, dust on snow, snow deposition ...
- **Call of subroutine (*snow_on_xxx*) ... (*snow_in_xxx*) ...**

Phase II: Snow cover scheme (MLS) – Q & A

- We need to calculate the fluxes in a different way! Problem?
- How to tackle the problem of different forcing for snow covered fractions?
 - Tile approach?
 - Updating T_SNOW only?
 - Other option?
- **Should we switch to ICON (ASAP)???**
- Are we missing something?
- What other problems might be solved using MLS, but need investigation?
 - Snow data assimilation with SMRT?
 - Compatibility with snow analysis?
 -



Thanks! Questions and/or comments?

Sascha Bellaire¹, Michael Lehning^{1,2}, Jean-Marie Bettems³, Oliver Fuhrer³

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