

# Assimilation of 3D radar reflectivity into the Local model

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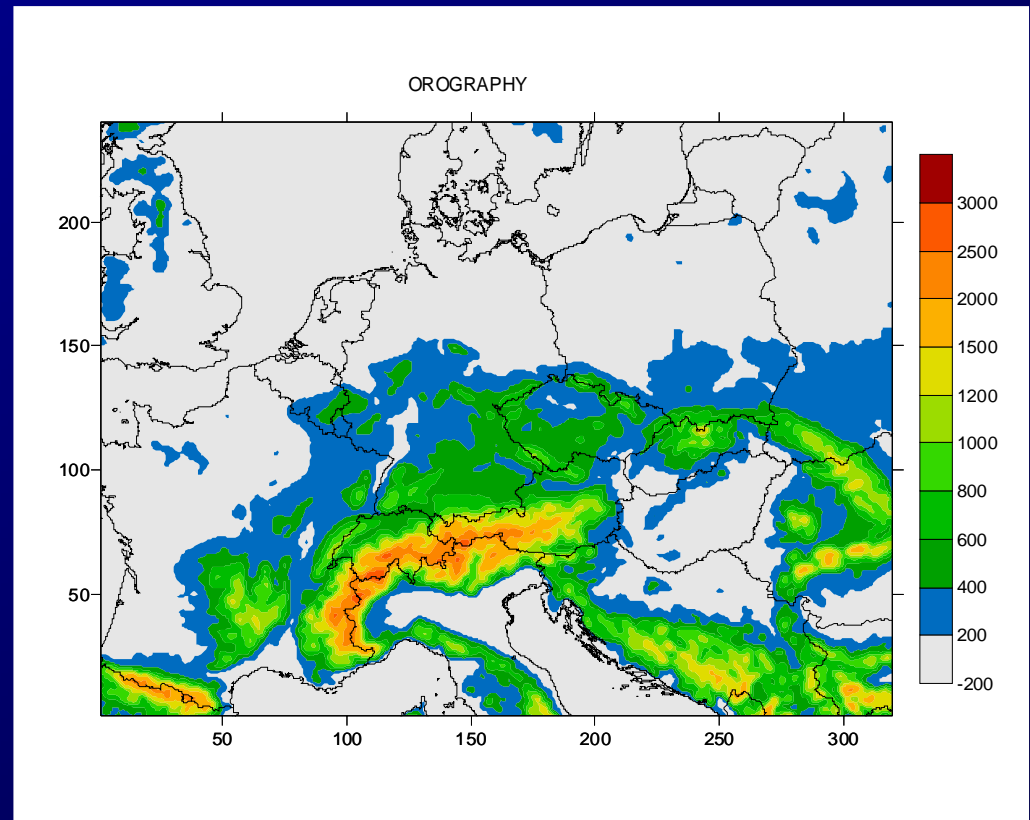
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# Content

- LM model
- Radar data
- Assimilation methods: 2D and 3D
- Case studies
- Conclusions

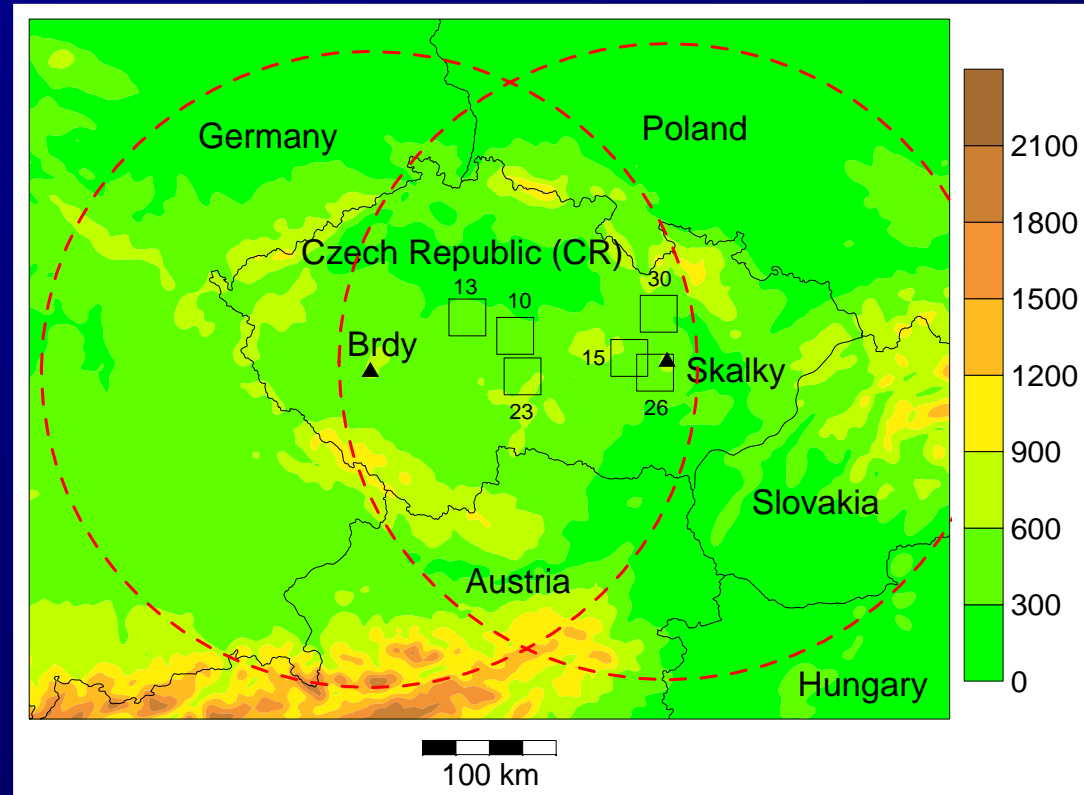
# Implementation of the LM (version 3.18)

- $\Delta h = 11$  km
- 35 vertical levels
- ECMWF analyses
- Tiedtke's cumulus parameterization
- Prognostic variables:
  - rain water
  - ice



# SLM

- $\Delta h = 2.8$  km
- 35 vertical levels
- Explicit precipitation
- Prognostic variables:
  - rain water
  - ice
  - graupels

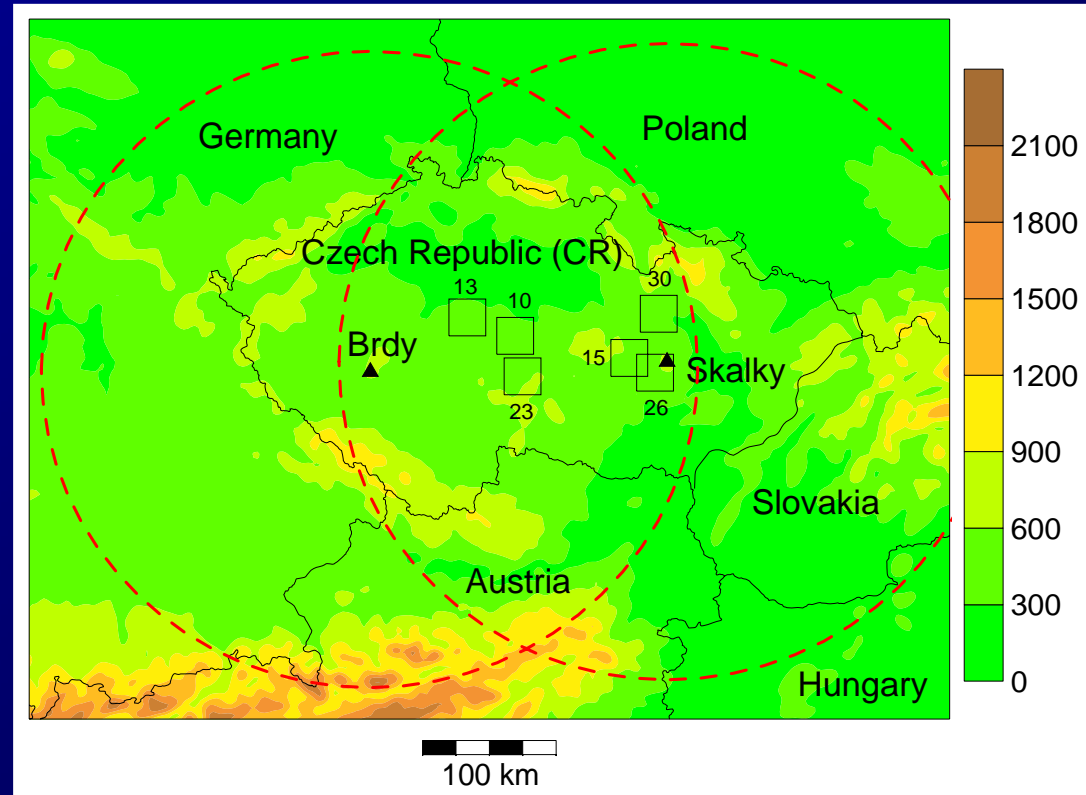


# Radar data

- Brdy, Skalky  
C-band
- Resolution 1km x 1km
- $\Delta t=10$  min.
- Z-R  $\alpha=200, \beta=1.6$

$$Z = \alpha R^\beta$$

- CAPPI 2km (ground precipitation)
- CAPPI at 1, 1.5, 2, ..., 14 km



# Assimilation of radar reflectivity

WVC: Correction of  $q_v$  is based on the difference  
 $D = \text{"model precipitation"} - \text{"observed precipitation"}$   
unit: [kg/kg]  
(Falkovich et al., 2000) and (Sokol and Rezacova, 2006)

IF(  $D > 0$  )

$$q_{v,k}^{\text{new}} = q_{v,k} + \text{DIF}$$

$$\text{DIF} = w(z) * \text{MIN}( \alpha * D, \delta )$$

$$\delta = \text{MAX}( \varepsilon^+ * q_{v/k}^s(T_k) - q_{v/k}, 0 )$$

ELSE

$$q_{v,k}^{\text{new}} = q_{v,k} - \text{DIF}$$

$$\delta = \text{MAX}( q_{v,k} - \varepsilon^- * q_{v/k}^s(T_k), 0 )$$

$\varepsilon^+$ ,  $\varepsilon^-$ ,  $\alpha$  ... constants,  $k$  ... vertical index,  $z$  ... elevation  
 $w(z)$  is weighting function

# Types of assimilation methods

- WVC 2D

$D = \text{"model prec."} - \text{"observed prec."}$   
at the ground

- WVC 3D

$D = \text{"model prec."} - \text{"observed prec."}$   
at single model levels

- WVC VIL (3D)

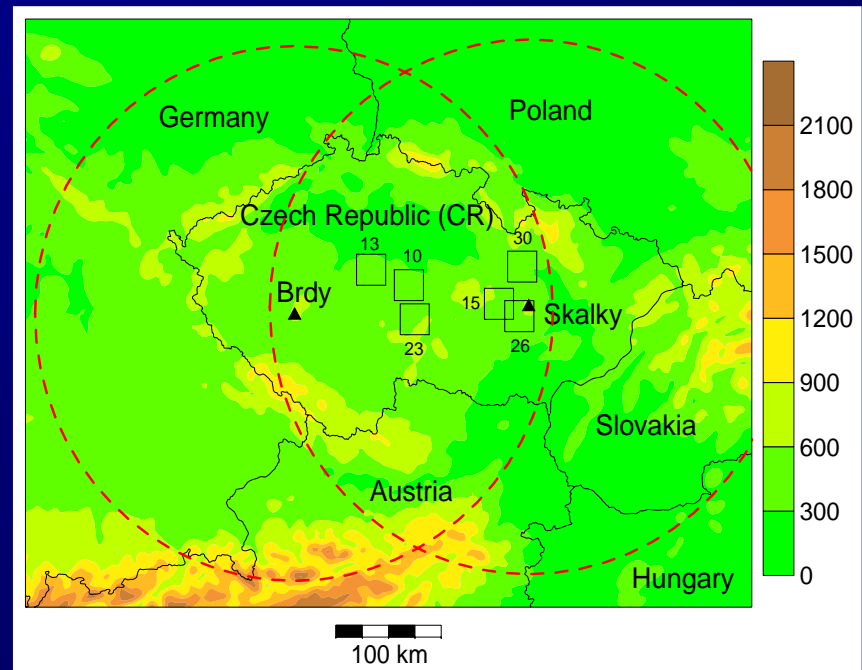
$D = \text{"model VIL"} - \text{"observed VIL"}$

VIL – vertically integrated water vapour content

- LHN (2D)

# Case studies

- Six convective storms from the CR (10 June 2004 14UTC, 13 July 2002 14UTC, 15 July 2002 15UTC, 23 May 2005 14UTC, 26 May 2003 12UTC, 30 May 2005 14UTC)
  - Duration approx. 3 hr,  $[T, T+3]$
  - Local storms
  - Observed precipitation by rain gauges: 50 – 170 mm
  - The centres of the events  $(g_1, g_2)[T, T+3]$
- 
- Only radar data are assimilated and used in the verification
  - Radar underestimates precipitation



# Verification - evaluation

- Event oriented
- 3h precipitation [T,T+3]
- Methods of verifications
  - “by eye”
  - “objective evaluation” (using subjective criterions)

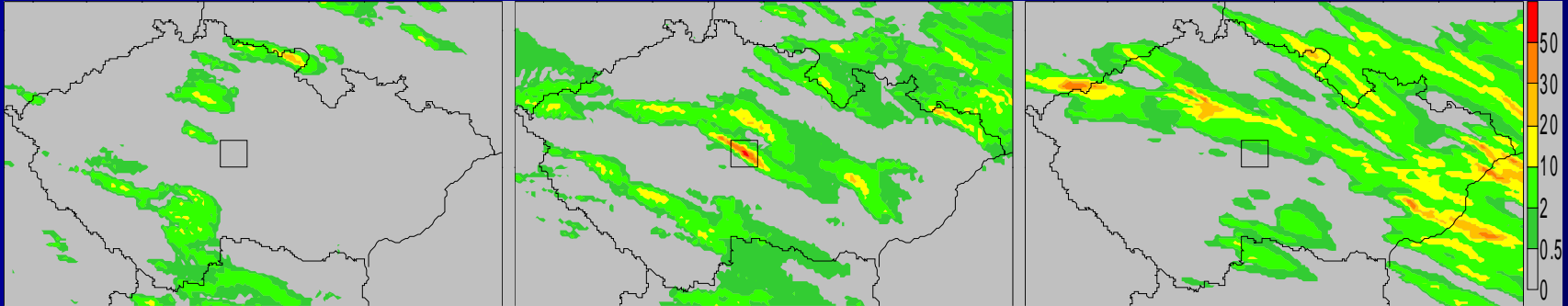
(i) Verification “by eye”

# 10 June 2004 - Observed precipitation T=14 UTC

[T-1,T]

[T,T+3]

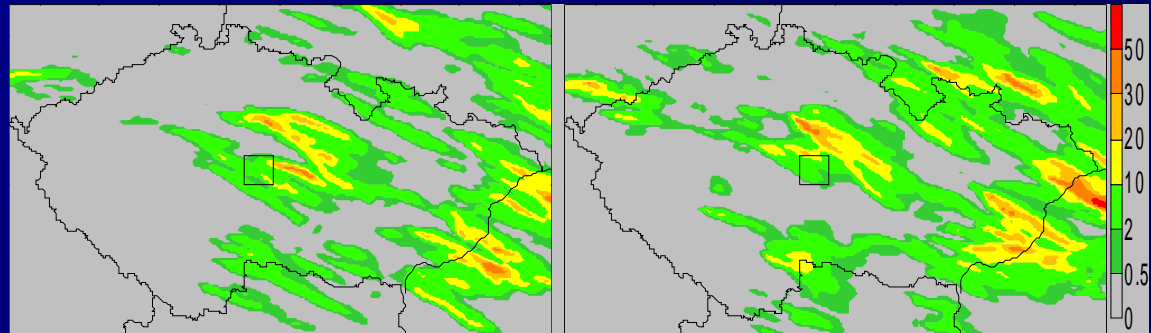
LM



LHN

WVC 2D

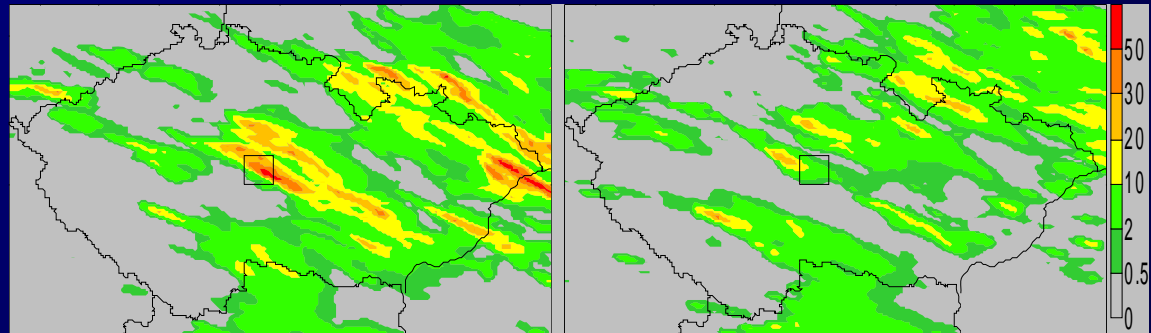
2D



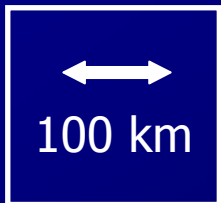
WVC 3D

WVC VIL

3D



Forecasted precipitation with assimilation

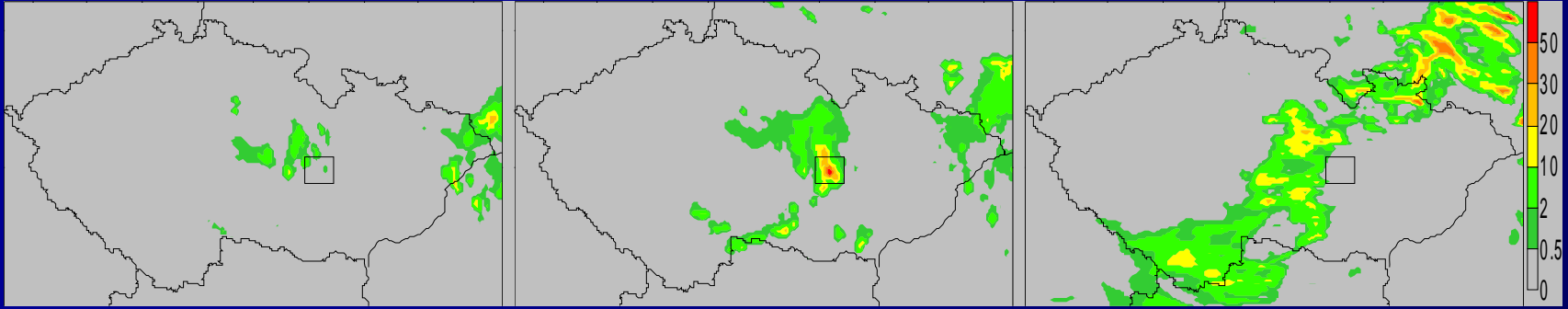


# 15 July 2002 - Observed precipitation T=15 UTC

[T-1,T]

[T,T+3]

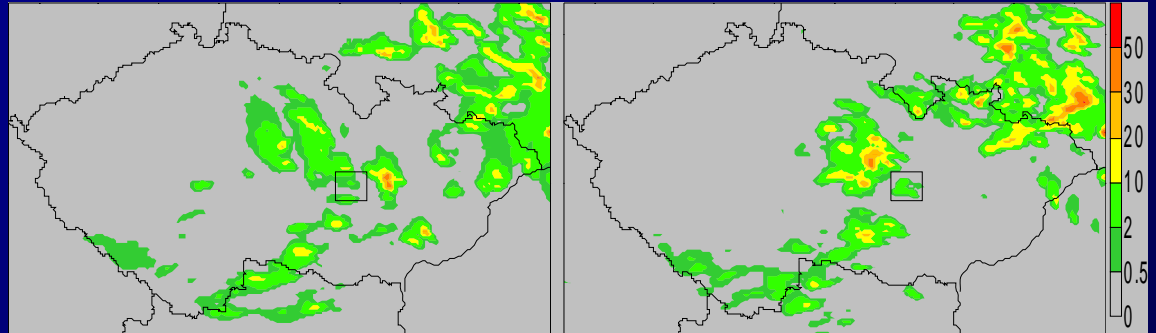
LM



LHN

WVC 2D

2D

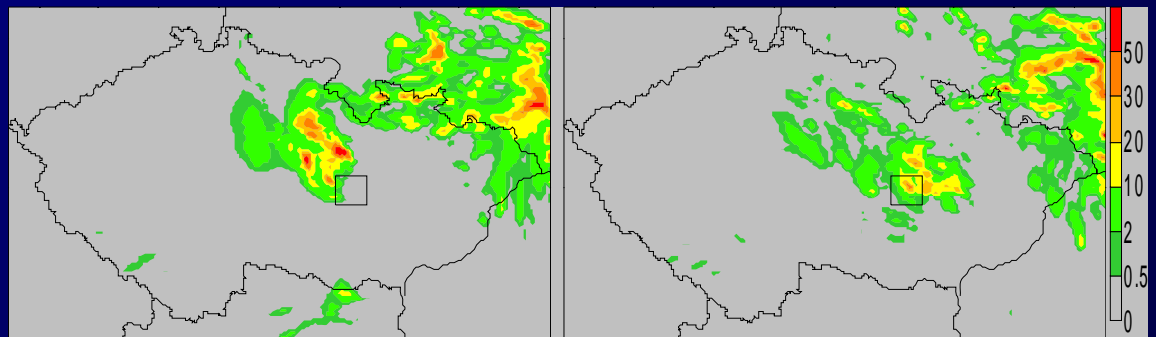


Forecasted precipitation with assimilation

WVC 3D

WVC VIL

3D

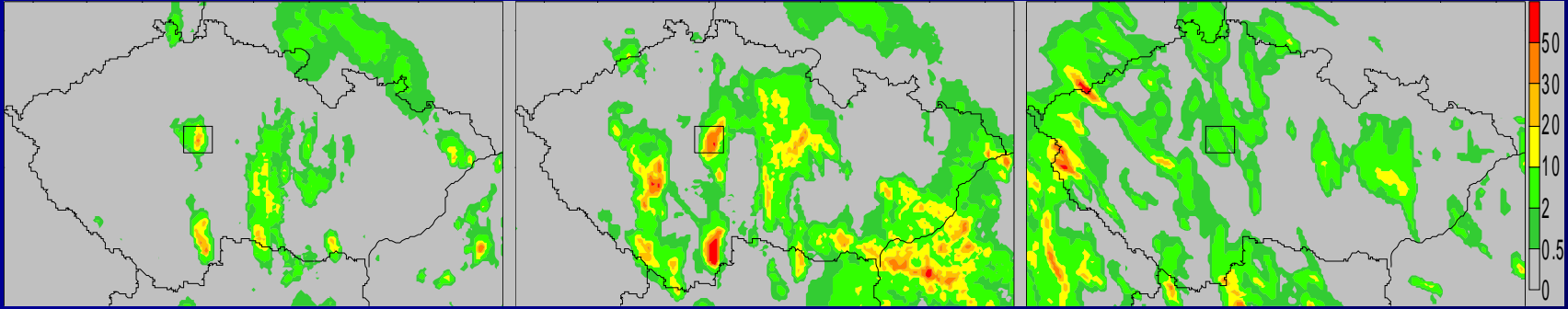


# 13 July 2002 - Observed precipitation T=14 UTC

[T-1,T]

[T,T+3]

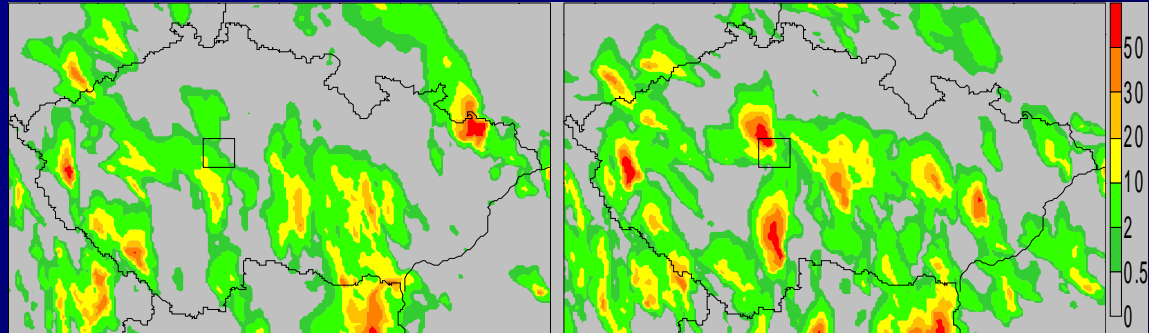
LM



LHN

WVC 2D

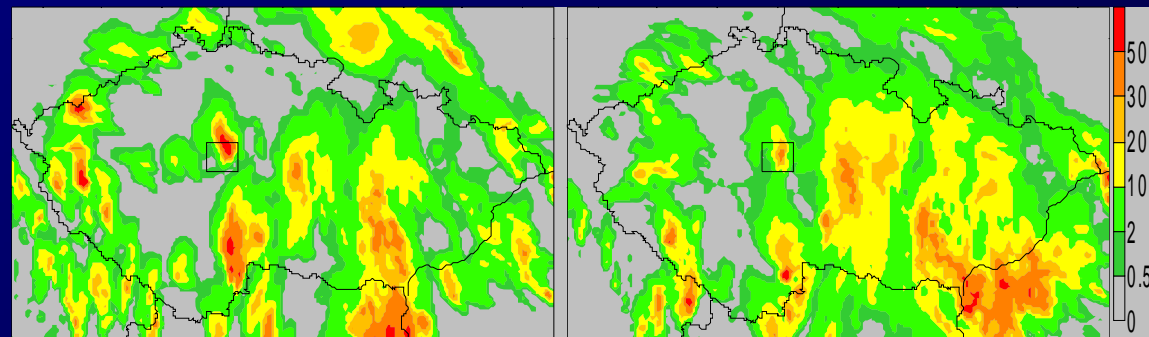
2D



WVC 3D

WVC VIL

3D



Forecasted precipitation with assimilation

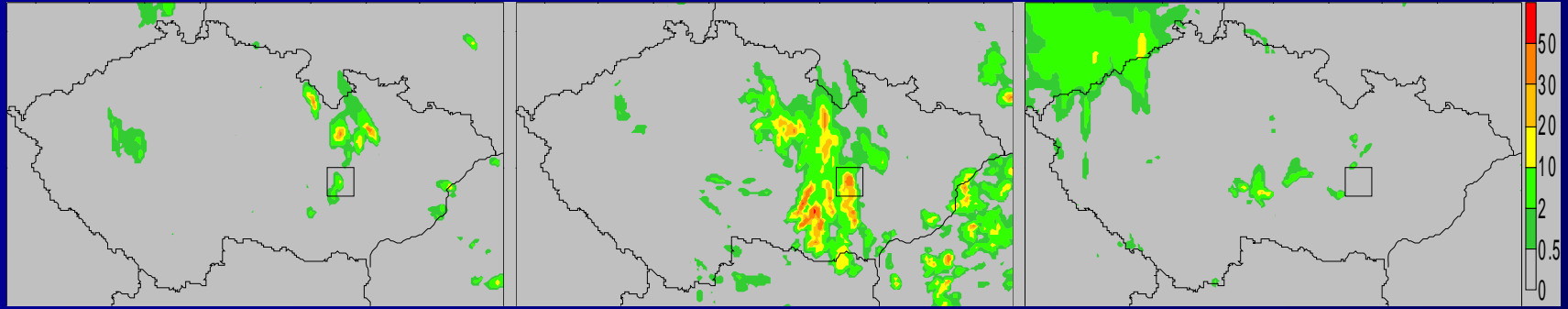


# 26 May 2003 - Observed precipitation T=12 UTC

[T-1,T]

[T,T+3]

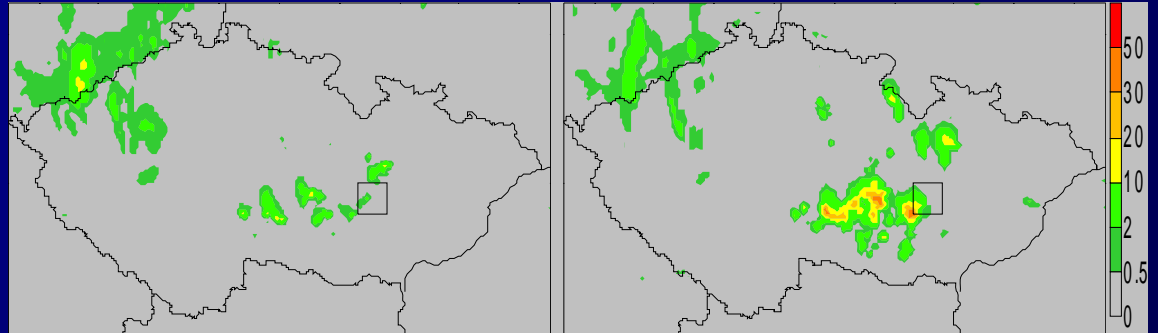
LM



LHN

WVC 2D

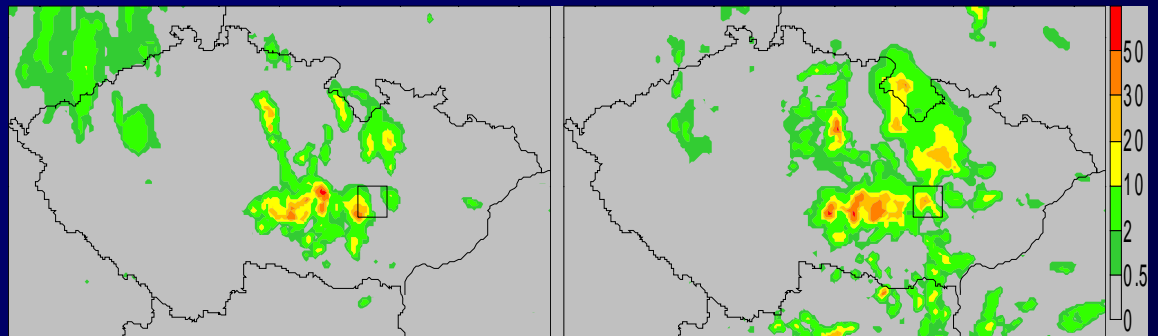
2D



WVC 3D

WVC VIL

3D



Forecasted precipitation with assimilation

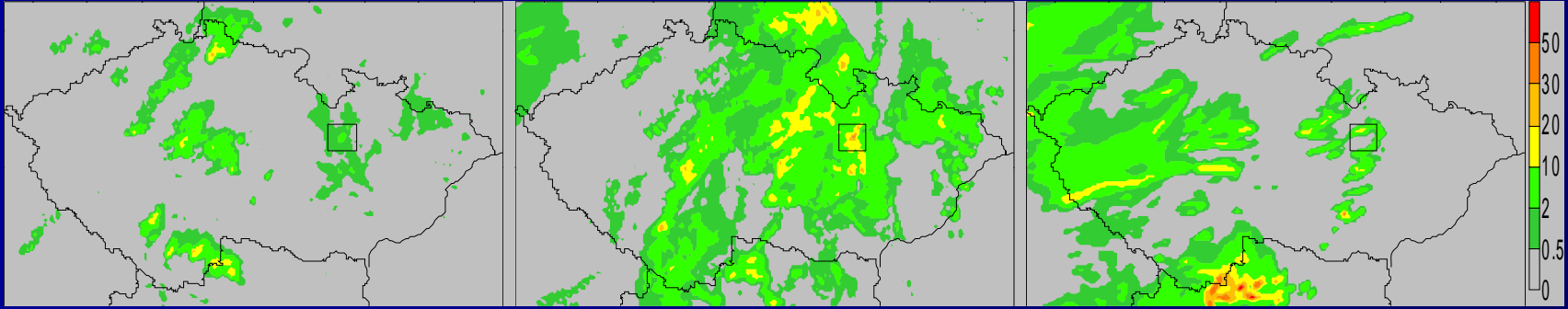


# 30 May 2005 - Observed precipitation T=14 UTC

[T-1,T]

[T,T+3]

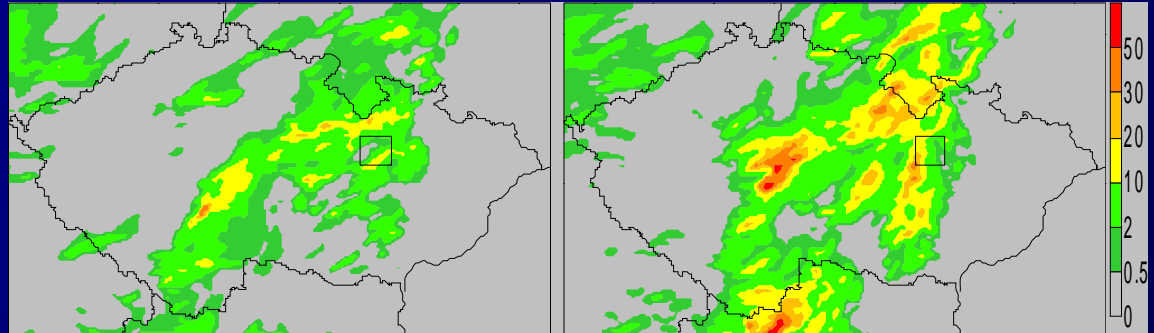
LM



LHN

WVC 2D

2D

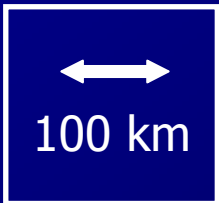
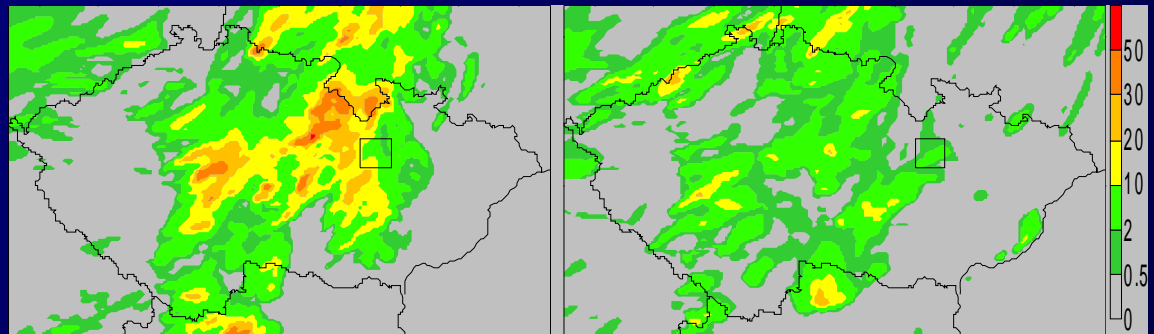


Forecasted precipitation with assimilation

WVC 3D

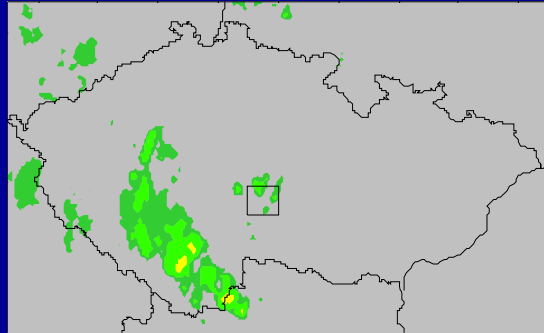
WVC VIL

3D

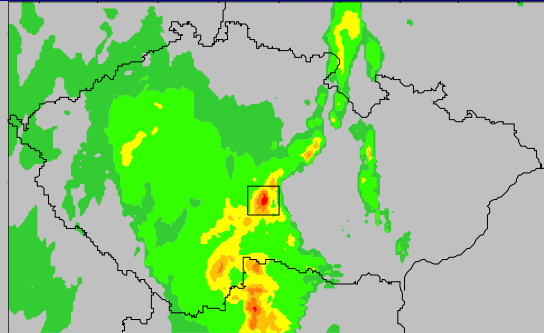


# 23 May 2005 - Observed precipitation T=14 UTC

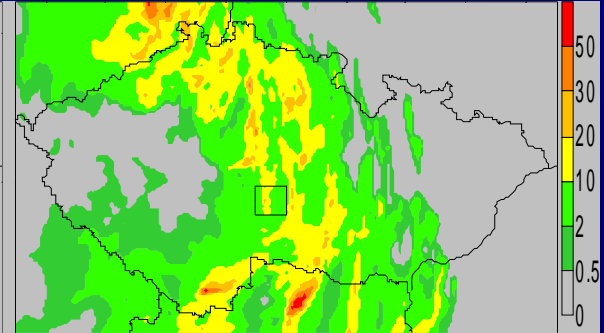
[T-1,T]



[T,T+3]

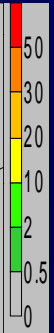
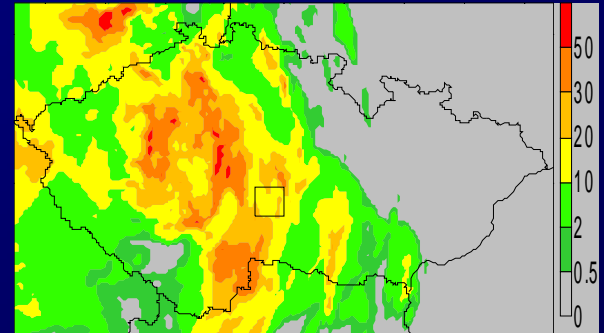


LM



LHN

WVC 2D

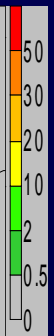
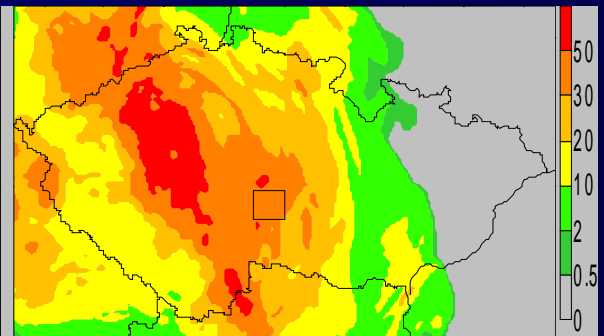
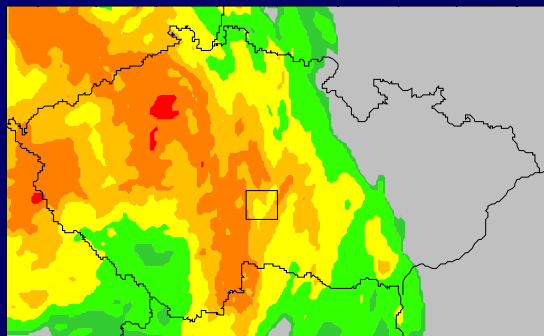


2D

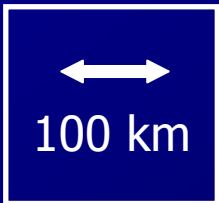
Forecasted precipitation with assimilation

WVC 3D

WVC VIL



3D



(ii) “Objective approach”

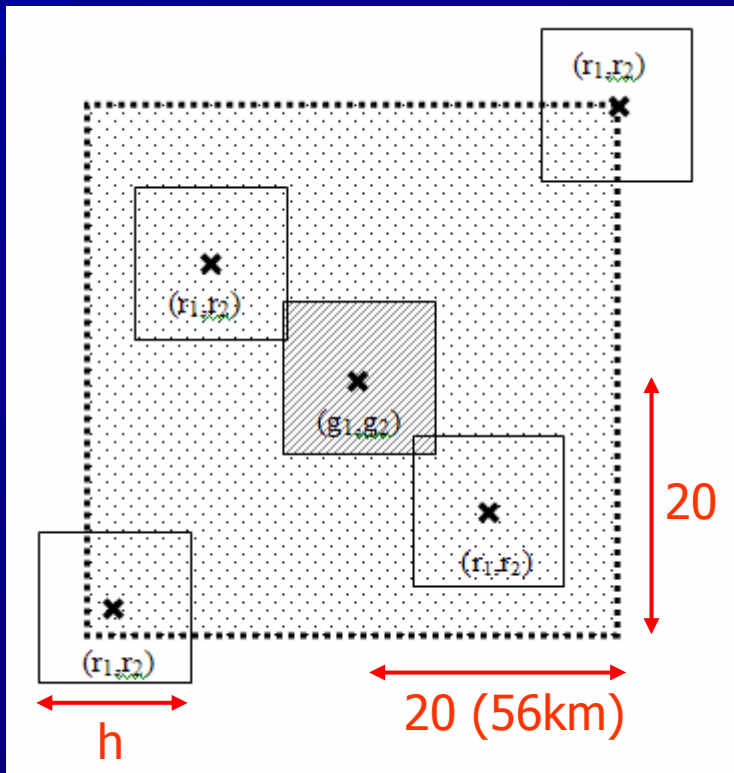
# SRMSE

## Assumptions:

The storm position as well as the structure are not expected to be forecasted accurately.

But we want:

- The forecasted storm should be close to the observed one.
- Distribution of precipitation values should be similar.



## Verification method

- Observed  $\{o_i\}$  and forecasted  $\{f_i\}$  grid values from squares  $h \times h$  are compared:

$$o_1 \leq o_2 \leq \dots$$

$$f_1 \leq f_2 \leq \dots$$

$$\text{RMSE}(o, f)$$

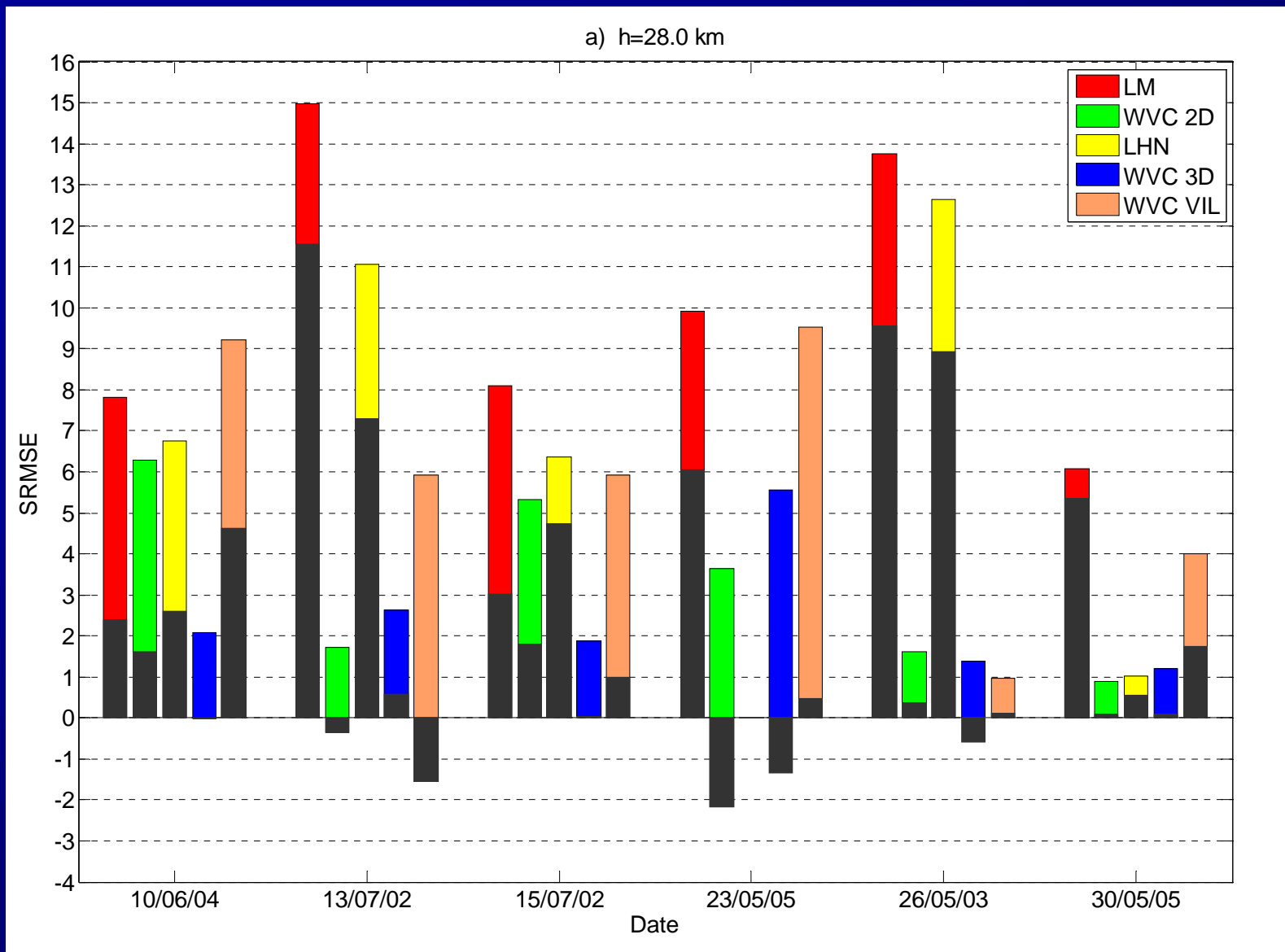
- Observed values are from the fixed squared in the centre  $(g_1, g_2)$

- Forecasted values are from various squares  $(r_1, r_2)$

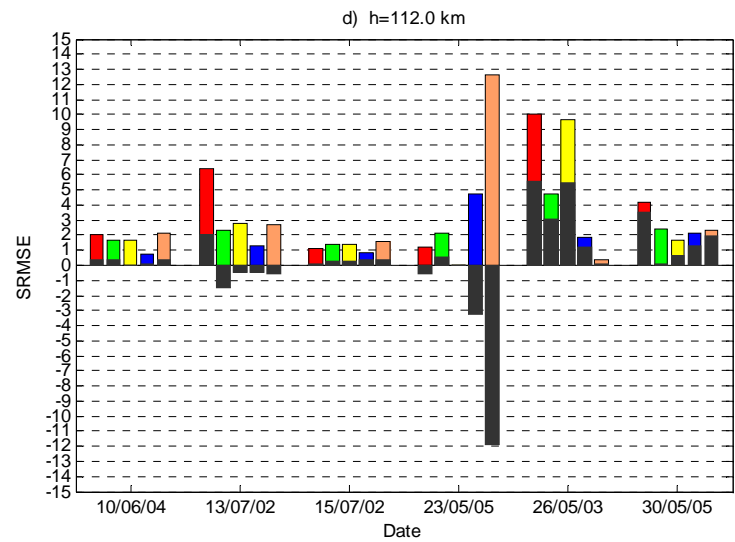
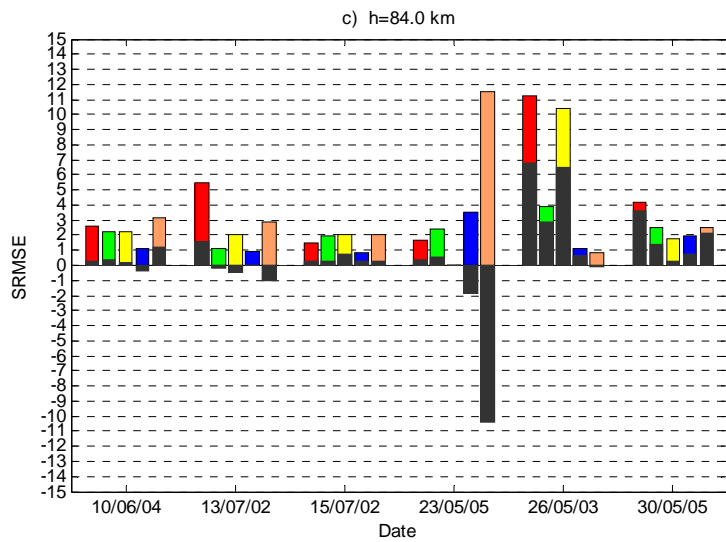
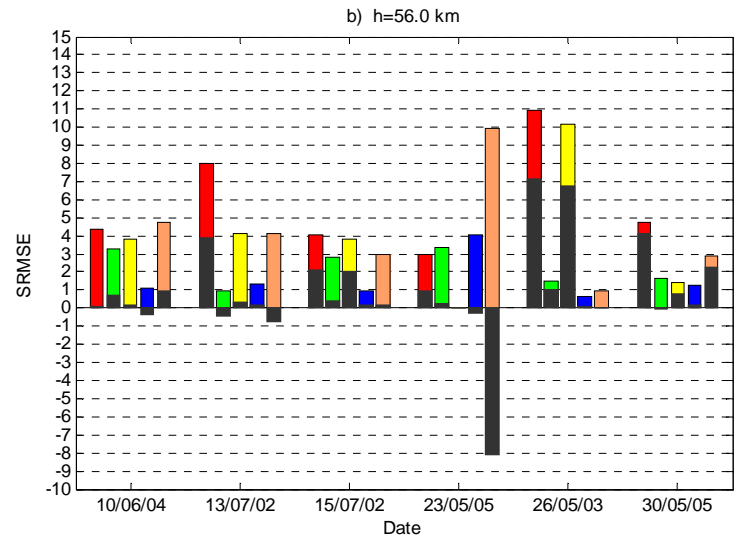
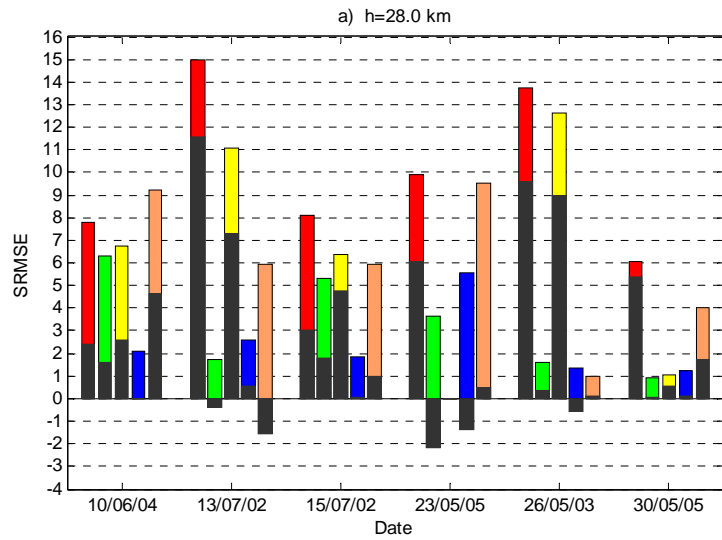
- $\text{SRMSE} = \min \{ \text{RMSE}(o, f) \}$

$$\text{abs}(r_1 - g_1) \leq 20, \text{abs}(r_2 - g_2) \leq 20$$

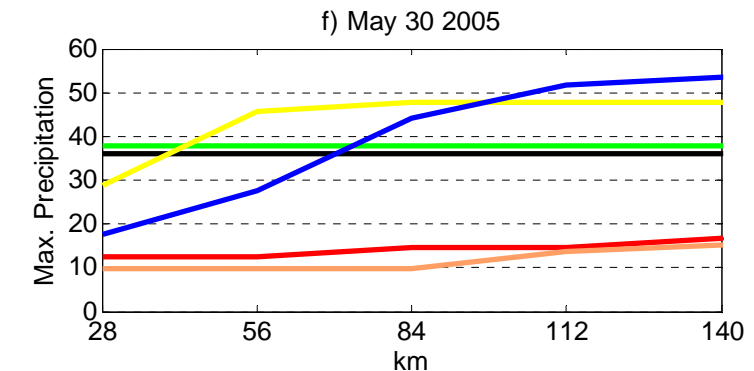
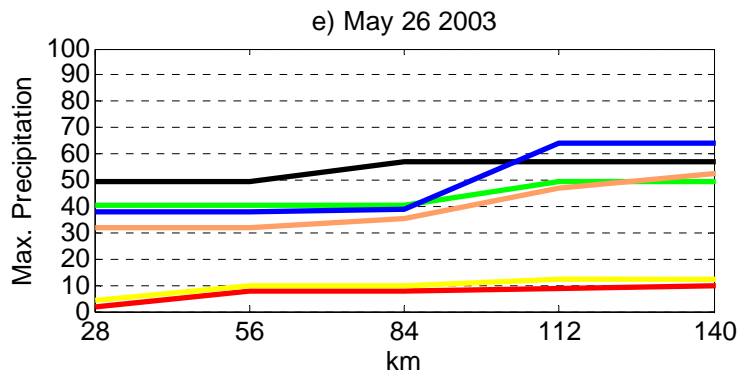
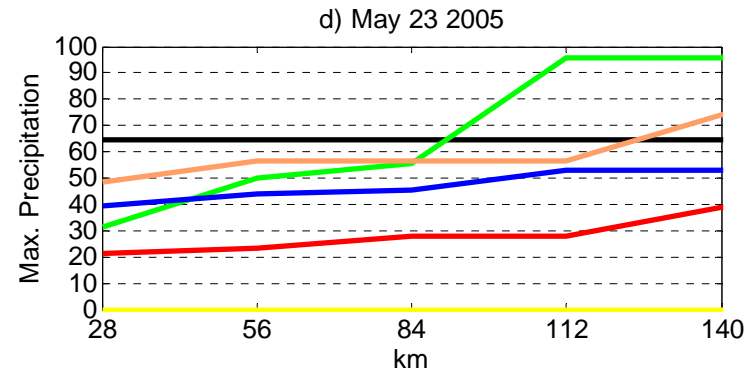
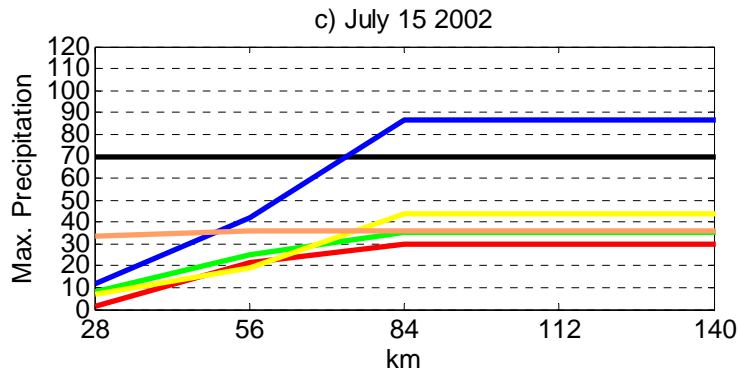
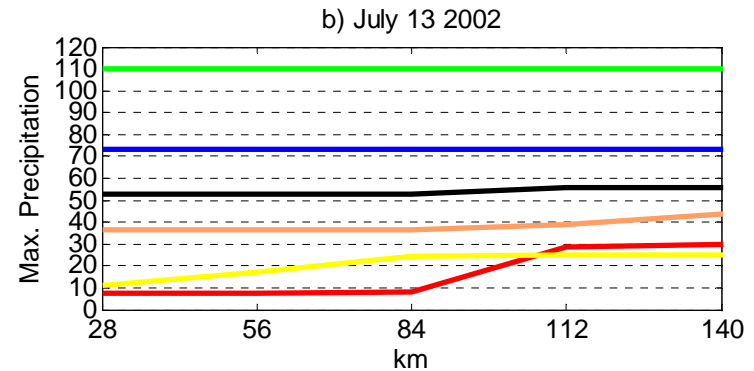
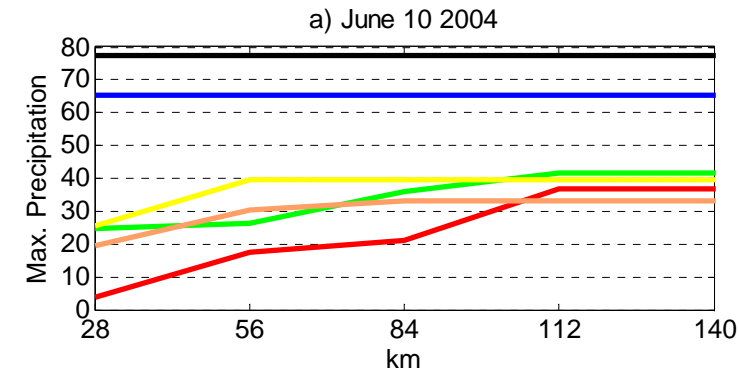
# SRMSE 28km x 28km



# SRMSE



# Maximum in squares $h=28, 56, 84, 112, 140$ km



— OBS — LM — WVC 2D — LHN — WVC 3D — VIL

# Conclusions

- The assimilation significantly improves the forecast.
- 2D methods:
  - LHN and CWV methods are comparable when LM forecast is good
  - If LM does not forecast precipitation and observed precipitation during the assimilation period is small then the LHN has problems to develop precipitation.
- 2D versus 3D methods:
  - Preliminary results
  - VIL is not good
  - WVC 3D seems slightly better
- The forecast is successful if first indications of the storm are observed by radars.
- Nowcasting of precipitation

# Thank you

Acknowledgement:

DWD (LM COSMO)

CHMI (radar and gauge data)