

Status of COSMO-ART & ICON-ART

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Development and applications of ICON-ART





Implementation of a point source



At a given point:		<pre><pntsrc id="RNDFACTORY"> <lon type="real">2.351667</lon> <lat type="real">48.856667</lat> <substance type="char">testtr</substance> <source_strength_type="real">1.0 <height type="real">1.0 <height type="real">1.0 <height type="real">2014-03-29T00:00:00 <starttime type="char">2014-03-29T00:00:00</starttime> <endtime type="char">2014-03-29T01:00:00</endtime> </height></height></height></source_strength_type="real"></pntsrc></pre>
Uniform emission profile:	ſ	<pre><pntsrc id="Eyjafjalla"> <lon type="real">-19.36</lon></pntsrc></pre>
		<lat type="real">63.63</lat> <substance type="char">testtr</substance>
Between surface and given height	4	<pre><source_strength type="reat">1.0</source_strength> <height type="reat">-6000. /height> <upit type="char">kg s=l</upit></height></pre>
		<pre><starttime type="char">2014-03-29T04:00:00</starttime> <endtime type="char">2014-03-29T05:00:00</endtime></pre>
	L	
Between given height_bot and height		<pre><pre><pre>clop type="real">=19 4</pre></pre></pre>
		type="real">63.59
		<substance type="char">testtr</substance>
	4	<source_strength type="real">1.0</source_strength>
		<pre><height bot="" type="seal">1491.</height></pre>
		<unit type="char">kg s-l</unit>
		<pre><starttime type="char">2014-03-29T00:00:00</starttime></pre>
		<pre><endtime type="char">2014-03-29T01:00:00</endtime></pre>

Implementation of a point source



Emission profile:

```
Edit View Bookmarks Settings Help
!DOCTYPE tracers SYSTEM "sources_selTrnsp.dtd">
<sources>
 <pntSrc id="Eyjaf1">
   <lon type="real">-19.62</lon>
   <lat type="real">63.63</lat>
   <substance type="char">ash_insol_acc</substance>
   <source strength type="real">500.0</source strength>
   <height type="real">9000.</height>
    <height bot type="real">1666.</height bot>
   <emiss profile type="char">0.0076 * [z_star] - 0.5 * sqrt(pi) * 0.9724 * 0.3078
                               * erf((0.4481 - [z star]) / 0.3078) / 0.524647415</emiss profile
   <unit type="char">kg s-1</unit>
   <startTime type="char">2010-04-03T00:00:00</startTime>
   <endTime type="char">2014-04-31T00:00:00</endTime>
 </pntSrc>
 <pntSrc id="Eyjaf2">
   <lon type="real">-19.62</lon>
   <lat type="real">63.63</lat>
   <substance type="char">ash_insol_coa</substance>
   <source_strength type="real">500.0</source_strength>
   <height type="real">-4000.</height>
   <emiss_profile type="char">0.5 / (4 * pi) * cos(2 * pi * [z_star] / 0.5) + [z_star]/emiss_profile>
   <unit type="char">kg s-1</unit>
   <startTime type="char">2010-04-03T00:00:00</startTime>
   <endTime type="char">2014-04-31T00:00:00</endTime>
 </pntSrc>
```

Examples for a point source







Implementation of vegetation fire emissions



Portugal 2017







https://worldview.earthdata.nasa.gov

Plume rise model (Freitas et al. 2006, Walter et al. 2014)











Overview of ice nucleation onset temperatures and saturation ratios





Hoose and Möhler, Atmos. Chem. Phys. Atmos. Chem. Phys., 12, 9817-9854, https://doi.org/10.5194/acp-12-9817-2012, 2012





m Contour plot - SPP - Height: 8000 m



Dust particle shapes

е









Models



Dust forecast models usually assume spherical dust. But spheres fail to reproduce the magnitude and angular distribution of the scattering.







Nousiainen and Kandler 2015

Dust particle shapes: optical properties



- **Tri-axial ellipsoids** better reproduce the laboratory measurements (Meng et al 2010).
- We use a mixture of 35 ellipsoid shapes with aspect ratio of 1.1 to 3.3.



Verification: AB 1064 nm on 30.08.2017







- Improvement of ABS for NSP particles
- Minor changes for AOD
- Particle size distribution seems to be more important in case of AOD

Hoshyaripour, G., Bachmann, V., Förstner, J., Steiner, A., Vogel, H., Wagner, F., Vogel, B. Accounting for Particle Non-sphericity in a Dust Forecast System: Impacts on Model-Observation Comparison, submitted to JGR



Climate Engineering by Arctic Winter CirrusThinning



after Storelvmo et al., 2013

Optical Properties of Hydrometeors







- 象 R2B09 (~5 km)
- two-moment microphysics: Seifert and Beheng (2006)
- 🐼 cloud optical properties: Fu et al. (1998), Fu (2007), Hu and Stamnes (1993)
- Inucleation: Barahona and Nenes (2009)
- heterogeneous nucleation: Phillips, et al. (2013)
- activation of CCN: Bangert et al. (2012)



Calipso and Halo flight tracks





Validation: Calipso





Validation: POLSTRACC, GLORIA





 H_2O

 $\mathbf{RH}_{\mathrm{ice}}$



Climate Engineering by Arctic Winter CirrusThinning



after Storelvmo et al., 2013

Gruber, S., U. Blahak, F. Haenel, Ch. Kottmeier, Th. Leisner, H. Muskatel, T. Storelvmo, and B. Vogel,

A process study on thinning of Arctic winter cirrus clouds with high-resolved ICON-ART simulations, in preparation

Applications of COSMO-ART





Impact of aerosols on clouds and atmospheric dynamics over southern West Africa





Deetz et al., 2018







Abbreviation	Description of Simulation
AIE _{0.1} ADE _{0.1}	$F_{AIE} = 0.1$ and $F_{ADE} = 0.1$
$AIE_{0.25}ADE_{0.25}$	$F_{AIE} = 0.25$ and $F_{ADE} = 0.25$ (clean case)
AIE _{0.5} ADE _{0.5}	$F_{AIE} = 0.5$ and $F_{ADE} = 0.5$
AIE _{1.0} ADE _{1.0}	$F_{AIE} = 1.0$ and $F_{ADE} = 1.0$ (reference case)
AIE _{2.0} ADE _{2.0}	$F_{AIE} = 2.0$ and $F_{ADE} = 2.0$
AIE _{4.0} ADE _{4.0}	$F_{AIE} = 4.0$ and $F_{ADE} = 4.0$ (polluted case)

Impact on radiation





Impact on physical properties and precipitation





Impact on shortwave radiation at surface







Aerosol strongly effects cloud effective radii.

LWC and precipitation show no significant change.

Strong impact of modified clouds on radiation.

Strong contribution from biomass burning in central Africa.

K. Deetz, H. Vogel, P. Knippertz, B. Adler, J. Taylor, H. Coe, K. Bower, S. Haslett, M. Flynn, James Dorsey, lan Crawford, Christoph Kottmeier, B. Vogel
Numerical simulations of aerosol radiative effects and their impact on clouds and atmospheric dynamics over southern West Africa, Atmos. Chem. Phys., 18, 9767-9788, https://doi.org/10.5194/acp-18-9767-2018, 2018
K. Deetz, H. Vogel, S. Haslett, P. Knippertz, H. Coe, B. Vogel
Aerosol liquid water content in the moist southern West African monsoon layer and its radiative impact

Atmos. Chem. Phys. Discuss., https://doi.org/10.5194/acp-2018-420, 2018

KIT-ART-Team







Current Projects



Research themes

- Impact of volcanic ash on atmospheric processes
- Dust-cloud-radiation feedback
- Scale dependency of aerosol cloud interaction
- Climate engineering
- Biomass burning aerosol
- Impact of sub pollen particles
- Emission driven annual cycles
- Chemistry-climate interactions, including PSCs
- Water isotopologues (weather/climate)
- Composition assimilation



Additional material Africa simulations





Comparison to observations







Additional material time spitting


Microphysics sensitivity to timestep



Model setup

- Idealised 2-hour simulation using COSMO v5.3
- 1-km resolution; 64 vertical levels; timestep 1-20 s
 - Weisman-Klemp thermodynamic profile; 2K warm bubble; linear shear
- Seifert & Beheng 2-moment microphysics Segal & Khain CCN activation
 - Two different aerosol settings:
 - clean = 100 CCN cm⁻³; continental = 1700 CCN cm⁻³

Total precipitation: aerosol and timestep effects





Why? Numerics!



Model dynamics calculated before microphysics



- Storm updraft -> adiabatic cooling
- cooling -> supersaturation
- Cooling = rate $x \Delta t$
- Supersaturation depends on timestep
- Microphysical processes now calculated using timestepdependent supersaturation
- Many processes affected
- Timestep-dependent results



Aerosols Trace Gases and Climate Processes

Timestep dependence in a simple model C = "Dynamics" dq_c Updraft q_c dt🗌 cooling **Autoconversion** condensation (microphysics) Condensation Aq_{c}^{2} = "simplified microphysics" **Consecutive Splitting** 0.07 (a) 0.06 Cloud water content [g kg⁻¹] 0.05 0.04 **ال**رة Cdt q_c 0.03 1s 6s 0.02 2s 7s 3s 8s 0.01 4s 9s 5s 10s 0.00 0 25 50 75 100 125 150 175 200 Time [secs]

COSMO with Simultaneous Splitting







Summary

simulations

Large and systematic effect of model timestep on convection-permitting

50% increase in precipitation50km change of maximum precipitation changing sign of aerosol impact

Caused by "Consecutive Splitting"
 Dynamics calculated first, then microphysics

- Supersaturation (or q_c) scales with timestep
- Results much better with "Simultaneous Splitting"



Affecting convection-permitting simulations ... in most (all?) models

Also affects NWP and process-studies

Solution: Changing input for microphysics – easy to change in model Andrew Barrett (andrew.barrett@kit.edu)
Acrosols Trace Gases and Climate Processes



Additional material aerosol module

Milestones







Aerosol, radiation and cloud physics

Convective uplift

Chemistry-climate feedbacks





The basic equation

$$\frac{\partial n(v_{p})}{\partial t} = -\underbrace{\nabla \cdot \mathbf{v} n(v_{p})}_{Advection} - \underbrace{\nabla \cdot \mathbf{c}_{p}(v_{p}) n(v_{p})}_{External \ forces} + \underbrace{\nabla \cdot D_{PAR}(v_{p}) \nabla n(v_{p})}_{Diffusion} + \frac{1}{2} \int_{0}^{v_{p}} \beta(v_{p} - \widetilde{v}_{p}, \widetilde{v}_{p}) n(v_{p} - \widetilde{v}_{p}) n(\widetilde{v}_{p}) d\widetilde{v}_{p} - \int_{0}^{\infty} \beta(v_{p}, \widetilde{v}_{p}) n(v_{p}) n(\widetilde{v}_{p}) d\widetilde{v}_{p}}_{Coagulation} + \underbrace{\left[\frac{\partial}{\partial t} n(v_{p})\right]_{g}}_{Particle \ growth} + \underbrace{\frac{\dot{n}_{s}(v_{p})}{Sources/\sin}}_{Sources/\sin}$$

Friedlander (1977)





Atmospheric transport (advection, convection, turbulent diffusion)

Sedimentation

Washout

Emission

(sea salt, mineral dust, volcanic ash, pollen, radioactive material)

Optical properties





\odot Condensation (explicit of H₂SO₄)

Nucleation

Gas-aerosol partitioning



	SA	IA	MA	Sa	la	Ma	Sc	lc	Mc
SA	SA	MA	MA	Sa	Ma	Ma	Sc	Mc	Mc
IA		IA	MA	Ma	la	Ma	Mc	Ic	Mc
MA			MA	Ma	Ma	Ma	Mc	Mc	Mc
Sa				Sa	Ma	Ma	Sc	Mc	Mc
la					la	Ma	Mc	Ic	Mc
Ма						Ma	Mc	Mc	Mc
Sc							Sc	Mc	Mc
lc								Ic	Mc
Mc			Î						Mc

e.g. intermodal coagulation:

$$Ca_{0,ij} = \int_0^\infty \int_0^\infty \beta(d_1, d_2) n_i(d_1) n_j(d_2) \, \mathrm{d}d_1 \mathrm{d}d_2$$



free molecular regime:



$$\begin{aligned} \operatorname{Ca}_{0,ij}^{\mathrm{fm}} &= \int_{0}^{\infty} \int_{0}^{\infty} \beta_{\mathrm{fm}}(d_{1}, d_{2}) n_{i}(d_{1}) n_{j}(d_{2}) \, \mathrm{d}d_{1} \mathrm{d}d_{2} \\ &= M_{0,i} M_{0,j} K_{\mathrm{fm}} b_{0}^{(1)} \sqrt{d_{gi}} \Biggl[e^{\frac{1}{8} ln^{2}(\sigma_{gi})} + \sqrt{\frac{d_{gj}}{d_{gi}}} e^{\frac{1}{8} ln^{2}(\sigma_{gj})} \\ &+ 2 \frac{d_{gj}}{d_{gi}} e^{\frac{1}{8} ln^{2}(\sigma_{gi})} e^{\frac{4}{8} ln^{2}(\sigma_{gj})} + \frac{d_{gj}^{2}}{d_{gi}^{2}} e^{\frac{9}{8} ln^{2}(\sigma_{gi})} e^{\frac{16}{8} ln^{2}(\sigma_{gj})} \\ &+ \left(\sqrt{\frac{d_{gi}}{d_{gj}}} \right)^{3} e^{\frac{16}{8} ln^{2}(\sigma_{gi})} e^{\frac{9}{8} ln^{2}(\sigma_{gj})} + 2 \sqrt{\frac{d_{gi}}{d_{gj}}} e^{\frac{4}{8} ln^{2}(\sigma_{gi})} e^{\frac{1}{8} ln^{2}(\sigma_{gj})} \Biggr] \end{aligned}$$



$$\frac{\partial C_{SO_4,l}}{\partial t} = \frac{\pi}{6} \cdot G_l^3 \cdot C_{H_2SO_4}(t)$$

$$G_{l}^{3} = \frac{G_{l,fm}^{3} \cdot G_{l,nc}^{3}}{G_{l,fm}^{3} + G_{l,nc}^{3}}$$

$$G_{l,fm}^3 = \frac{6}{\pi} \frac{\pi \alpha \bar{c}}{4} M_l^2$$
$$G_{l,nc}^3 = \frac{6}{\pi} 2\pi D_v M_l^1$$



near continuum regime:

$$\begin{aligned} \operatorname{Ca}_{0,ij}^{\operatorname{nc}} &= \int_{0}^{\infty} \int_{0}^{\infty} \beta_{\operatorname{nc}}(d_{1}, d_{2}) n_{i}(d_{1}) n_{j}(d_{2}) \mathrm{d}d_{1} \mathrm{d}d_{2} \end{aligned}$$
$$&= M_{0,i} M_{0,j} K_{\operatorname{nc}} \left[2 + A_{i} \operatorname{Kn}_{g_{i}} \left(e^{\frac{4}{8} \ln^{2}(\sigma_{g_{i}})} + \frac{d_{gj}}{d_{gi}} e^{\frac{16}{8} \ln^{2}(\sigma_{gi})} e^{\frac{4}{8} \ln^{2}(\sigma_{gj})} \right) + A_{j} \operatorname{Kn}_{g_{j}} \left(e^{\frac{4}{8} \ln^{2}(\sigma_{gj})} + \frac{d_{gi}}{d_{gj}} e^{\frac{16}{8} \ln^{2}(\sigma_{gj})} e^{\frac{4}{8} \ln^{2}(\sigma_{gi})} \right) + \left(\frac{d_{gi}}{d_{gj}} + \frac{d_{gj}}{d_{gj}} \right) \left(e^{\frac{4}{8} \ln^{2}(\sigma_{gj})} \right) \left(e^{\frac{4}{8} \ln^{2}(\sigma_{gi})} \right) \right] \end{aligned}$$



- Calculation of the loss of sulfuric acid on already existing particles
- Calculation of a critical concentration according to Wexler (1994).

$$c_{crit} = 0.16 \cdot \exp \begin{bmatrix} 0.1 \cdot T - 3.5 \frac{RH}{100} - 27.7 \end{bmatrix}$$

Remaining mass above c_{crit} nucleates and forms new particles

Gas-aerosol partitioning





ISORROPIA, Nenes, Fountoukis

Simulated SO₂ and NH₃ concentration





Sulfate and ammonia concentration













New aerosol module for ICON-ART developed and realized (testing phase)

Mode structure allows large range of complexity:

reduced aerosol module for NWP-applications detailed aerosol module for research

Seamless in the vertical direction











Observations in the Asian Monsoon July 2017

Very high values (> 1 ppbv) of NH_3 measured (12 km - 14 km)

 \sim 40 times higher than maximum NH₃ values measured by MIPAS-Envisat

Inhomogeneous horizontal distribution





Sören Johansson for the GLORIA Team





Comparison with observations



GLORIA NH₃ 2017-07-29

ICON-ART NH₃ 2017-07-29



Sören Johannson & Michael Höpfner

Poster: Carmen Ullwer et al., Investigation of the distribution of aerosol-forming trace gases in the UTLS region with ICON-ART



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- ¹ KIT, Institute of Meteorology and Climate Research – Troposphere Research
- ² KIT, Institute of Meteorology and Climate Research
 - Atmospheric Trace Gases and Remote Sensing
- ³ Deutscher Wetterdienst (DWD)
- ⁴ KIT, Steinbuch Centre for Computing

What makes ICON and ICON-ART unique?



- Seamless in horizontal and vertical scales (troposphere-mesosphere)
- Seamless in time (seconds-decades): (LES) Weather Climate







Development and implemention of forward operators of

natural and anthropogenic aerosol

Vertical profile of attenuated backscatter





Link between observation and simulation



$$P(z) = C_L \frac{\beta(z)}{z^2} exp\left(-2\int_0^z \alpha(z')dz'\right)$$

Frequently the Lidar ratio S is used to derive the backscatter coefficient

$$\beta(\lambda) = \frac{\alpha(\lambda)}{S}$$



Preferable: Direct calculation of the backscatter coefficient

Some hypothesis and evidence



- Evidence of biomass burning plumes in central and south Africa from June to September.
- According to Mari et al, 2008, there is a long-range biomass transport which is carried Westward by a jet at 700 hPa between 2 to 4 km
- High-pressure region to the West of African continent might play a role to mix it into the boundary layer. (Haslett, 2017)



Organic aerosol wo/w vegetation fire





Impact on CDN and r_e





Conceptual model





- Clouds bring aerosols from aloft into PBL.
- 7.46% increase of CDNC over the DACCIWA region and 32.92% over the marine domain.
- **50 w/m2** decrease in direct surface incoming SW radiation
- Mass concentration flux rate can reach **5 μg m-**² **s**⁻¹.
Dust optical properties in ICON-ART



Prognostic (for direct radiative effect):



Diagnostics (for AOD and AB):



Verification: Mean AOD July 2017







Impact of particles on radiation





Calculated optical properties







- Particle size distribution (emitted and transported)
- Variable median diameter and its impacts on optical props
- Dynamic land surface properties for emissions
- Parameterization of convective dust emission

Hoshyaripour, G., Bachmann, V., Förstner, J., Steiner, A., Vogel, H., Wagner, F., Vogel, B. Accounting for Particle Non-sphericity in a Dust Forecast System: Impacts on Model-Observation Comparison, submitted to JGR



Operational pollen forecast at MeteoSwiss (web page and app!)

Simulation of air quality in the area of Karlsruhe and comparison with detailed measurements (IMK-AAF)

🎯 empa

 Investigation of timestep dependency of aerosol-cloudconvection modelling study results:
Barrett et al 2018: "One Step at a Time: How Model Timestep Significantly Affects Convection-Permitting Simulations" submitted to Journal of Advances in Modelling Earth Systems (JAMES)