JAN WACLAW PARFINIEWICZ

IMGW Poland

# 1 Introduction

During July-September 2006 two versions of COSMO-Model: 3.5 and 3.19 (both in terrain following coordinates) have been tested in parallel. Altogether 90 precipitation forecasts from parallel runs starting at 00 UTC for 30h have been completed and compared against daily totals GTS SYNOP data over Europe (the runs 20060701 and 20060729 have been dropped out). The models were run with diagnostic precipitation scheme version as it has been set into the operational IMGW model LM\_3.5 14 km grid version (taken as reference), which is stably running there since about 2 years.

# 2 Standard statistics

The averaged results of the statistical indices used for quantification of precipitation forecasts LM\_3.5 vs. LM\_3.19 are the following (PC - proportion correct, HIT - probability of correct rain signal, FAL - false alarm, BIAS - systematic error, HSS - Heidke skill score, PSS - Peirce skill score, ETS - Equitable threat score, CHI - the measure of cross relationship with  $\chi^2$  distribution, DEV - standard deviation error, and, CC - cross correlation coefficient):

LM:	$\mathbf{PC}$	HIT	FAL	BIA	HSS	$\mathbf{PSS}$	ETS	CHI	$\mathbf{C}\mathbf{C}$	DEV
3.5	0.81	0.75	0.20	1.26	0.52	0.55	0.37	0.30	0.50	5.16
3.19	0.81	0.74	0.20	1.21	0.52	0.54	0.36	0.30	0.50	5.28

Table 1: Standard statistics: averaged values

An example of ETS diagram is presented in Fig. 1. The vertical lines indicate dates when significant errors occurred (in red for LM\_3.5, in green for LM\_3.19).

As standard statistics are not sensitive enough to assess model behavior in extreme cases, it is necessary to work with all spectrum of errors.

### 3 Comparing overall 30h forecast errors distributions: LM\_3.5 vs. LM\_3.19

Distribution of errors in both models was based on more than 50000 sample points alltogether. In Fig. 2 two diagrams of precipitation errors are presented; left: monotonically arranged (ranked) by self - order from smallest (negative) to largest (positive); right: by classes of errors' magnitude. As the errors of order < -2, 2 > (the thick band for numbers between  $\sim 5000 - 44000$  on right panel) are negligible (however it may comprise cases of forecasts with severe precipitation), the further analysis will concentrate on separating extreme cases from dominating "dust" by zooming right and left parts of the diagram.



Figure 1: ETS diagram



(a) Errors by selforder

(b) Errors by classes

Figure 2: Overall 30h forecast precipitation errors distributions

In Fig. 3 investigation of the two border distributions is presented in the following way: left for values under predicted ones (dry branch) and right for overshoots (wet branch).

The further way to proceed is clear: select doubtful, significant cases and attempt to develop new model version that would clarify situation (or enlarge range of clarity). A control list of 162 point-cases errors > 40 [mm/day] was selected, which gives altogether 56 dates to monitor model quality progression (the administration of such a list is a separate challenge). Each of such cases has its own metric and graphical illustration. An example of the extremely imperfect LM\_3.19 forecast is the 8th of July, when daily total error for Chemnitz, Germany (10577) was 121 [mm/day], while for LM\_3.5 "only" 53 [mm/day] (Fig. 4). The illustration shows influence of singular error signal onto water accumulation over a given station. When taking accumulated water amount for an enough long period of aggregation as an ultimate indicator of forecast quality, one may see how huge forecast collapses are then smoothly compensated by a number of less significant events.



(a) Errors by selforder: left side

(b) Errors by selforder: right side

Figure 3: Investigation of the two border distributions



Figure 4: Accumulated precipitation and daily totals: Chemnitz Germany (10577)

#### 4 Comparing accumulated precipitation water errors over Europe

- a) Distribution of errors over Europe : LM\_3.5 vs. LM\_3.19. (Fig. 5)
- b) Relative Predominance LM\_3.19 over LM\_3.5 regarding accumulated water (Fig. 6). To asses each particular relationship between errors over given geographical point, the simple predominance factor RP is proposed:

 $RP = (Relative Predominance (|DelLM_3.5|)/(|DelLM_3.5| + |DelLM_3.19|)$ 

where DelLM\_(given version) means difference between 30 h forecast of daily total (i.e. 30h -06h model forecasts of total precipitation) and its realization over given station. If the number of errors in LM\_3.5 is large, the RP of LM\_3.19 approaches the value 1. Calculated RP of LM\_3.19 of accumulated water allows for ultimate geographical review of errors over Europe (Fig. 6).



Figure 5: Comparing Accumulated Precipitation Water Errors over Europe



rPREDOMINANCE Lm\_3.19 over Lm\_3.5 of Accum.Precip.Errors: JUL-SEP2006

Figure 6: Relative Predominance LM\_3.19 over LM\_3.5 regarding accumulated water

#### 5 Summary

Despite there is no evidence that Total\_Precipitation LM\_3.19 30h forecast errors are significantly smaller than in the LM\_3.5 version, the accumulated water errors during the analyzed season indicate clear predominance of the newer version. This result, obtained for summer season, indicates that configuration applied at IMGW till now, should be changed to account for new developments - relevant to finer mesh size (2.8 km), the CLM developments and changes in the new multi-layer soil model. Similarly, the diagnostic precipitation parametrisation should be re-tested against the prognostic one in the near future.

#### Acknowledgments

I appreciate the input to this work by my valued colleagues: Uli Schättler, Marek Lazanowicz, Jürgen Steppeler.