



User Seminar 2011, COLOBOC Workshop



# Development of snow parameterization: definition of initial snow density and accounting of fractional snow coverage within a cell

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## Methods and data

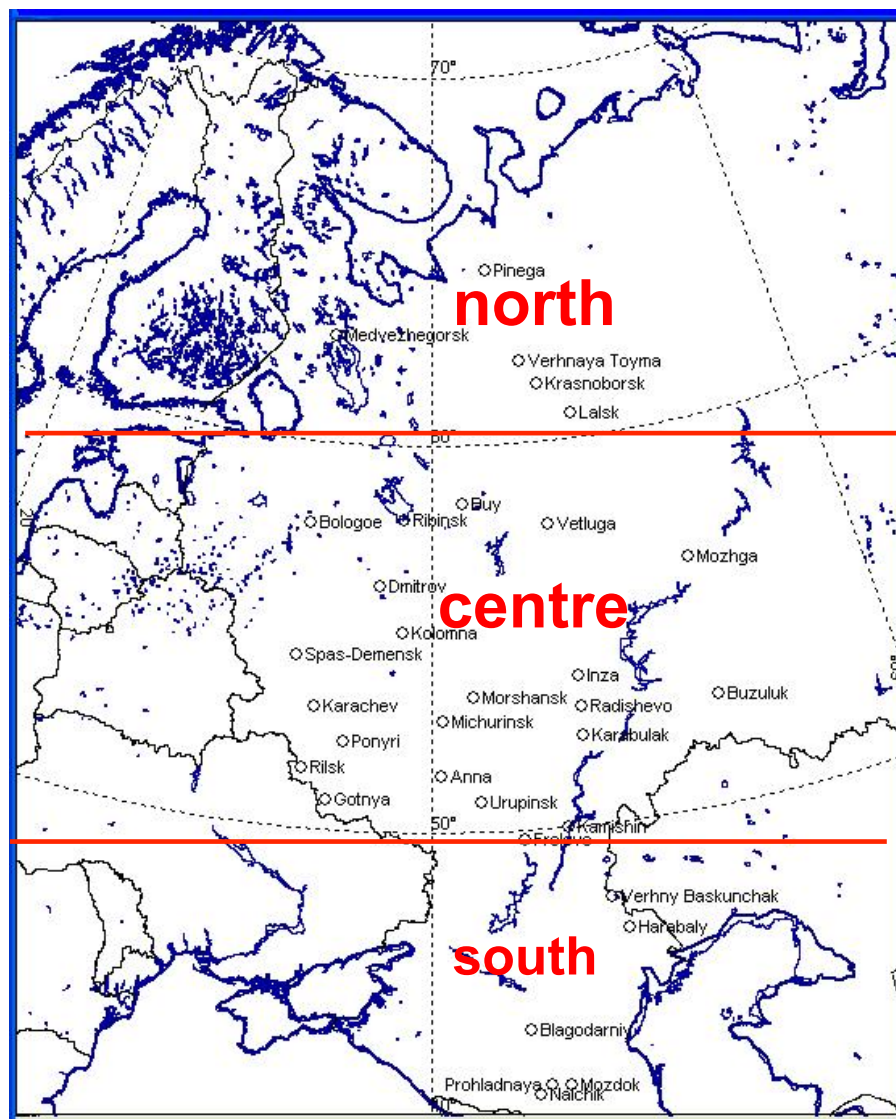
- model COSMO-Ru 14km resolution, version 4.2
- Integration period – 78h (from 00 UTC)
- Data:
  - station SYNOP measurements (36 stations)
  - decade measurements of snow survey on Roshydromet's stations (33 stations)







## Area of study: European Part of Russia







**Problem:**

**an overestimation (twice as much) of initial snow water equivalent (SWE) from GME (DWD)**



**too much snow mass in the model in comparison with the reality**



**underestimation of T2m forecasts (up to 10°C) during snowmelt period**



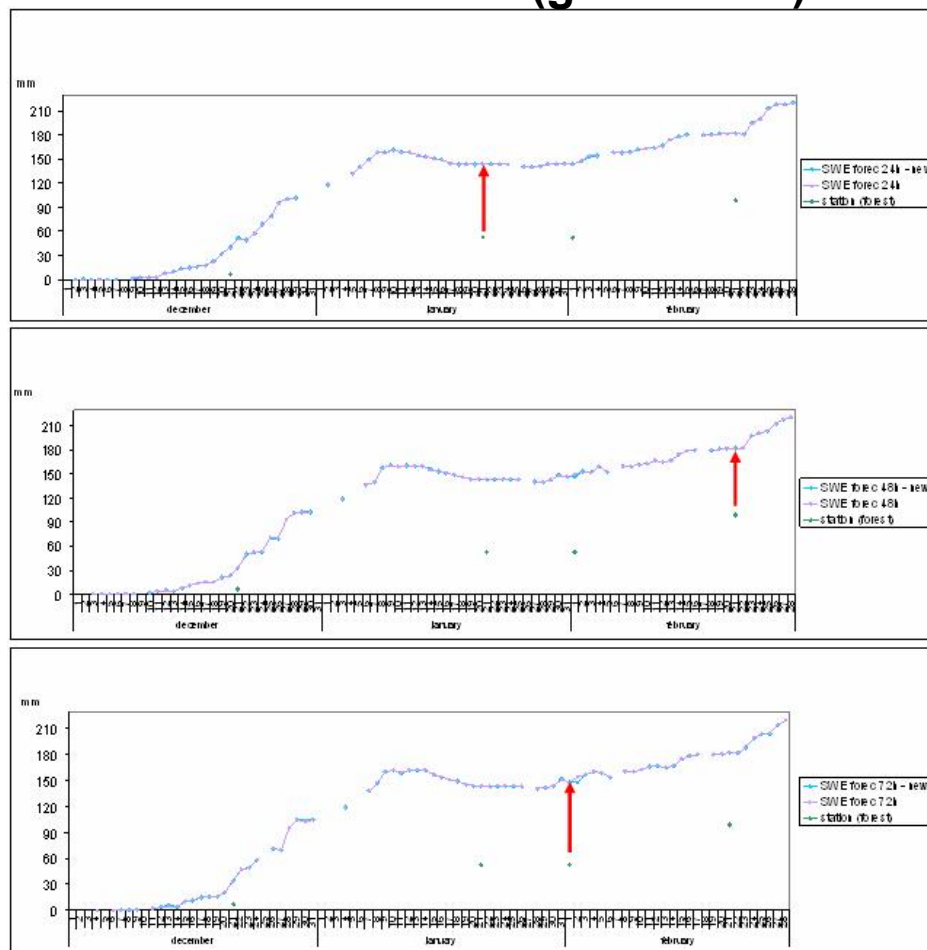




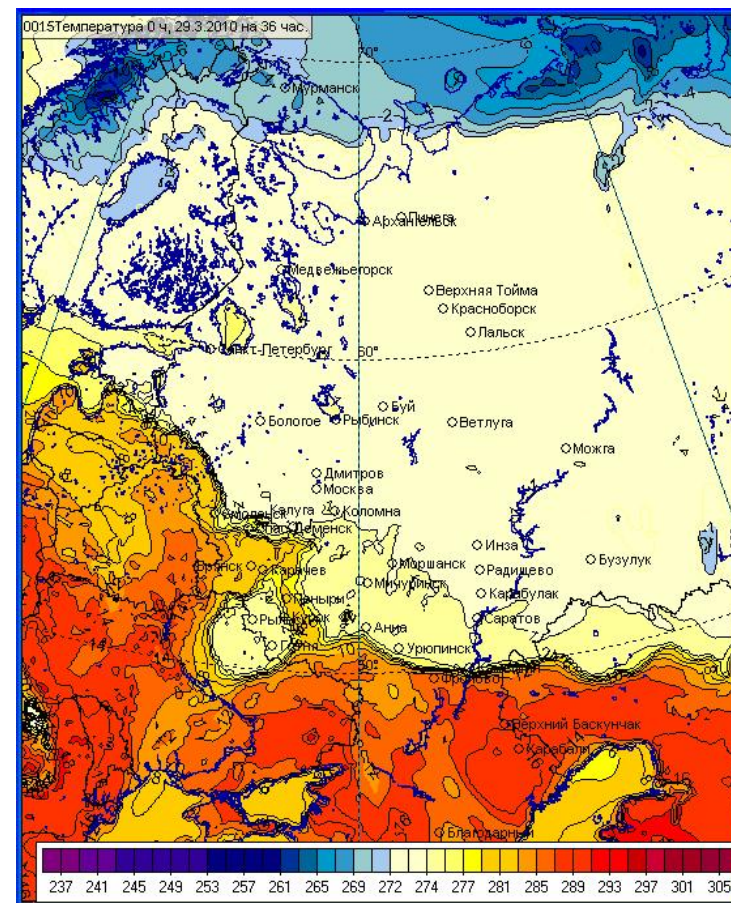
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## SWE forecasts and station measurements (green dots)



## T2m, 36h forecast. 29 March 2010



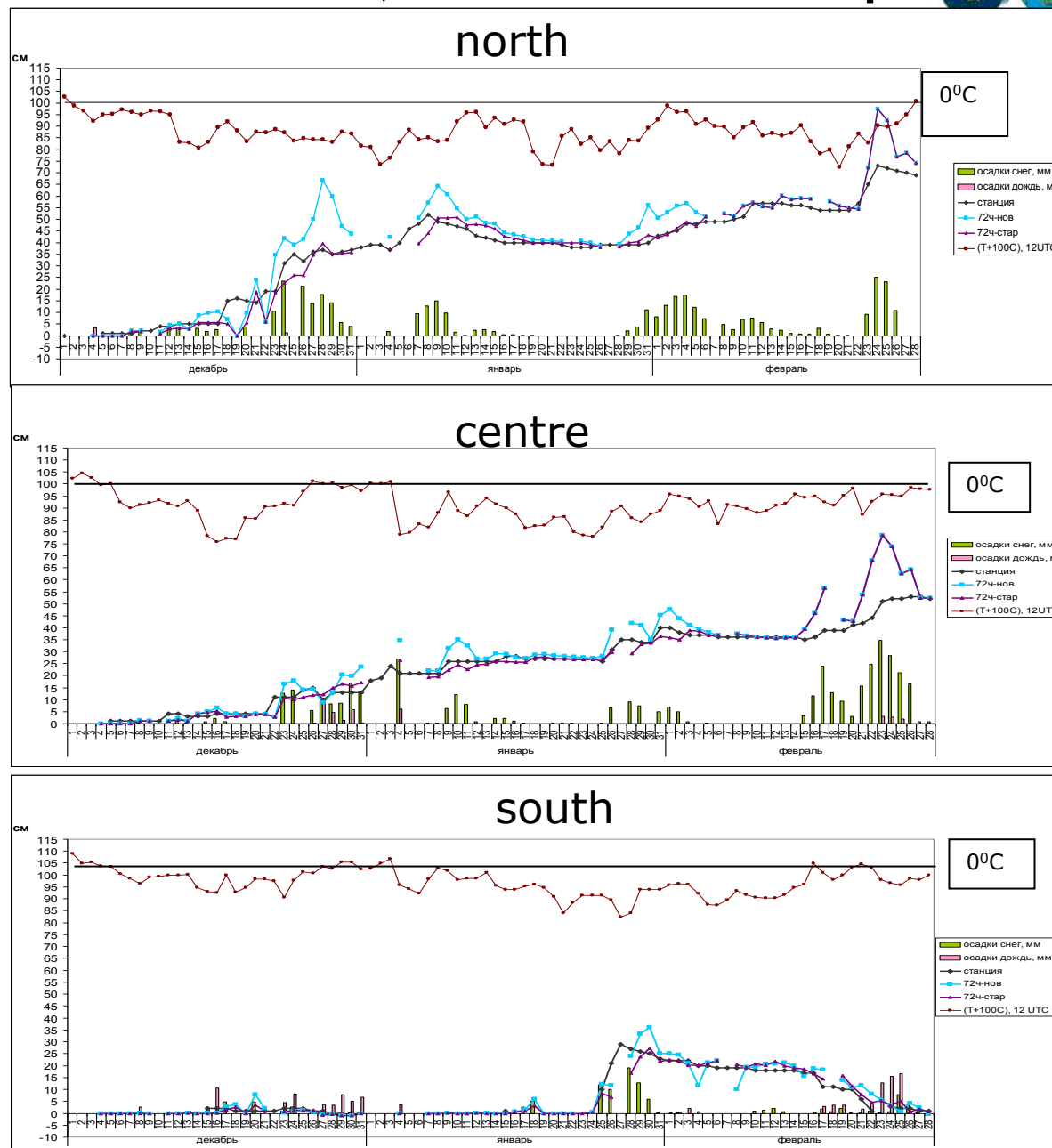




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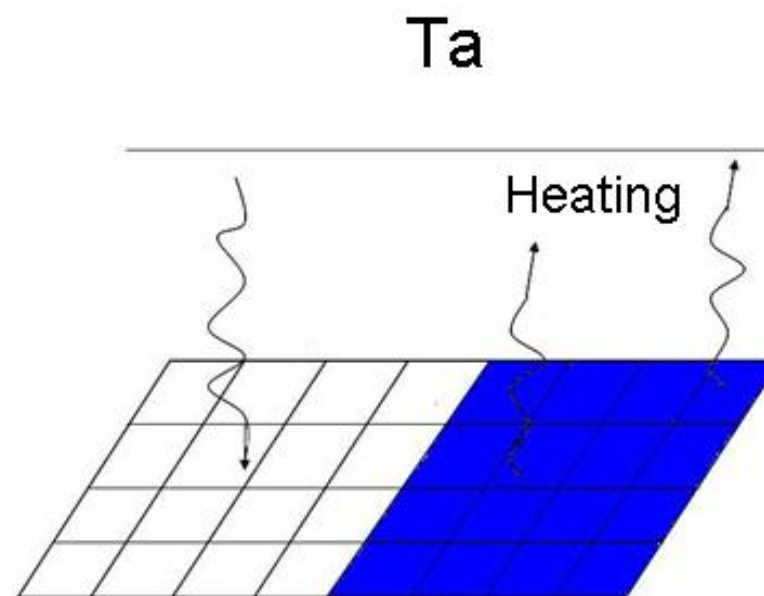
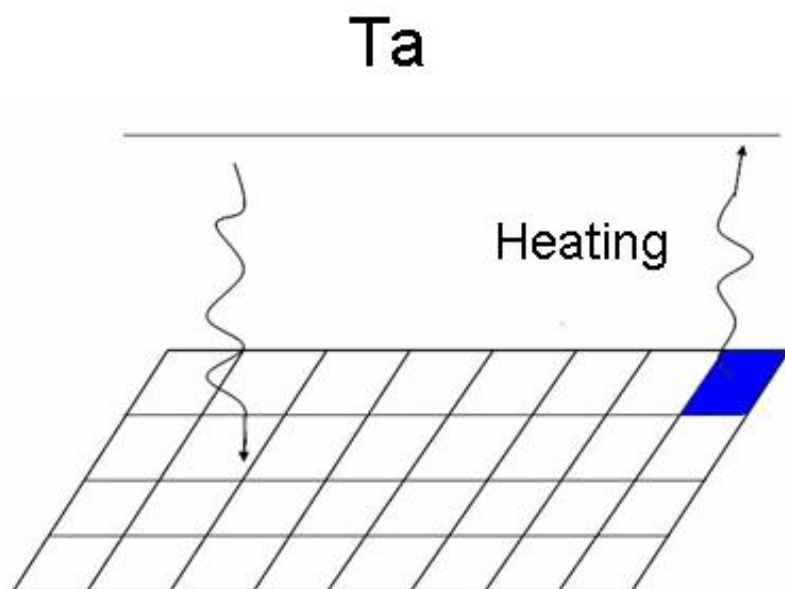
**Hsnow**







# Cell heating







# Surface fraction covered by snow (in TERRA within COSMO-Ru)

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, \frac{W_{snow}}{\delta_s}))$$

$\delta_s = 0.015$  parameter

$W_{snow}$  - snow water equivalent (SWE)







## Possibilities for improving T2m during snowmelt period:

- To change SWE (there is a feedback in T2m)

Thus it's necessary to define snow density ( $\rho_{\text{snow}}$ ) correctly (since prognostic snow height corresponds to measurements, and  $\text{SWE} = H_{\text{snow}} * \rho_{\text{snow}}$ )

- To take into account snow fractional covering







## Different approaches for determination snow density:

- As constant (for example,  $250 \text{ kg/m}^3$ )
- Precipitation microphysics
- Laboratorial researches
- “Hydrological-historical” approach: snow density is a function of meteorological parameters during the whole snow period (first of all – temperature)



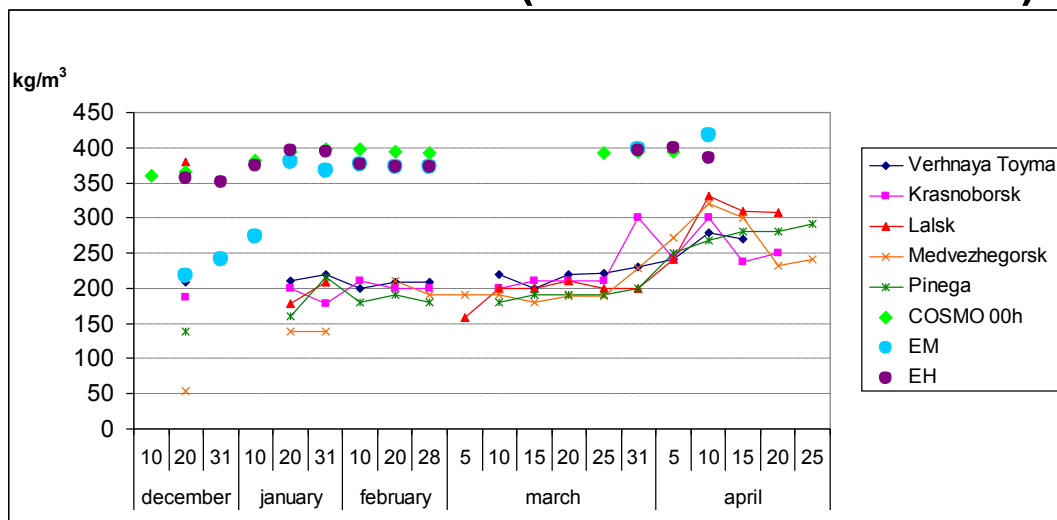




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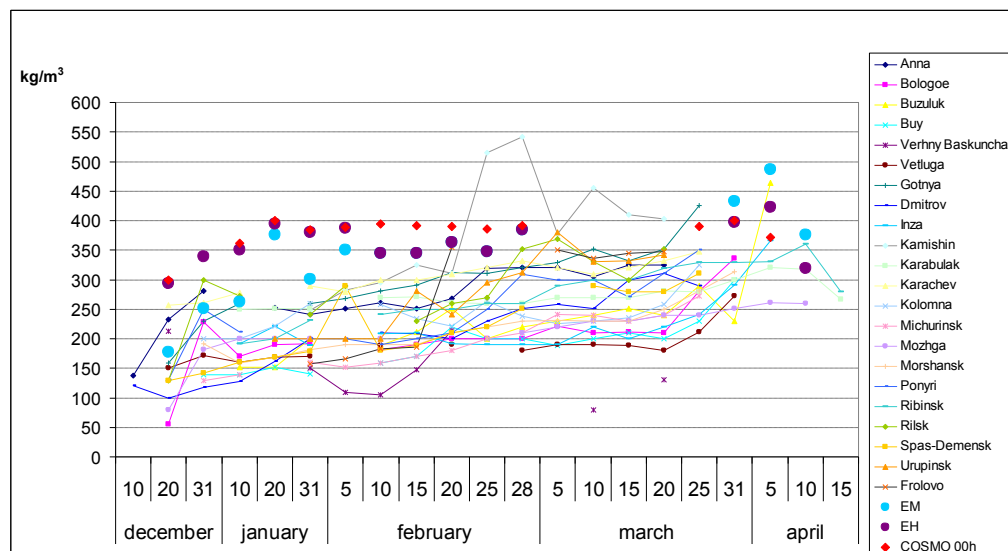


### Snow density ( $\text{kg/m}^3$ ): initial model values for 00h, 72h forecasts (two snow schemes) and station measurements



→ North region

Centre region



Initial and following snow density overestimation of the model COSMO-Ru

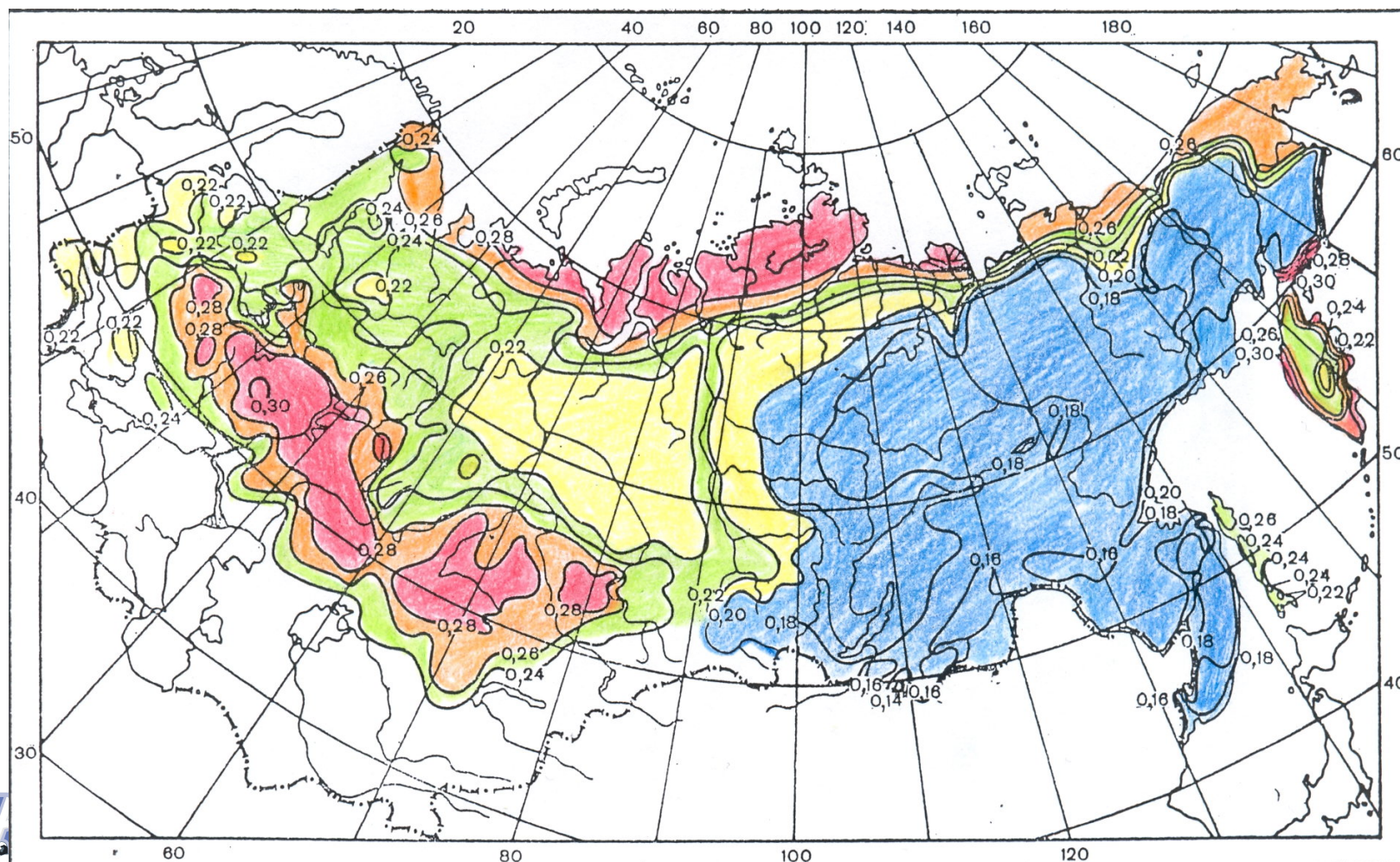




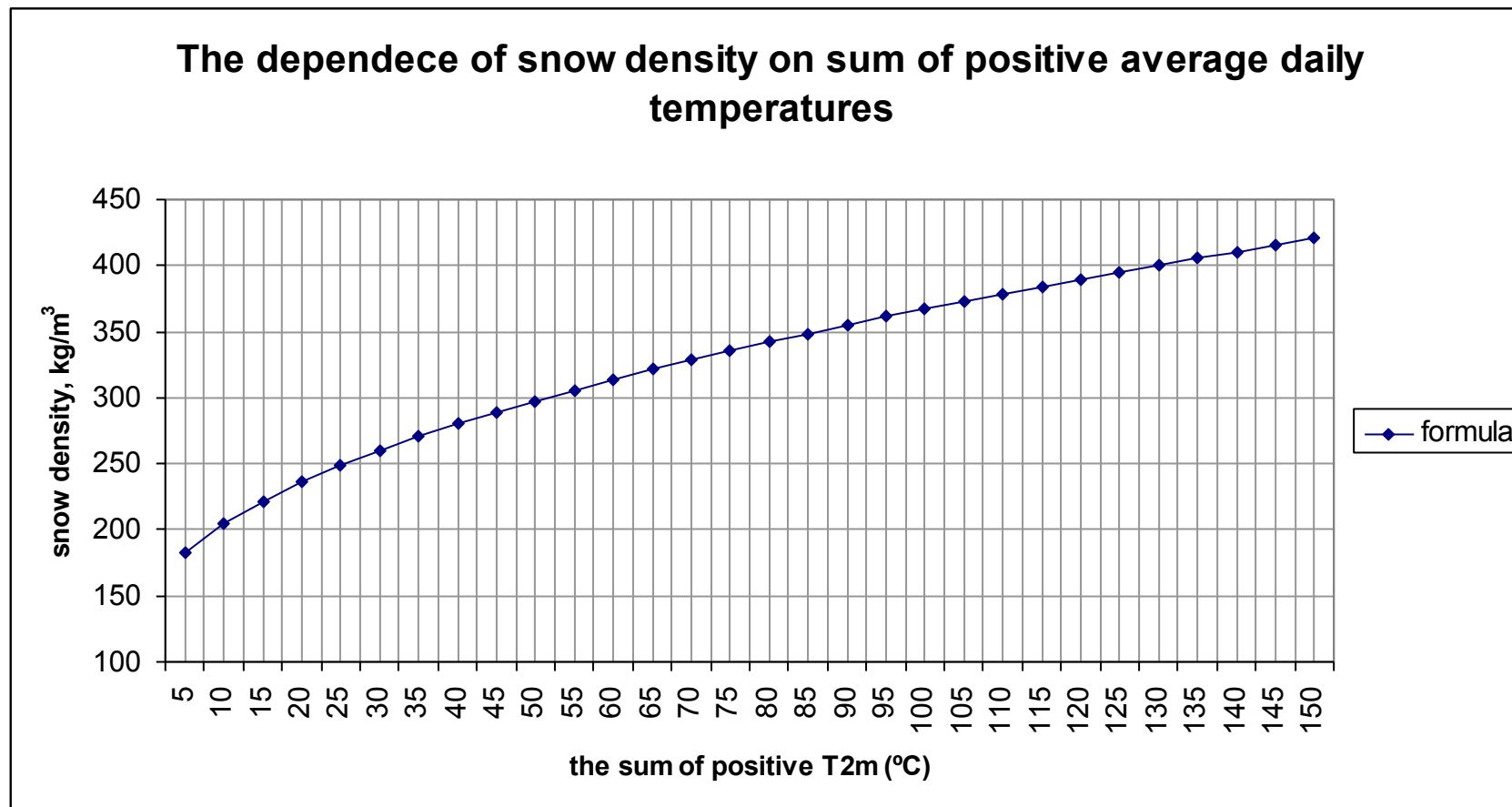
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# Average snow density ( $\text{g/cm}^3$ ) at maximum snow height







$$\rho = As + 23,72 \cdot \sqrt{SumT} \quad \text{'temperature' method}$$



As – empirical coefficient, As=130

**See *Dmitrieva N.* Snow density calculation using meteorological data // Meteorology and Hydrology, 1950, №2, pp.39-44**





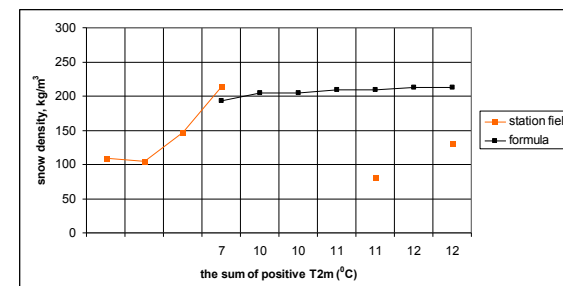
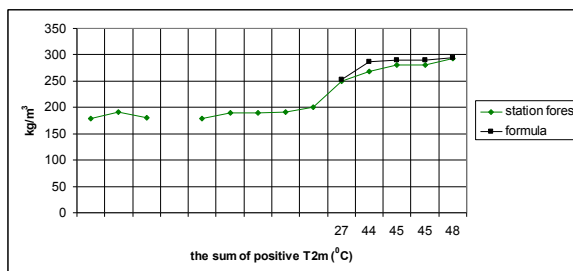
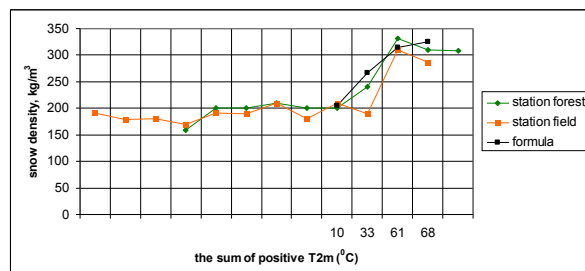
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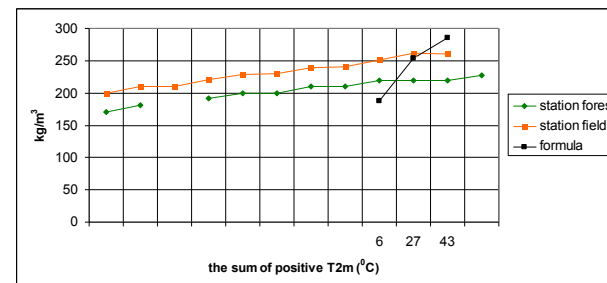
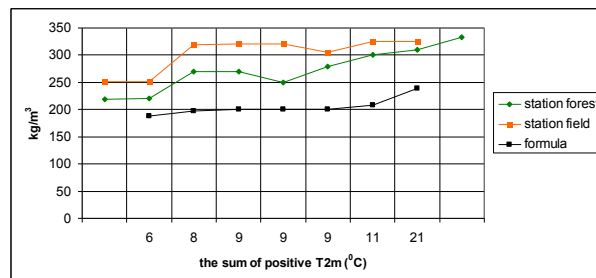
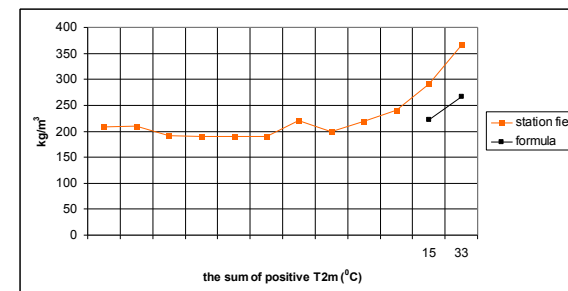
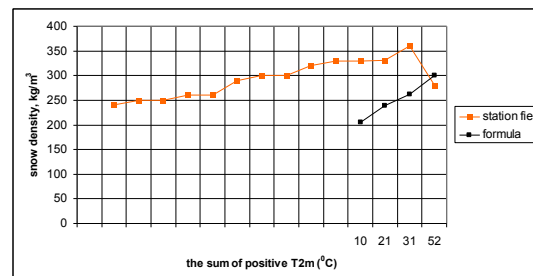
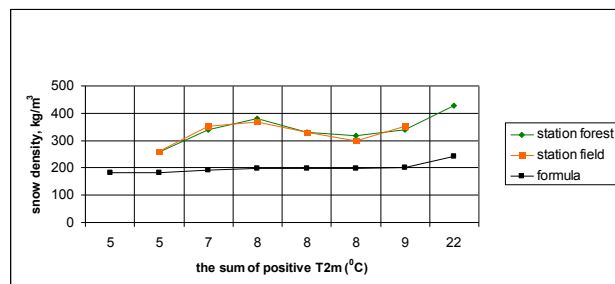
## Snow density, $\text{kg/m}^3$ : measurement and formula calculations

### North region

### South region



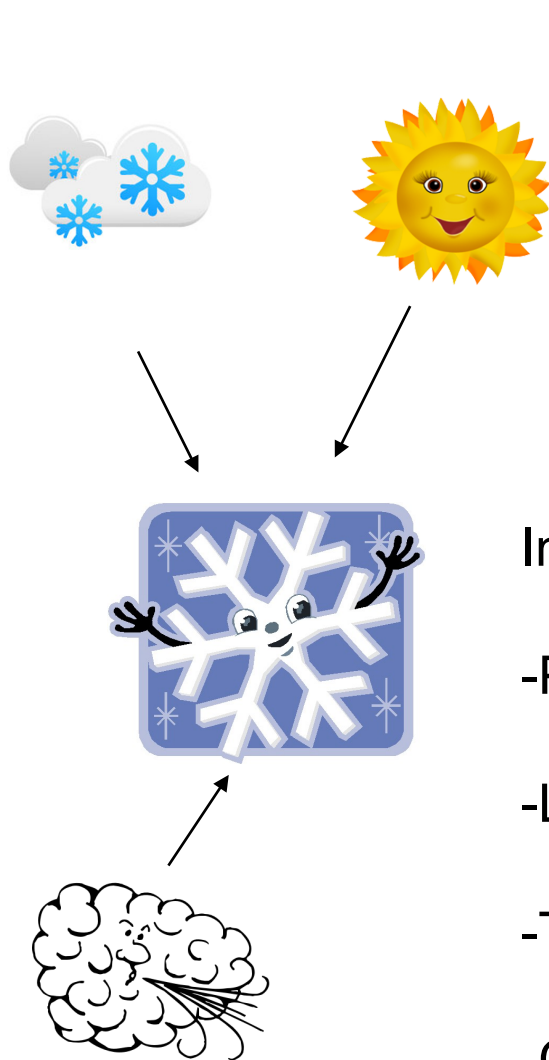
### Centre region





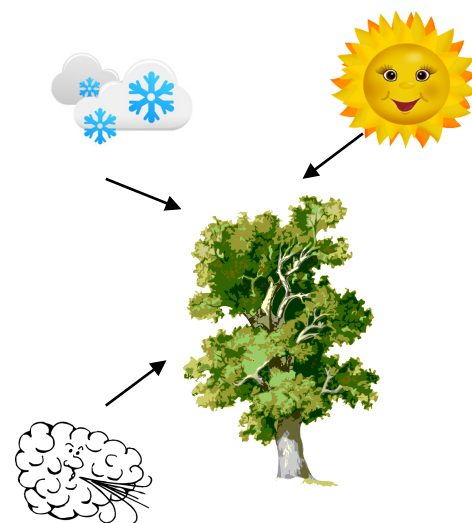
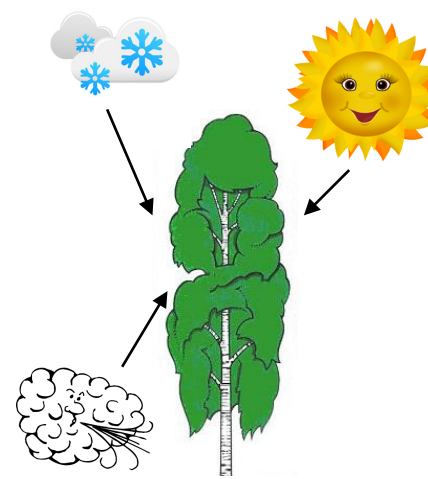
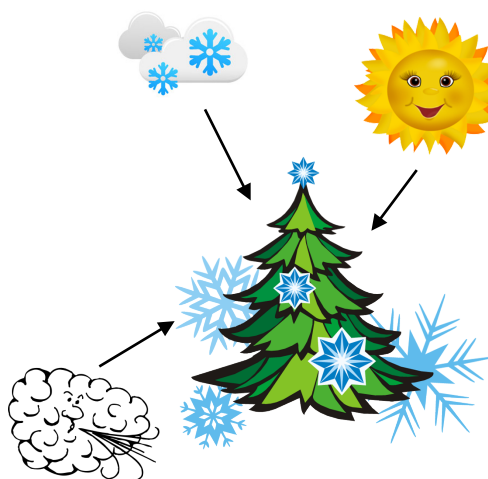


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Influencing factors:

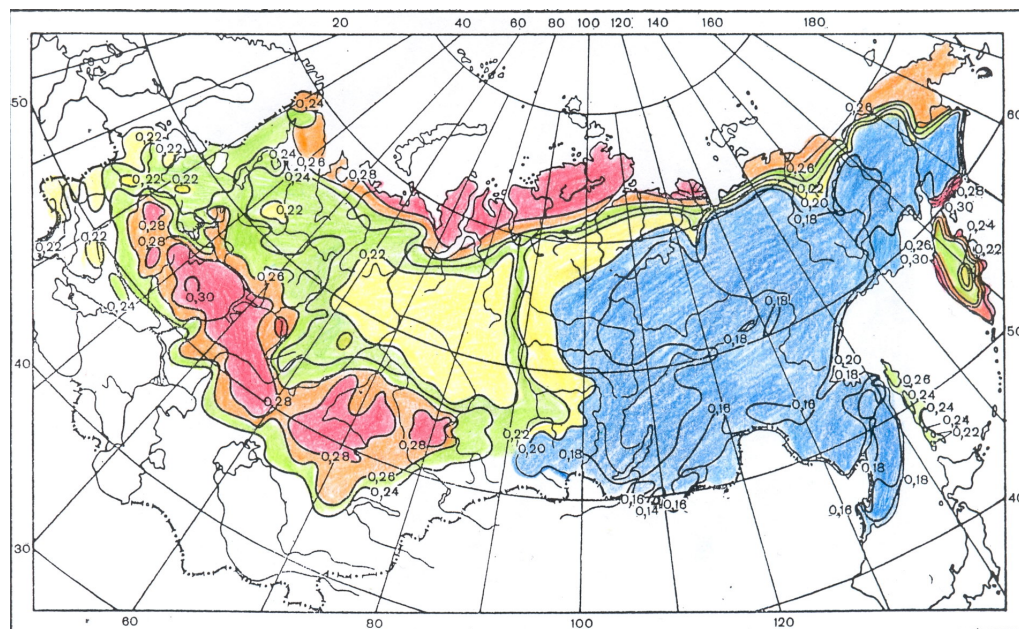
- Presence of thaws;
- Length of snow existence;
- The amount of precipitation;
- Compaction due to wind



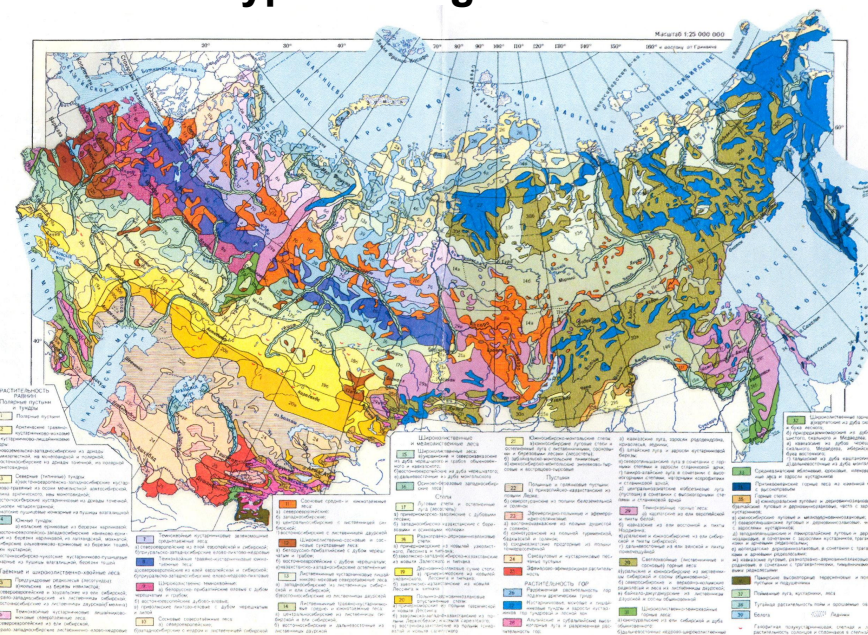


## Snow density depends on native zones

**Average snow density (g/cm<sup>3</sup>)  
at maximum ten-day snow height**



**Types of vegetation**



- 0,28-0,30** tundra, forest-tundra; south - steppe
- 0,26-0,28** north - forest-tundra; steppe and broadleaf forests (with predominance of birch)
- 0,24-0,26** taiga (spruce; spruce with a touch of oak); south - steppe

- 0,20-0,22** swamp vegetation, taiga (spruce); west - broadleaf forests (with predominance of oak and hornbeam)
- <0.20** taiga (with predominance of larch)







# Snow fractional covering: experiments

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, \frac{W_{snow}}{\delta_s}))$$

$\delta_s = 0.015$  parameter

$W_{snow}$  - snow water equivalent

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, 0.2))$$

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, 0.5))$$

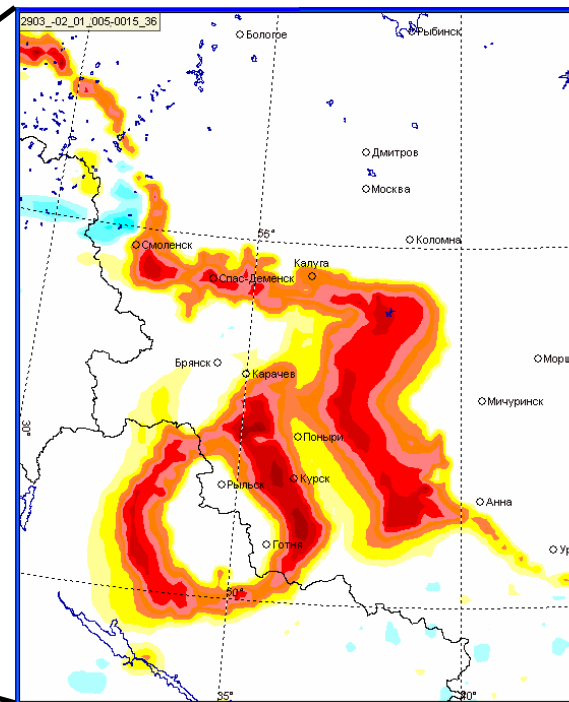
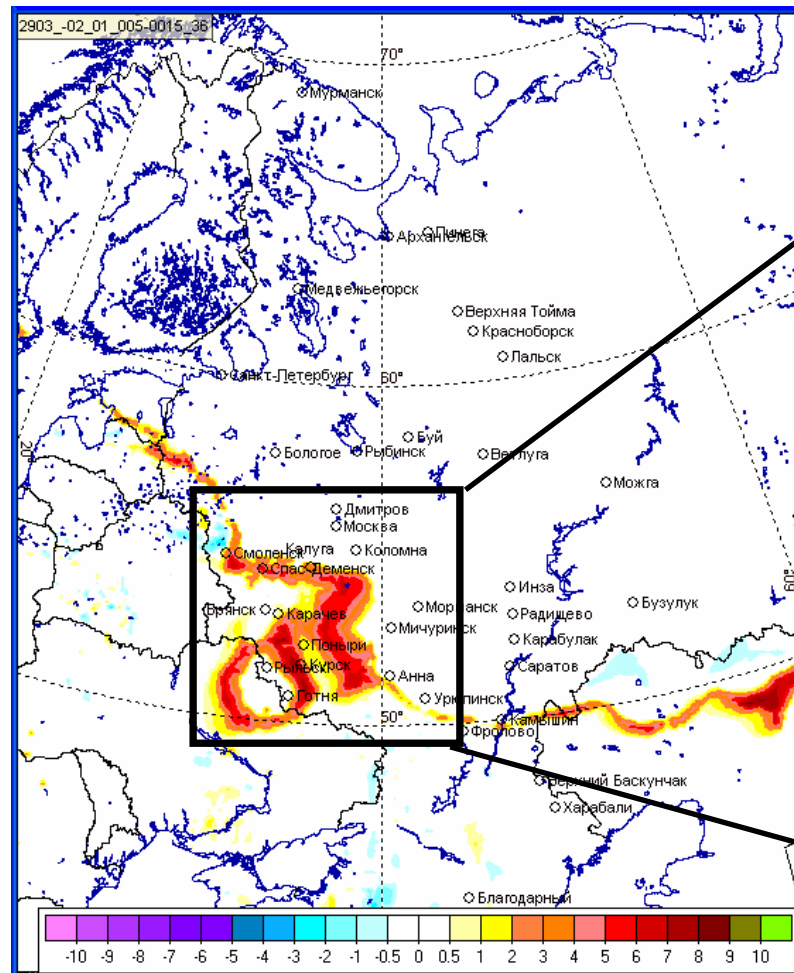




# $\Delta T_{2m}$ (29.03.10)

$$cf\_snow = 0.05m$$

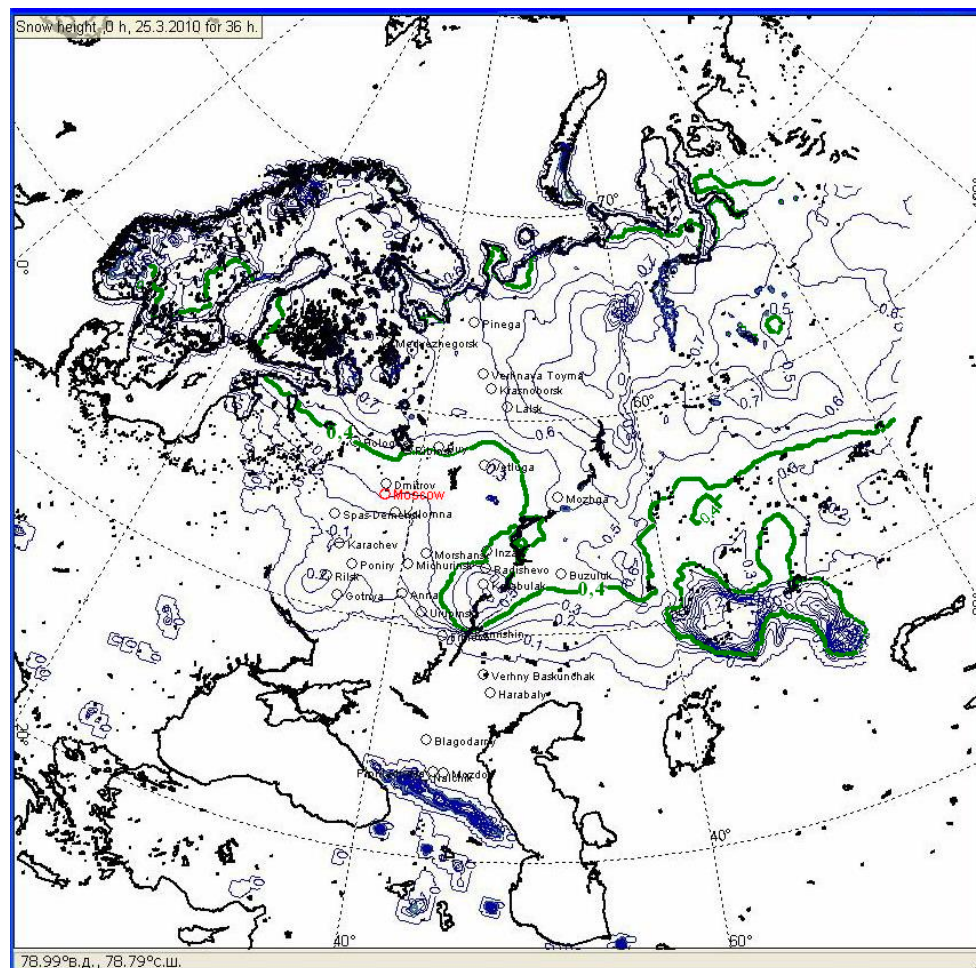
если  $T \geq -0,2^{\circ}C$ , то  $cf\_snow = 0.1m$







## Snow height (m), 36h forecast



According to model 36h forecast of snow height (see fig.), 40-cm snow exists on a wide territory in model as well as in the reality. However model snow mass is more big then according to observations (due to overestimated initial model SWE) and measured T2m is above 0°C







It was decided to change **longwave radiation** and **total forcing at snow surface** as they gave the main effect on the instability in calculating a forecast. So, only in these terms fractional cover was calculated as usual :

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, \frac{W_{snow}}{\delta_s})) \quad \delta_s = 0.015$$

For the rest of terms (net radiation, latent and sensible heat fluxes for snow and soil as well as evaporation and evapotranspiration) fractional cover was:

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, \frac{H_{snow}}{b_s})) \quad b_s = 0.4$$







# Total forcing at snow surface

$$G_{snow} = f_{snow} (c_p H_{snow} + LE_{snow} + Q_{rad,net} + G_p)$$

$G_{snow}$  total forcing at snow surface

$c_p H_{snow}$   $LE_{snow}$  sensible and latent heat fluxes at snow surface

$Q_{rad,net}$  net radiation at snow surface

$G_p = L \cdot P_r$  freezing rain

or  $G_p = -L \cdot P_{snow}$  melting snowfall

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, \frac{W_{snow}}{\delta_s}))$$
 surface fraction covered by snow







# Longwave radiation

$$Q_{lw} = \sigma(1 - \alpha) \cdot ((1 - f_{snow}) \cdot T_s + f_{snow} \cdot T_{snow})^4 + Q_{gr}$$

$Q_{lw}$  downward longwave radiation

$\sigma$  Boltzmann-constant

$\alpha$  thermal albedo ( of all soil types )

$$f_{snow} = \text{Max}(0.01; \text{Min}(1.0, \frac{W_{snow}}{\delta_s}))$$

surface fraction covered by snow

$T_s$  soil surface temperature

$T_{snow}$  snow surface temperature

$Q_{gr}$  thermal radiation at the ground

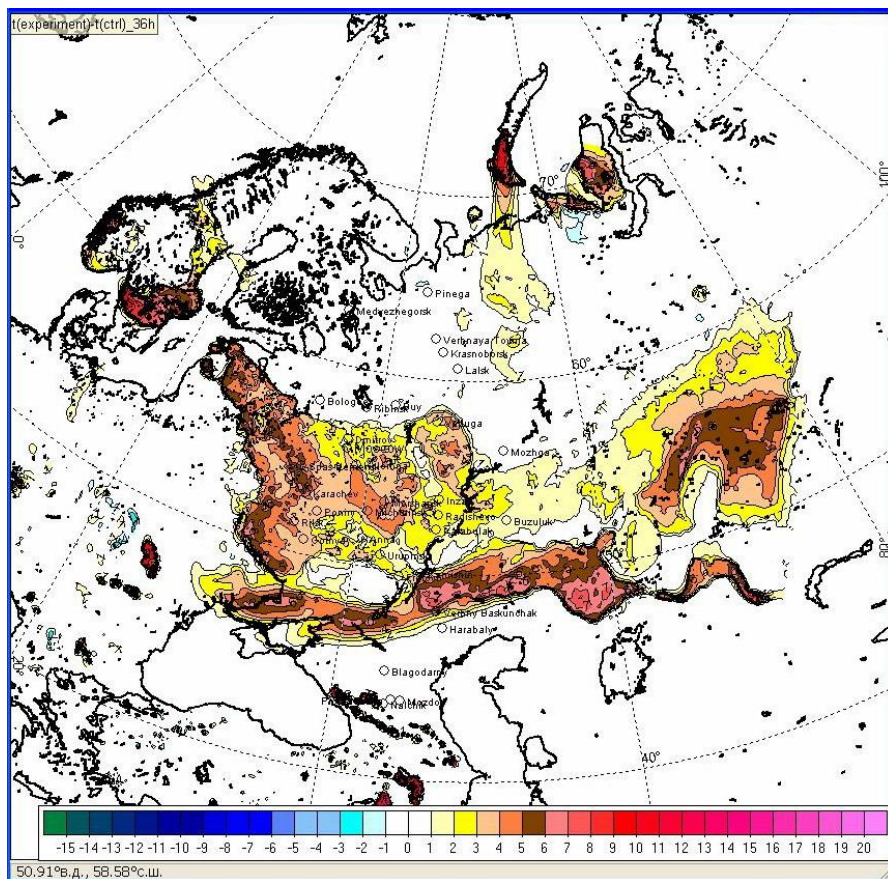




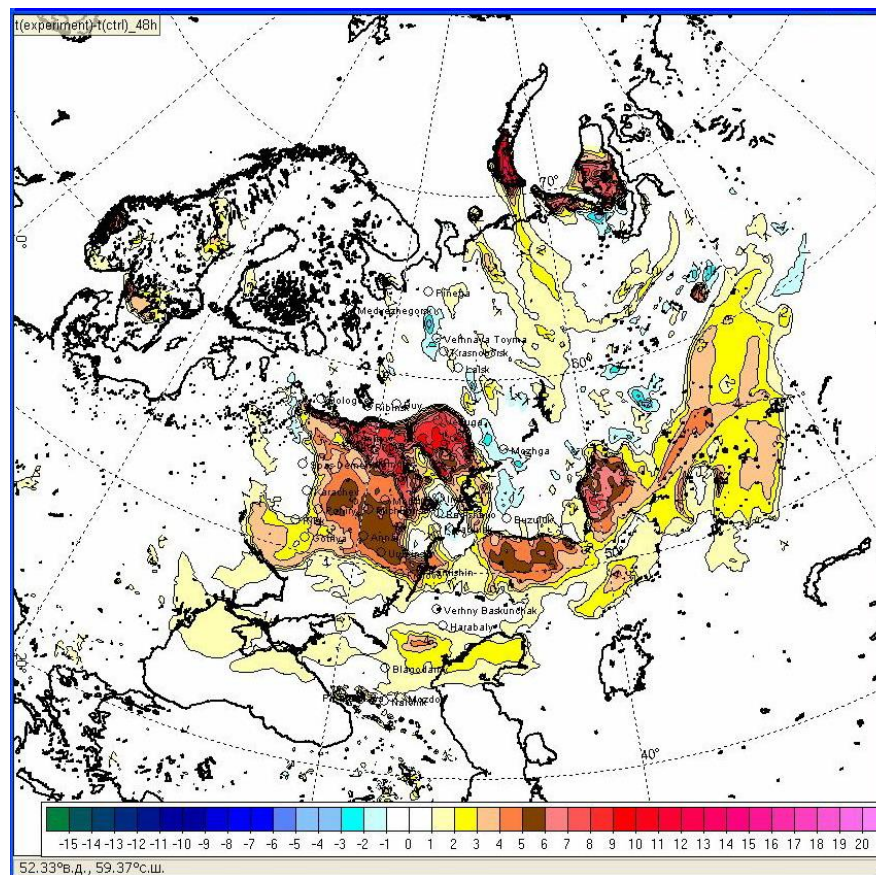


# T2m(experiment)-T2m(ctrl)

36h forecast



48h forecast



T2m temperature raised up to 10°C on the territory where according to model data snow cover was below 40 cm

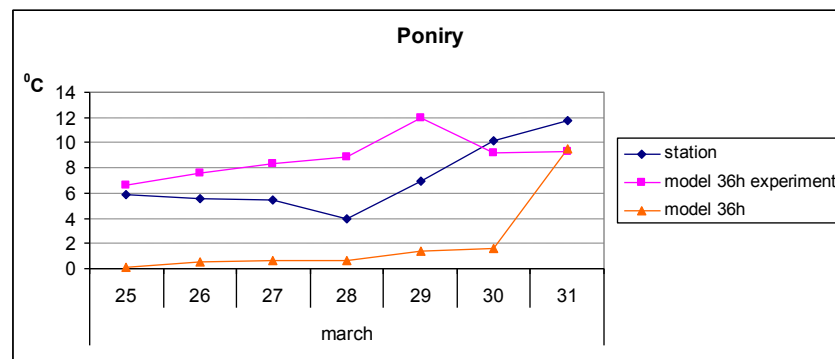
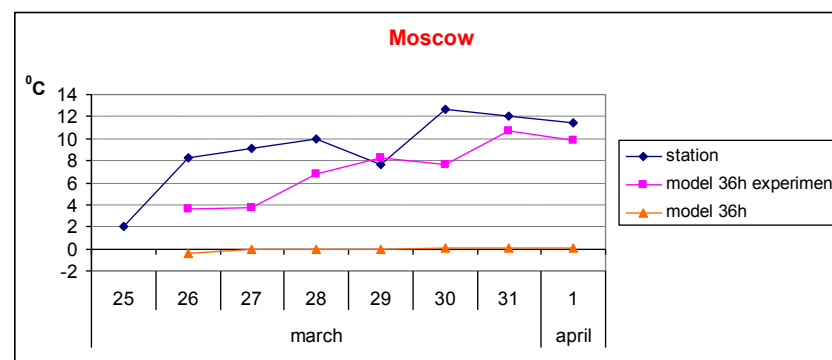
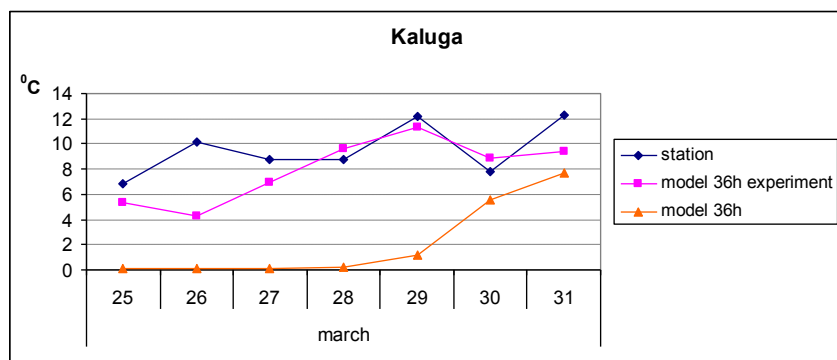








## For 60h forecast the tendency remains as for 36h







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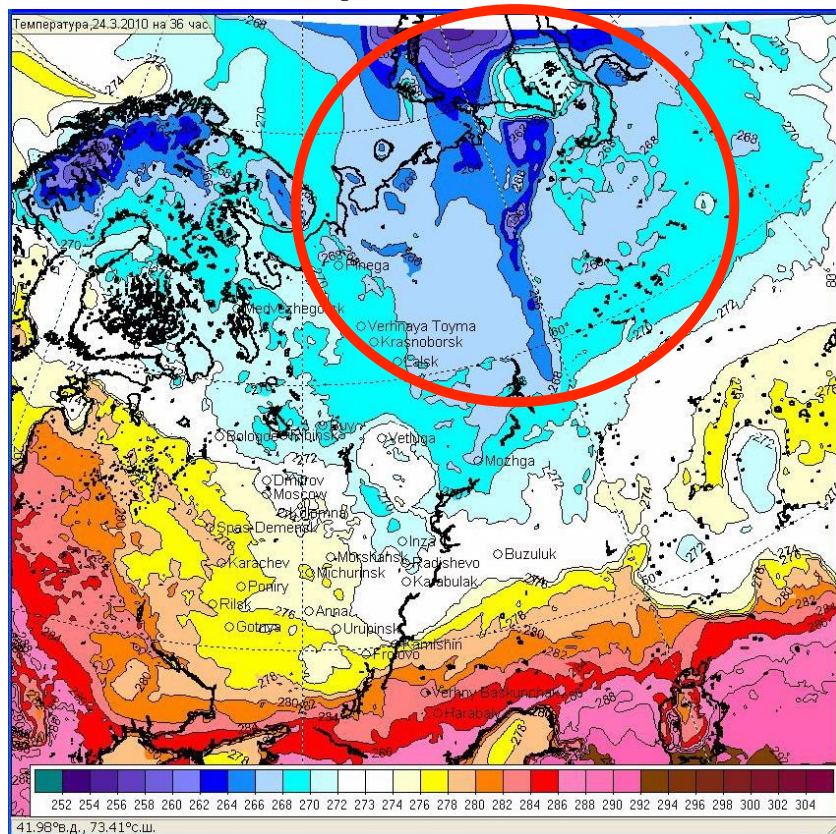
	AE experiment	AE ctrl model
Vetluga	-1,1	-6,5
Moscow	-2,7	-9,3
Kaluga	-1,5	-6,9
Karachev	0,4	-4,1
Rilsk	-0,4	-6,2
Anna	0,2	-6,2



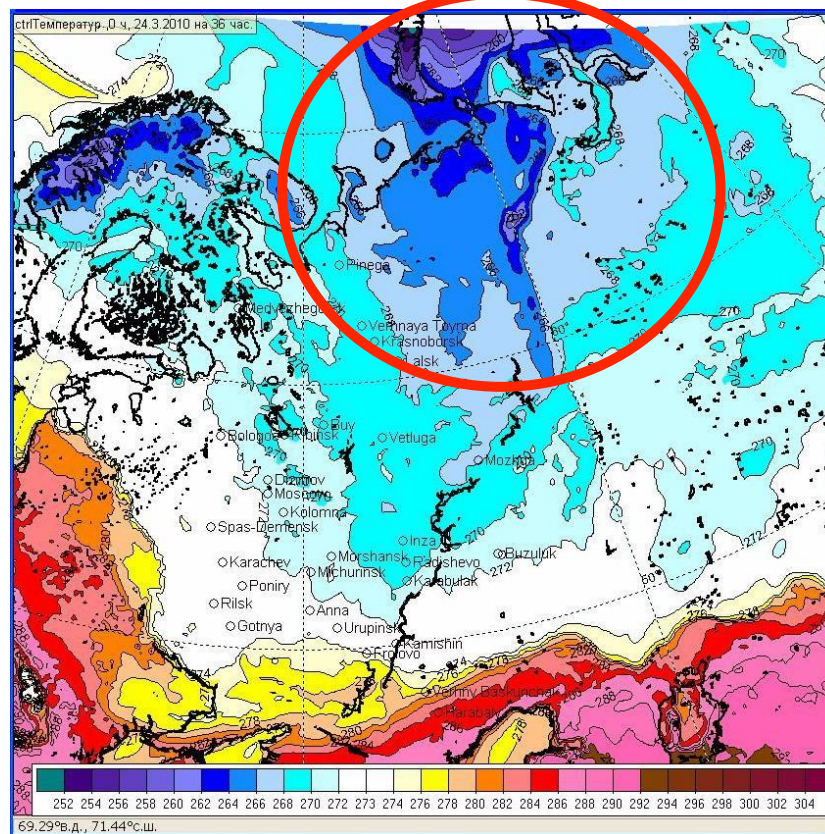


# T2m (K), 36h forecast

experiment



Ctrl model



Changes also occurred to northern regions, where snow height was more than 40 cm: areas with extreme T2m were reduced

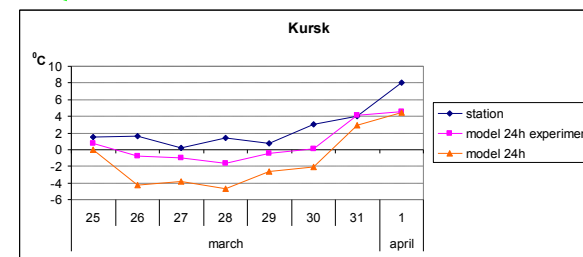
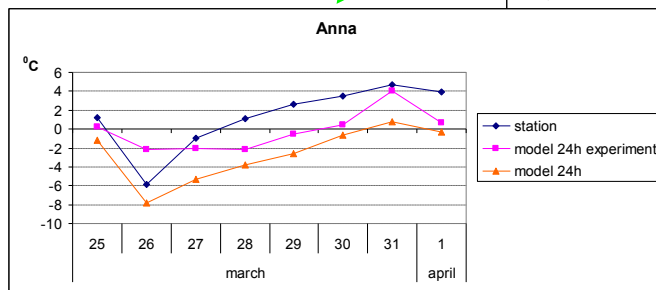
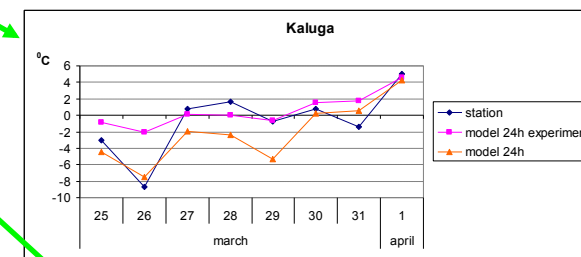
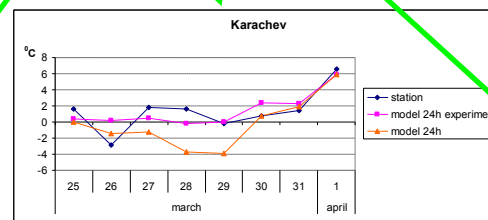
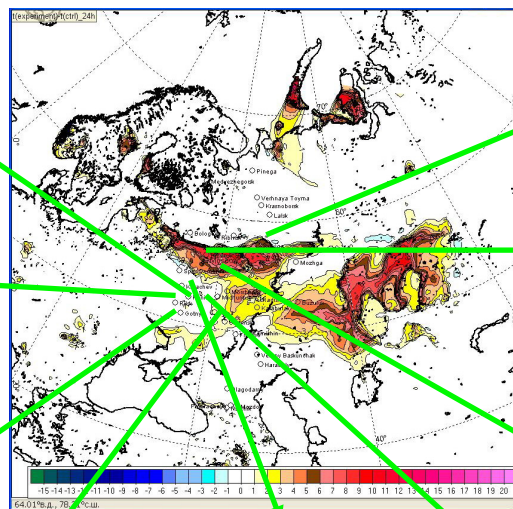
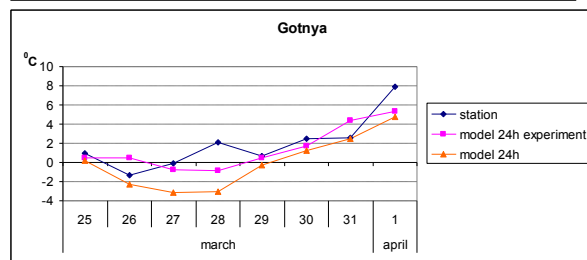
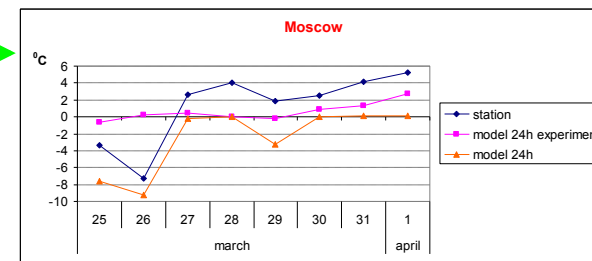
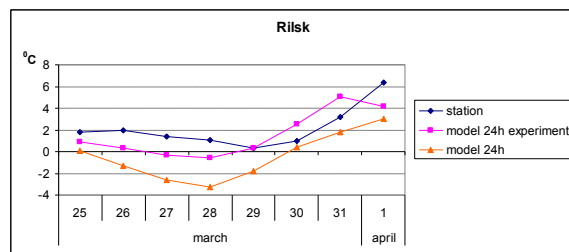
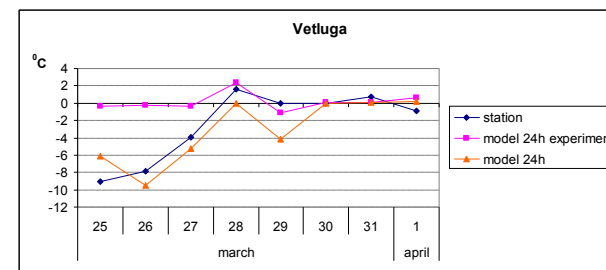
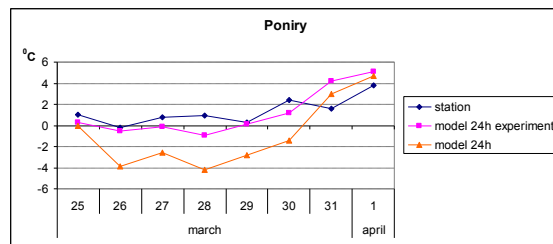




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Night T2m temperature ( $^{\circ}\text{C}$ ): measurements (blue), 24h forecasts – experiment (pink), ctrl model (orange)





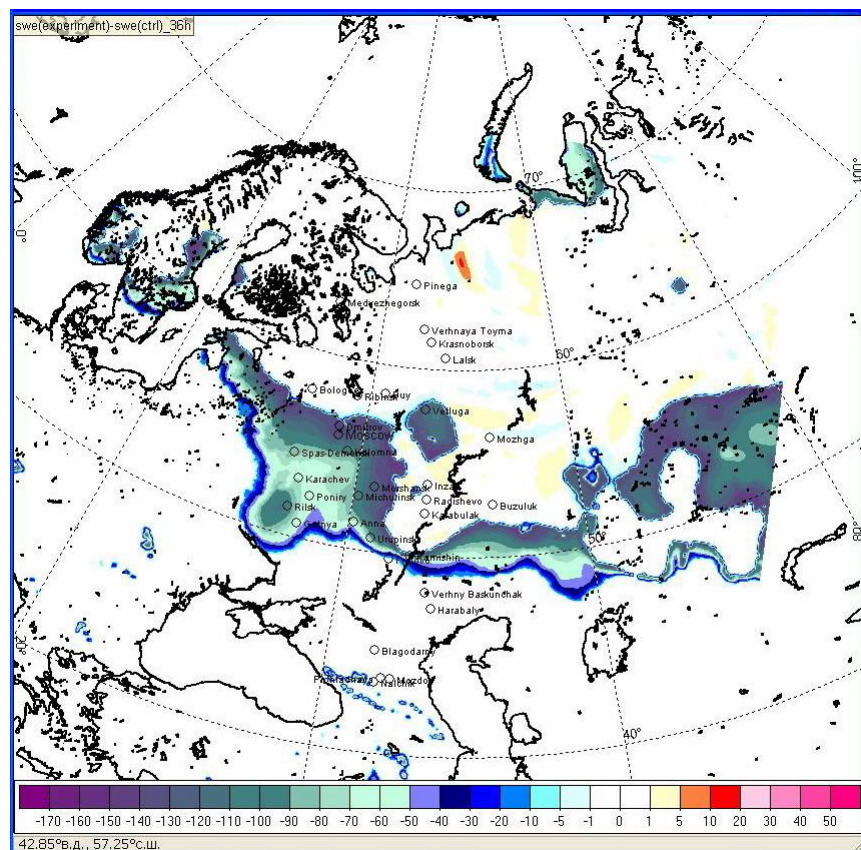


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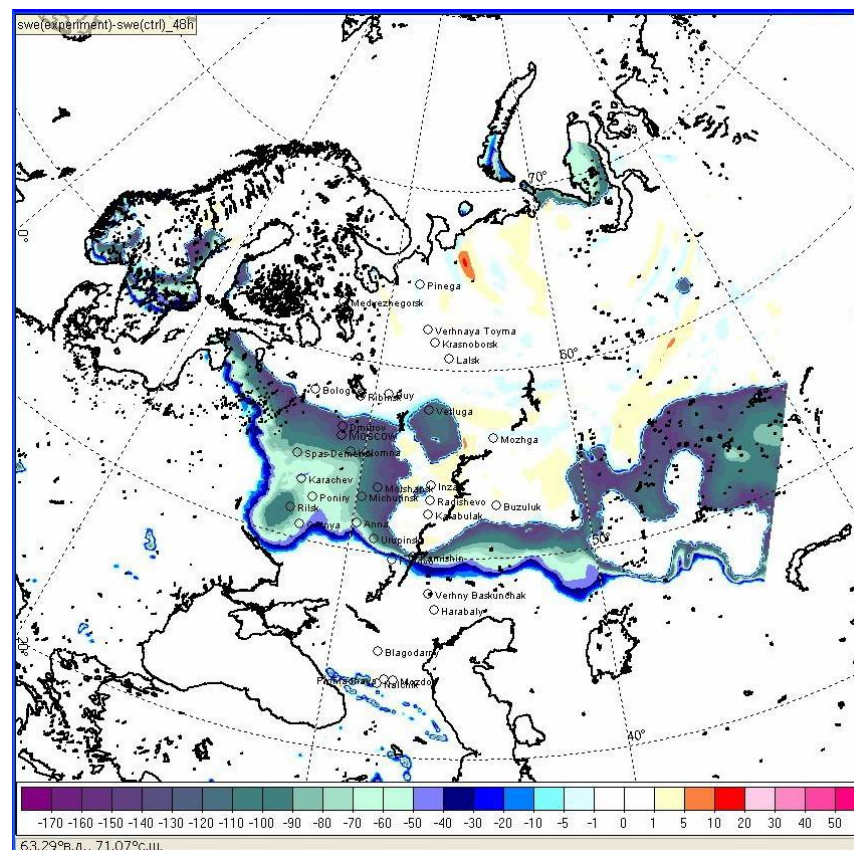


# SWE(experiment)-SWE(ctrl), mm

36h forecast



48h forecast

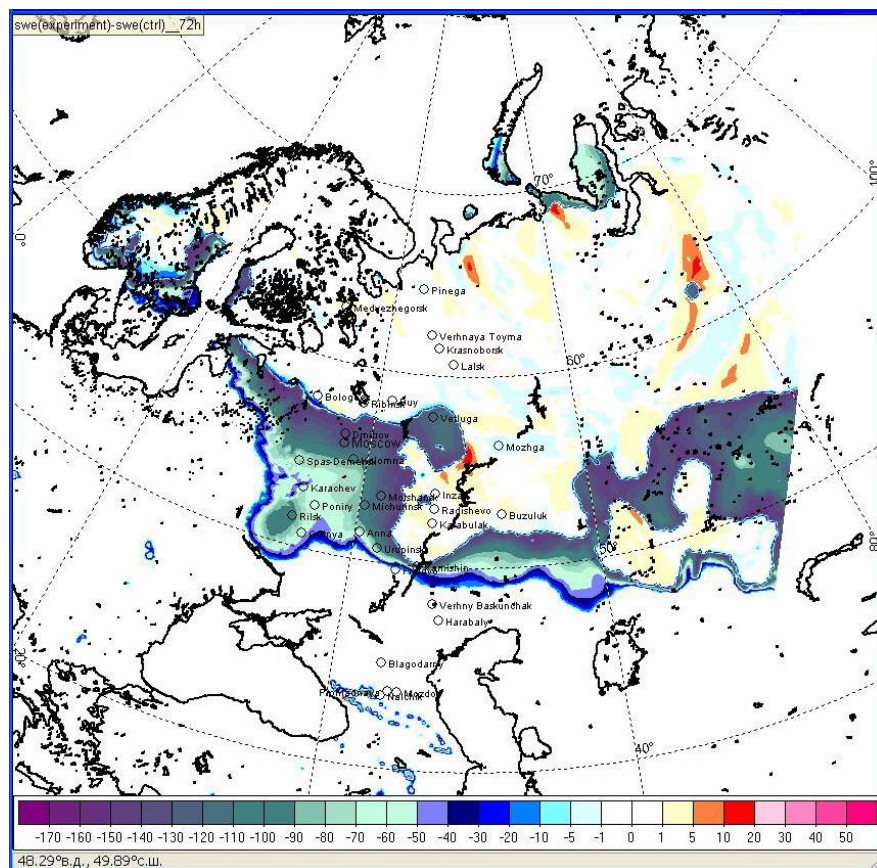




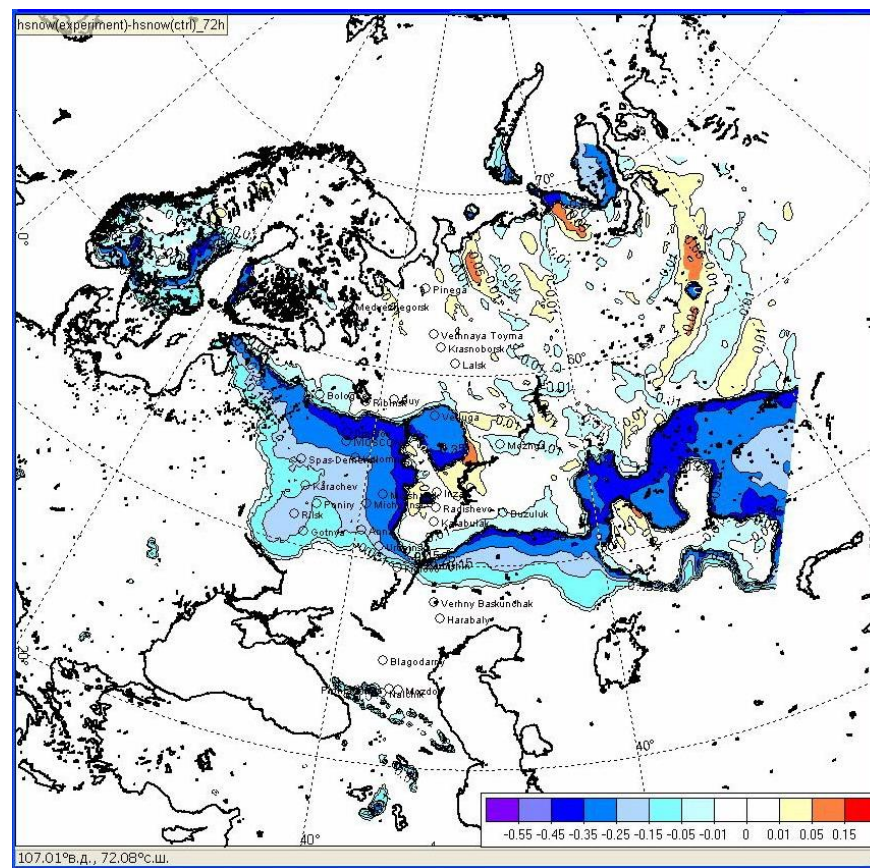


## 72h forecast

**SWE(experiment)-SWE(ctrl), mm**



**Hsnow(experiment)-Hsnow(ctrl), m**







# Conclusions

- SWE used as initial condition for COSMO-Ru run is overestimated, correspondingly, snow density is significantly overestimated as well
- It impacts the predicted near surface air temperature (T2m) during snowmelt period caused by the excessive snow mass, which prevents air to get warmer
- The more realistic algorithm of snow density calculation considers meteorological conditions during the period of snow presence (“temperature” method)
- Constants are in the empirical dependence of “temperature” method, and have geographical distribution close to configuration of vegetation-native zones
- Changes in algorithm of fractional snow coverage ( $SWE \rightarrow H_{snow}$ ) lead to possibility to simulate air warming more realistically (COSMO-Ru, version 4.2)

It is necessary further investigations of fractional snow coverage impact in the new COSMO-versions (4.13, 4.17)







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Thank you for your attention!

