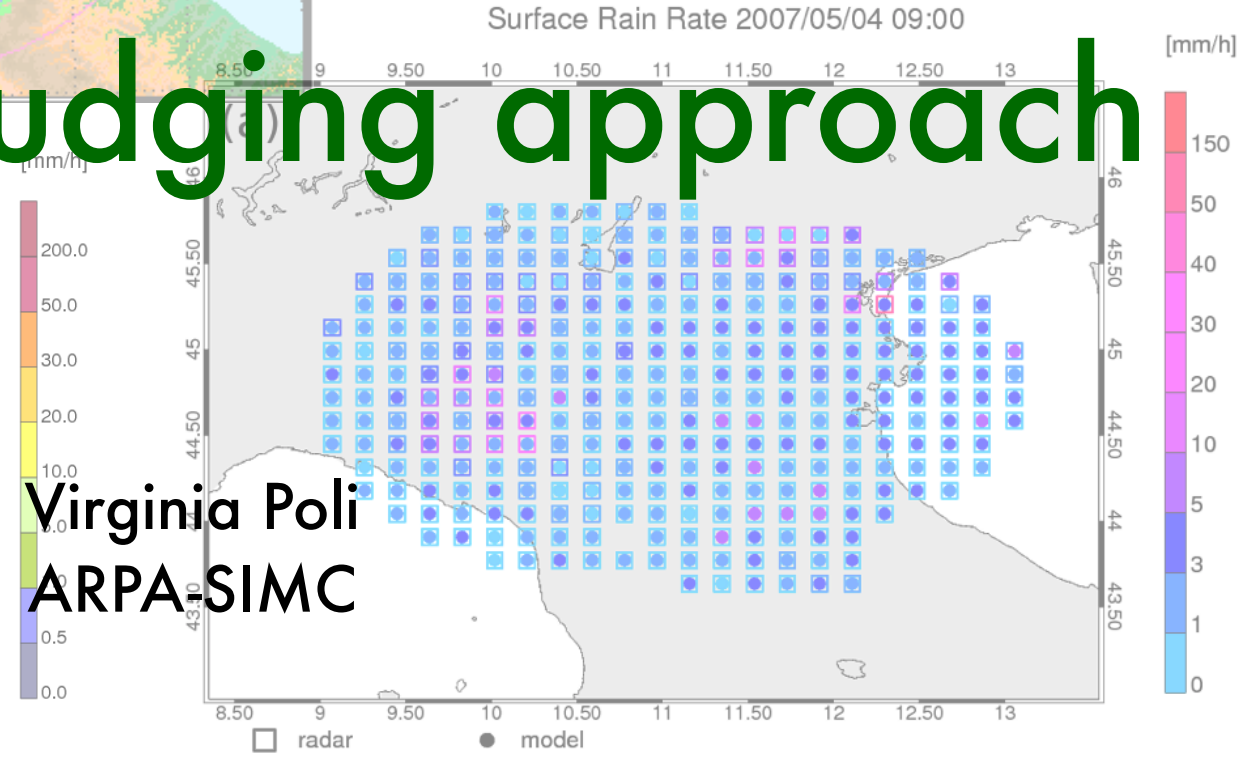
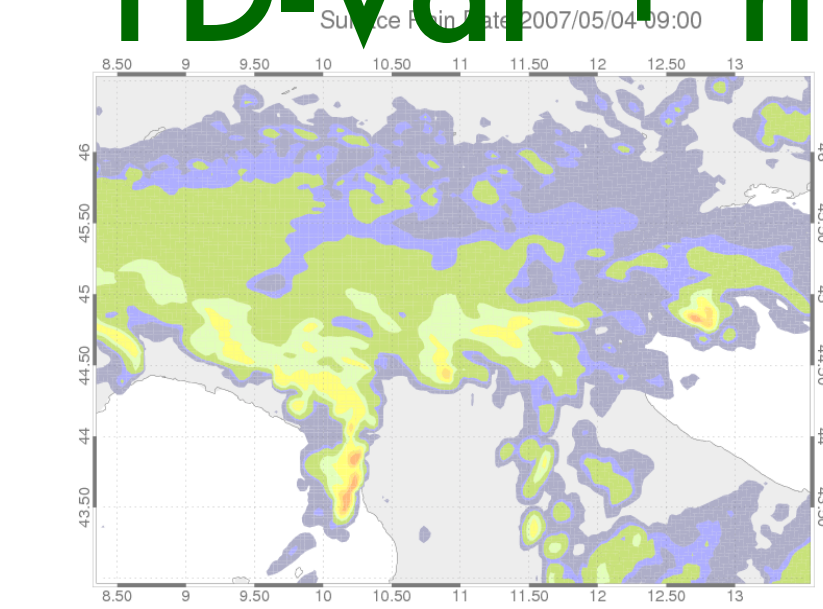
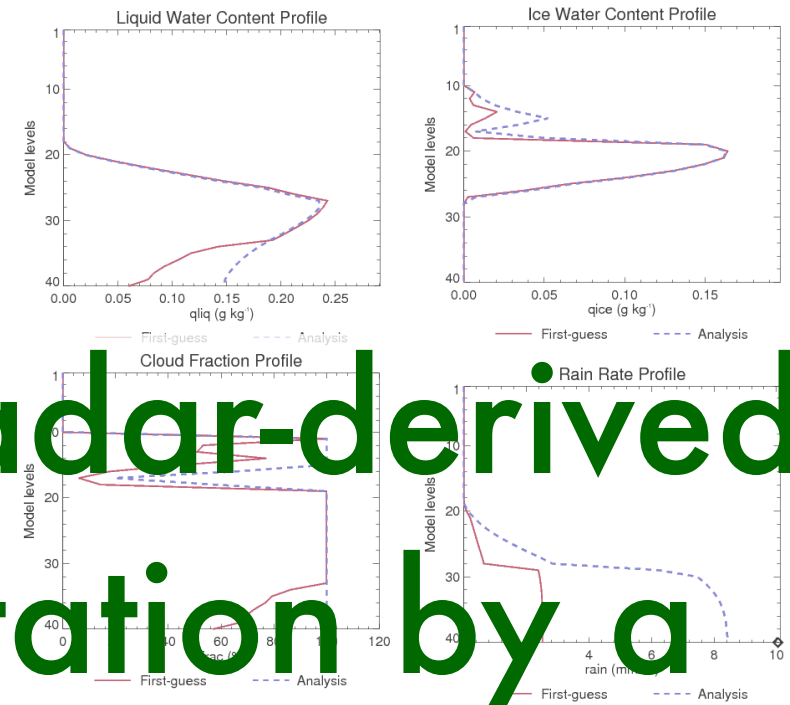
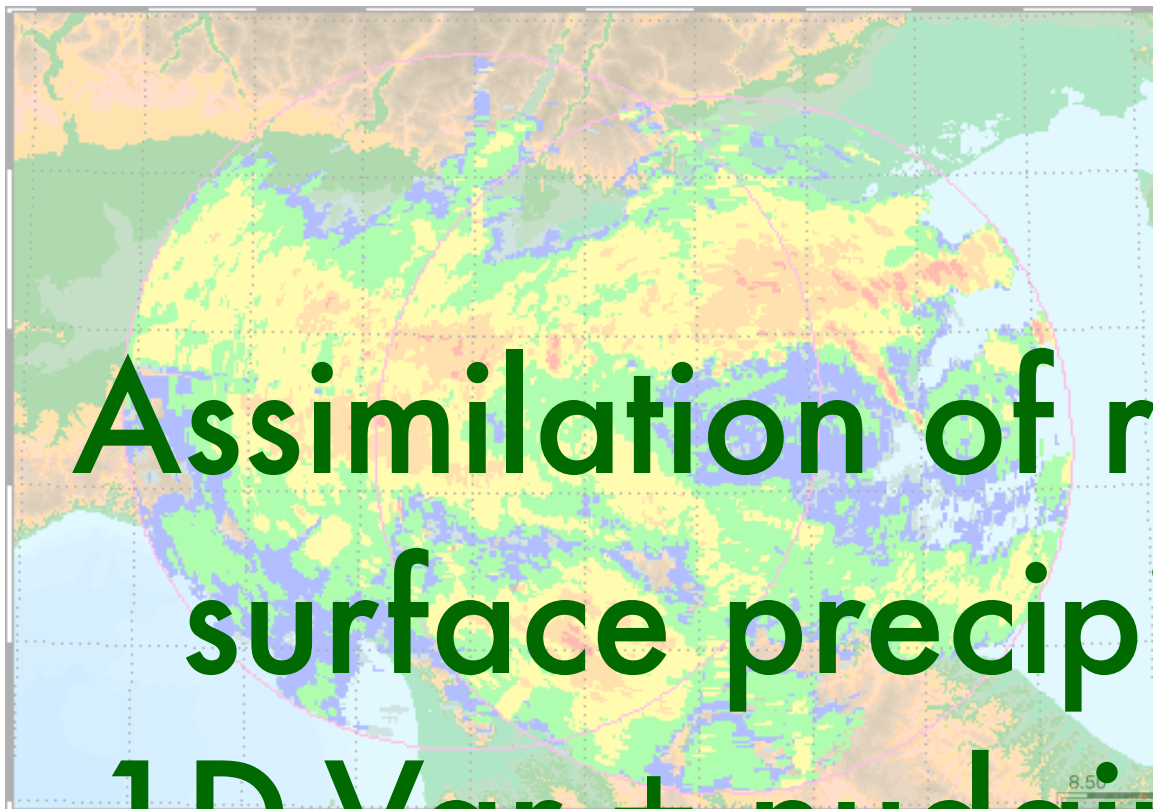


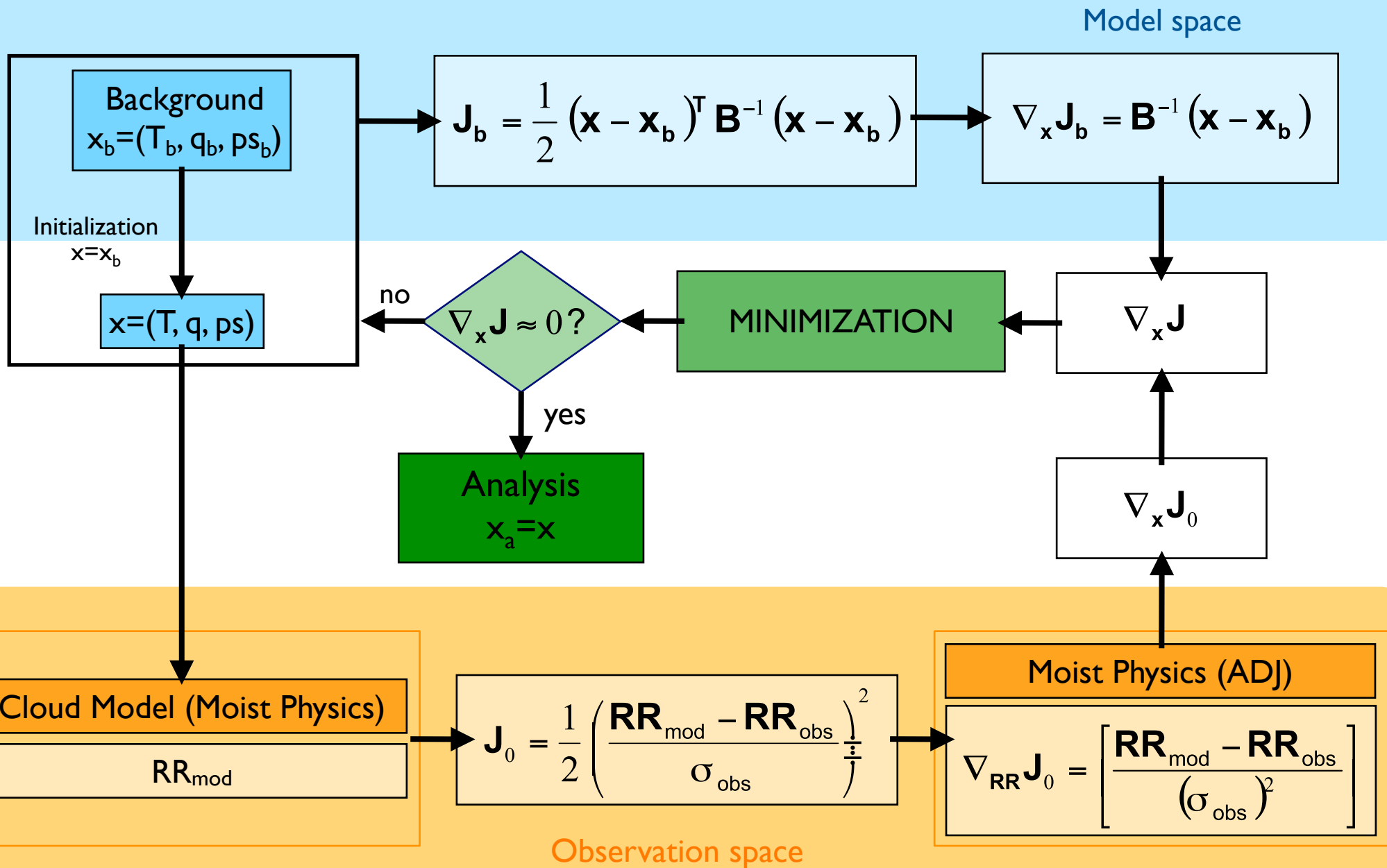
Assimilation of radar-derived surface precipitation by a 1D-Var + nudging approach



Virginia Poli
ARPA-SIMC

Variational assimilation

Convert observations (radar RR) in profiles of temperature and humidity and nudge them as “pseudo”-observations.

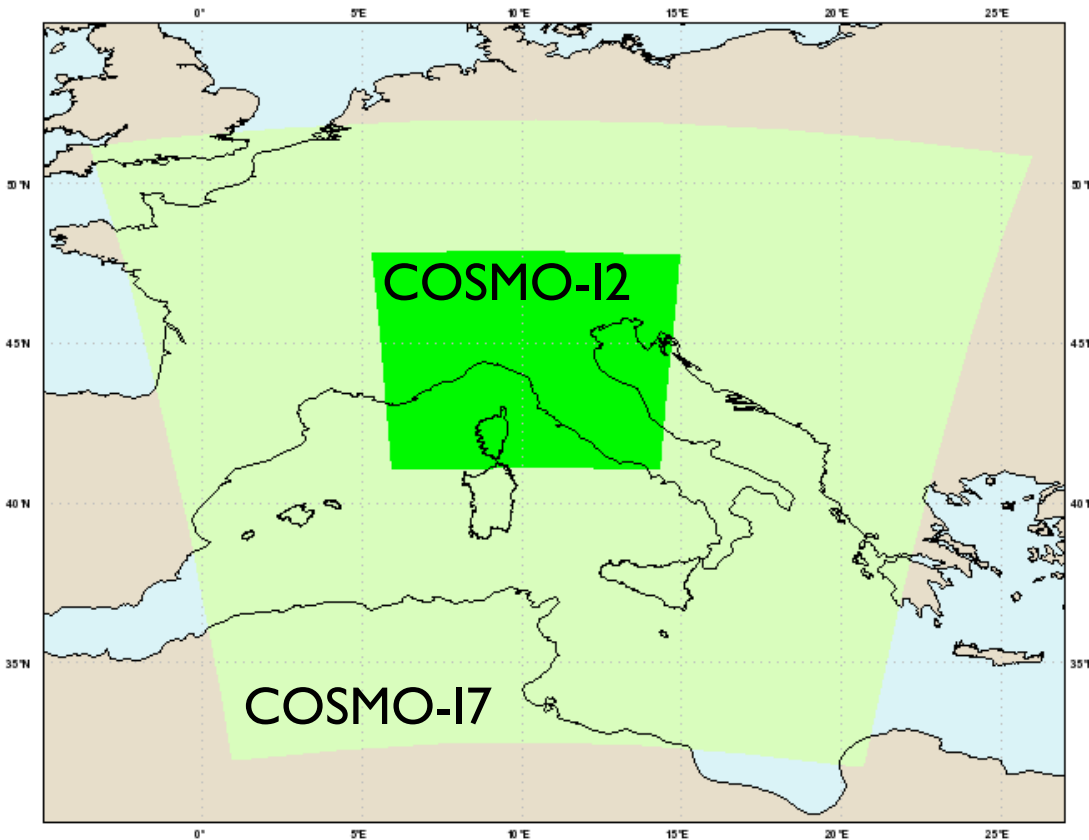


Numerical Model

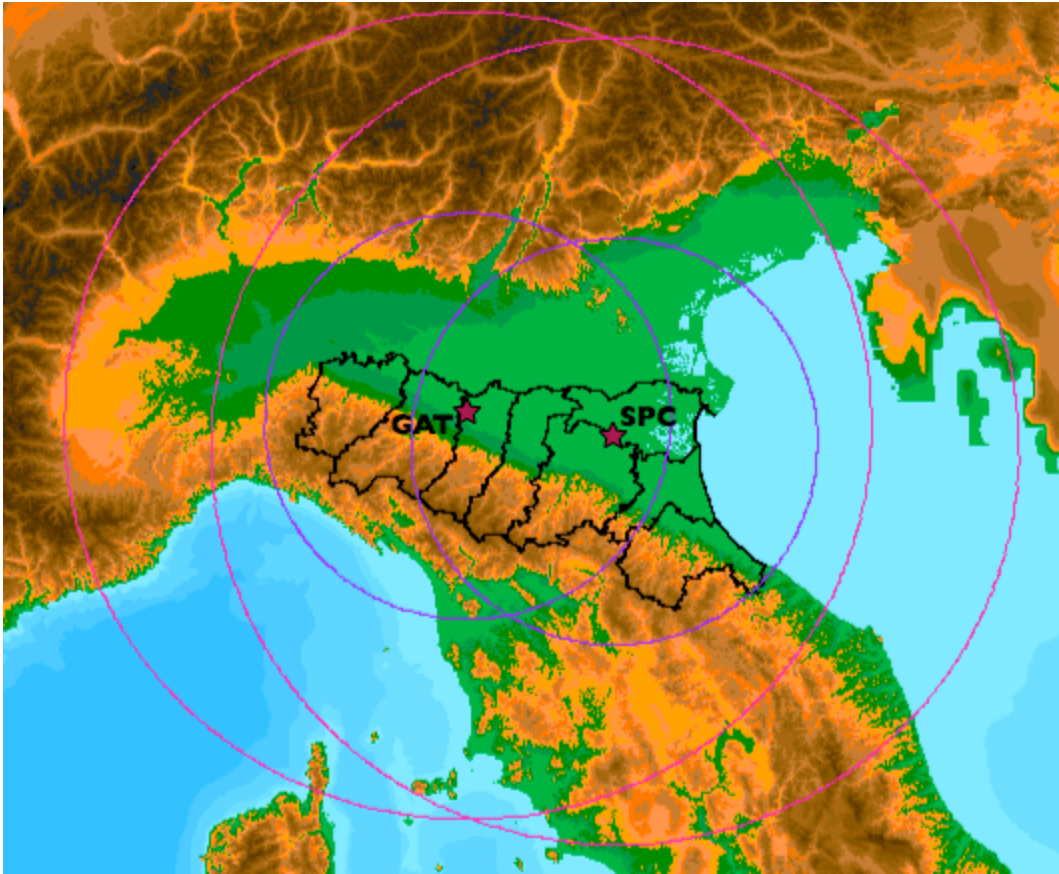
Numerical weather prediction model:

- COSMO-I2, version 4.0 (grid resolution = 2.8 km)
- nested on COSMO-I7 domain (grid resolution = 7 km)
- observations assimilation through nudging

Every 15 minutes, in the assimilation cycle, some parameters are selected (temperature profiles, humidity profiles, pressure profiles, sensible and latent heat fluxes at the surface) to be used as input in the variational algorithm.



Observations



Radar rain rates are derived from the two Emilia-Romagna polarimetric doppler C-band radars. Reflectivity measures are acquired from each radar with a repetition cycle of 15 minutes and a horizontal resolution of 1 km. Reflectivities values are then converted in instantaneous surface rain rate. The observational error has been estimated, as a first approximation, at about 30% of the measure.

The use of very high spatial and temporal data resolution should guarantee some improvements in the initial condition knowledge. Moreover, highly correlated observations can introduce not expected errors in the analysis, generating spurious structures.

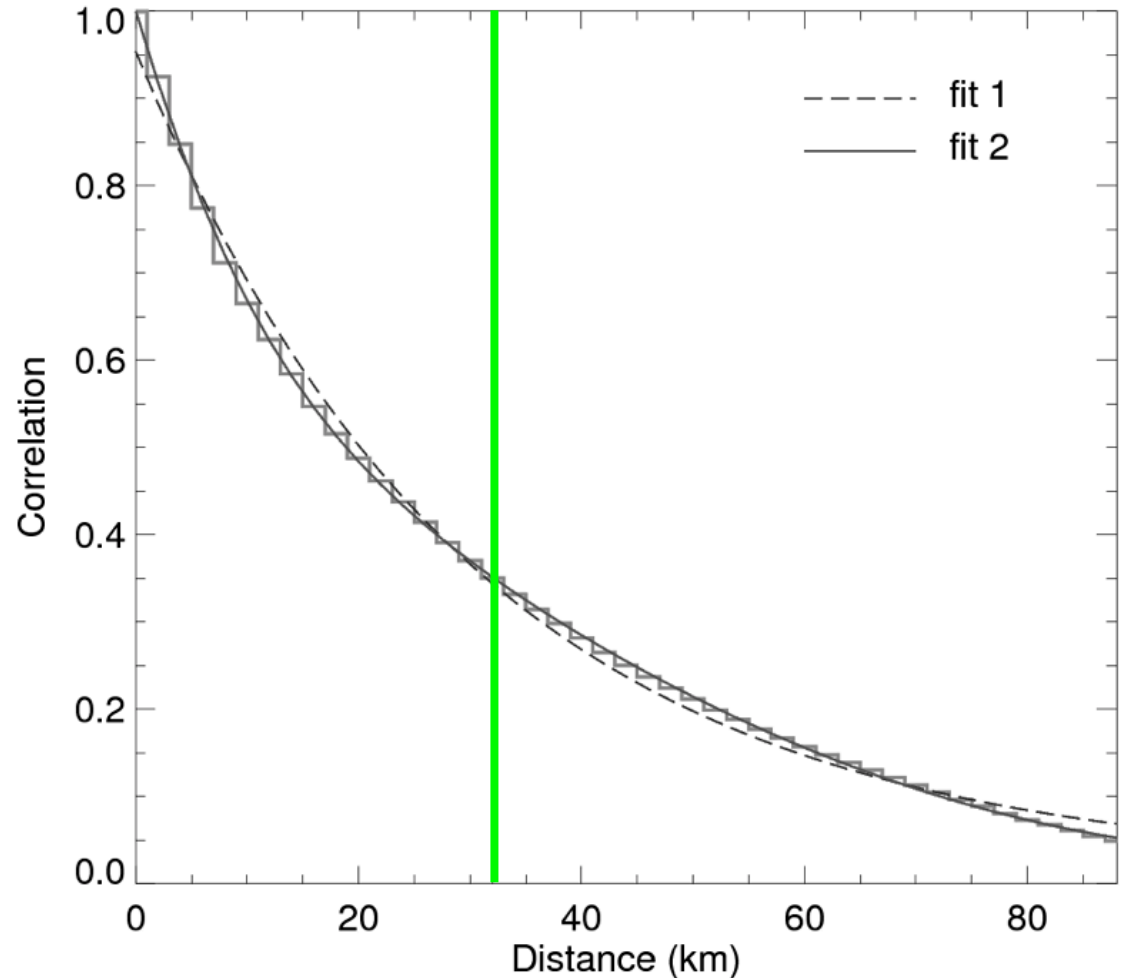
To reduce the total amount of data a **thinning** procedure is chosen.

Data thinning (I)

The spatial sampling is determined by the decorrelation length of the reflectivity auto-correlation field.

Using the data acquired over the whole test case period the auto-correlation histogram as a function of the distance has been generated. Smaller the distance higher the correlation between two points, being the correlation at zero distance equal to one by definition.

The trend is a monotonically decrease with no clear cut-off. The e-folding distance of 32 km could be chosen as decorrelation length as first approximation.

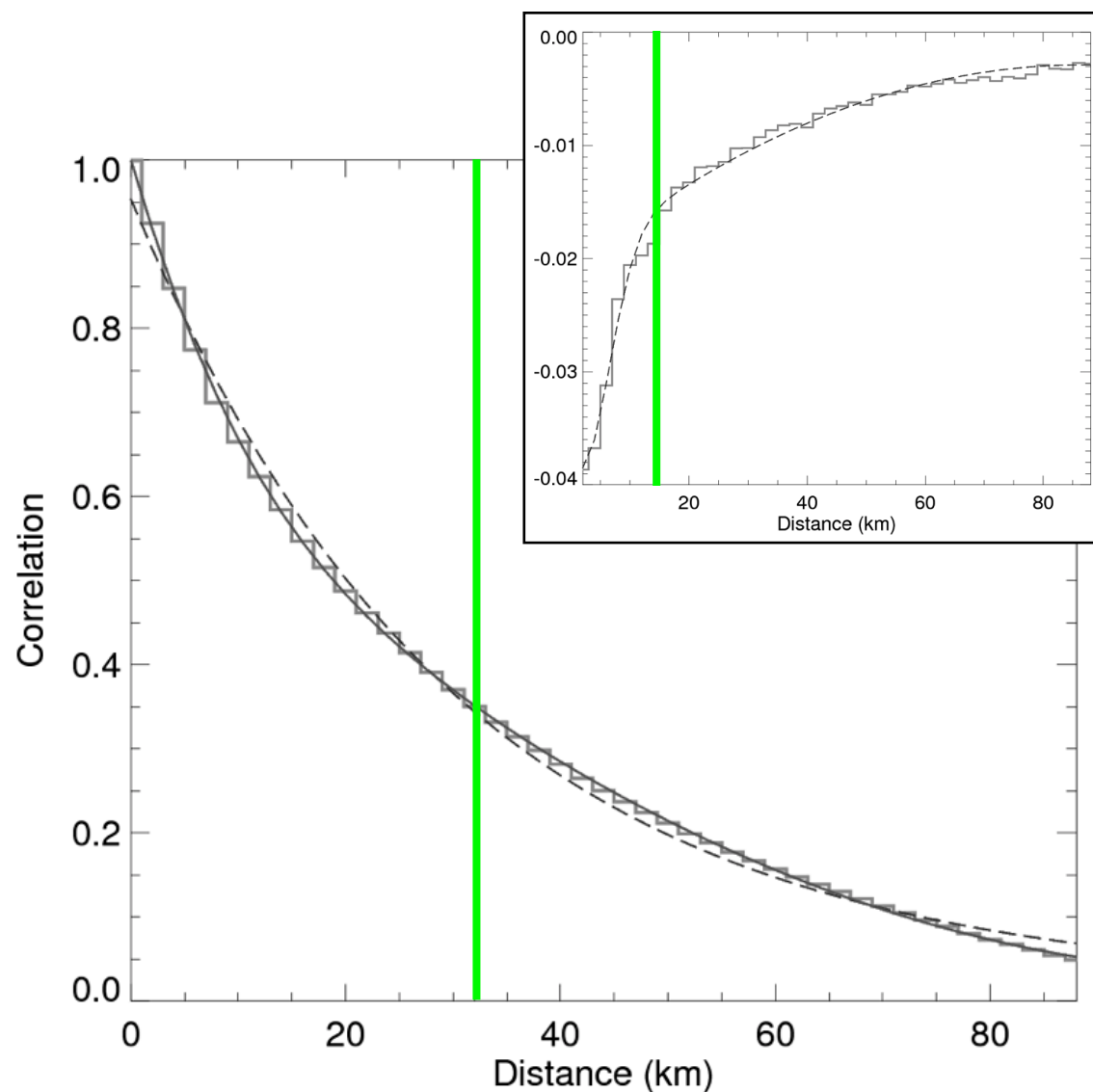


Data thinning (II)

The spatial sampling is determined by the decorrelation length of the reflectivity auto-correlation field.

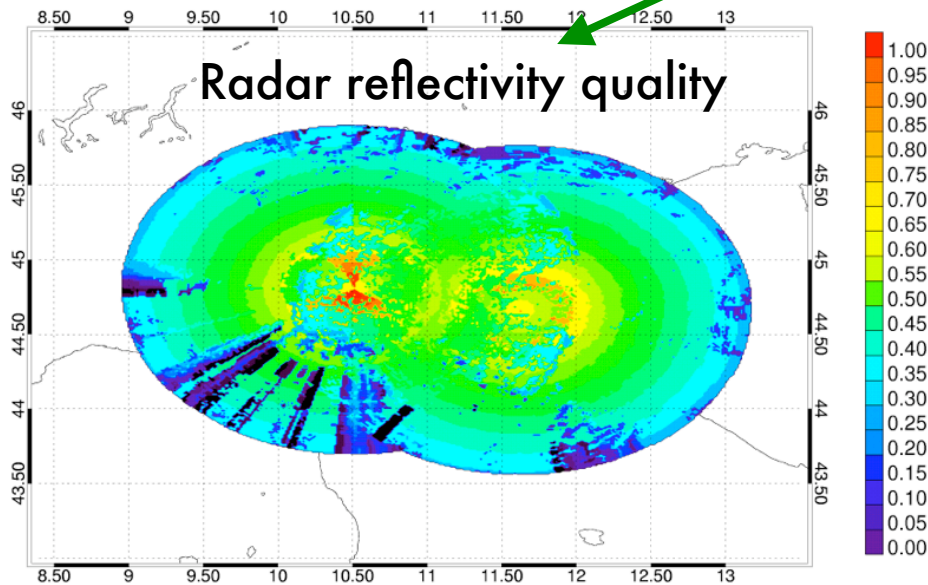
Using the data acquired over the whole test case period the auto-correlation histogram as a function of the distance has been generated. Smaller the distance higher the correlation between two points, being the correlation at zero distance equal to one by definition.

The trend is a monotonically decrease with no clear cut-off. The e-folding distance of 32 km could be chosen as decorrelation length as first approximation. **The derivative of the auto-correlation reveals discontinuity at about 15 km, which is chosen as new decorrelation length.**



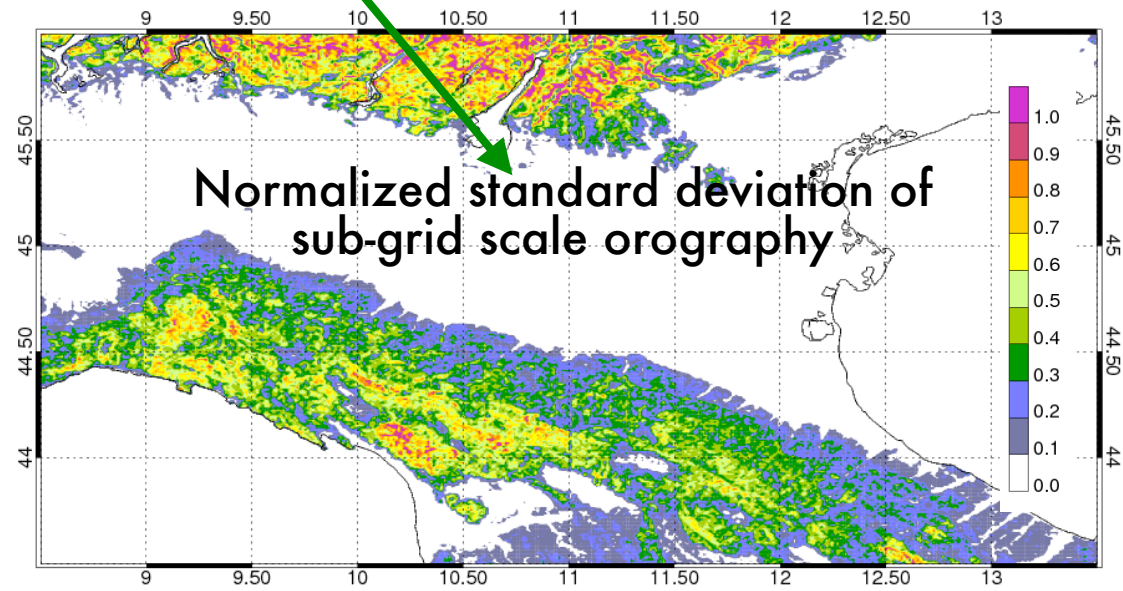
Observational error definition

$$Err = (1 - quality) + 0.2 \left(\frac{\sigma}{\sigma_{\max}} \right) quality$$



The quality is a single value that characterizes the final reliability of each radar datum (Fornasiero, 2006). It takes into account:

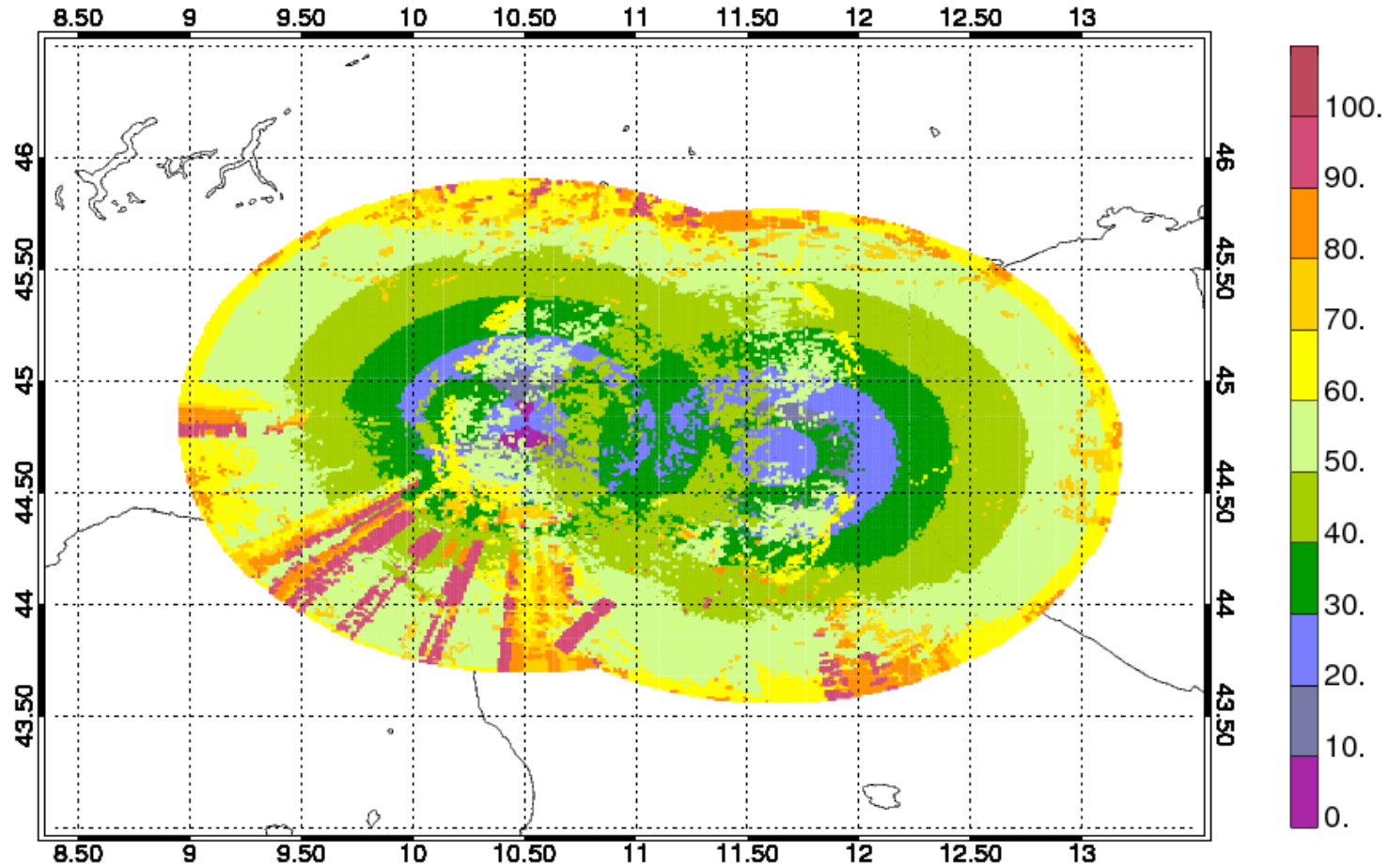
- Distance from radar
- Anomalous propagation
- Beam Blocking
- Path Integrated Attenuation

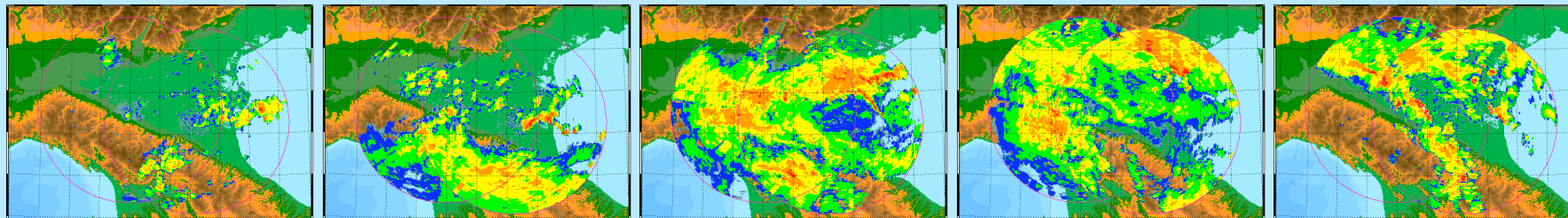


The dataset used to calculate the normalized sub-grid scale orography standard deviation is a NASA high resolution DEM (~90 m).

The maximum standard deviation used for normalization is chosen in the region of interest.

Error map





04/05/2007
00:00

04/05/2007
03:00

04/05/2007
06:00

04/05/2007
09:00

04/05/2007
12:00

CTRL run (nudging of the standard AOF)

Forecast

Model first-guess

Off-line ID-Var

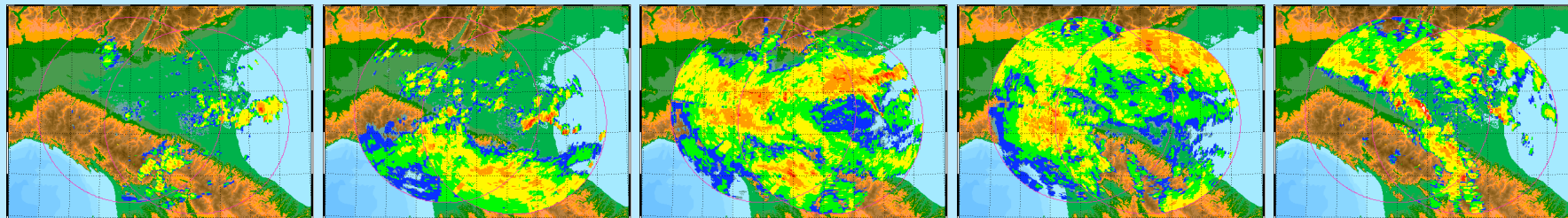
Humidity and temperature profiles

Database

New AOF generation

Experimental run with the nudging of the new AOF

Forecast



04/05/2007
00:00

04/05/2007
03:00

04/05/2007
06:00

04/05/2007
09:00

04/05/2007
12:00

CTRL

CTRL run (nudging of the standard AOF)

Forecast

Model first-guess

Off-line 1D-Var

Humidity and temperature profiles

Database

New AOF generation

EXP

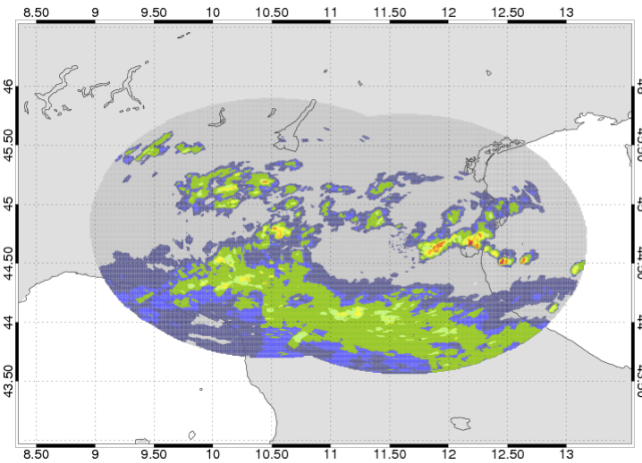
Experimental run with the nudging of the new AOF

Forecast

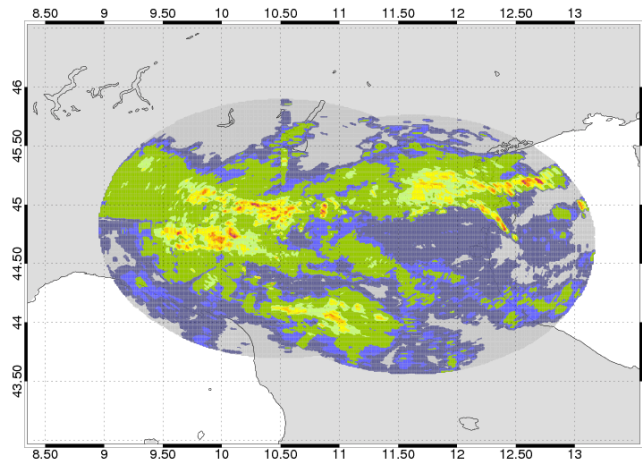
Case study

The selected case study is a convective episode in the Po valley which occurred between the 2nd and the 6th of May 2007. The strongest precipitation events occurred on the 4th of May and for this reason this day has been chosen for the assimilation experiment.

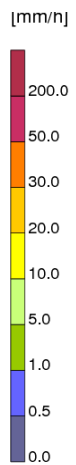
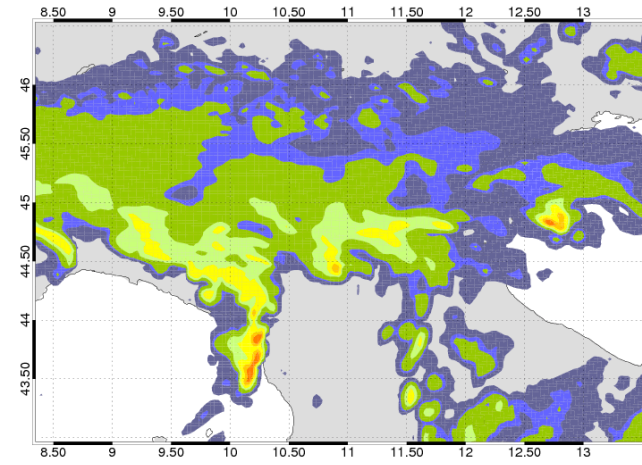
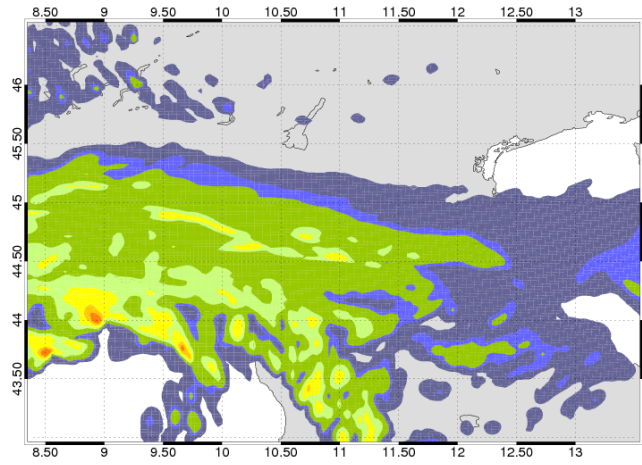
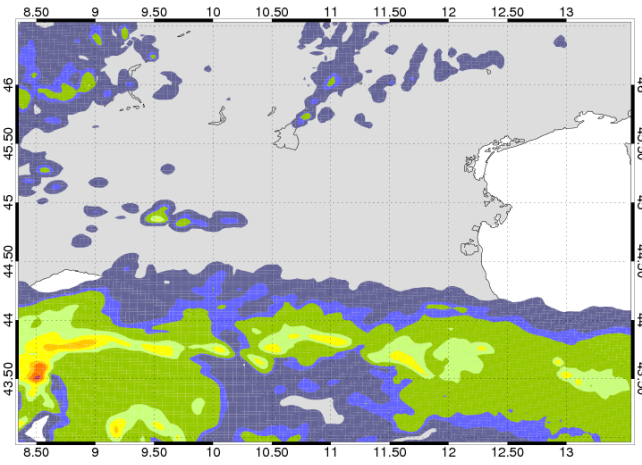
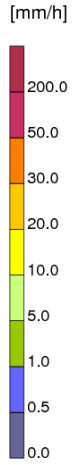
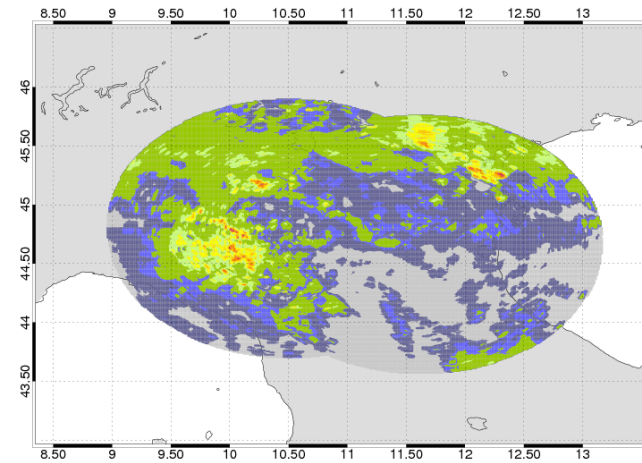
Surface Rain Rate 2007/05/04 03:00



Surface Rain Rate 2007/05/04 06:00



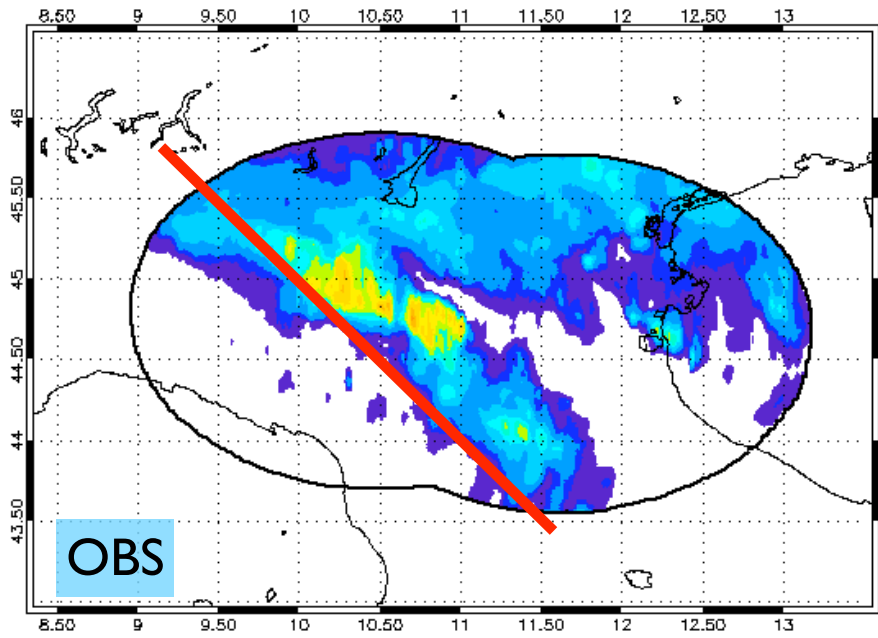
Surface Rain Rate 2007/05/04 09:00



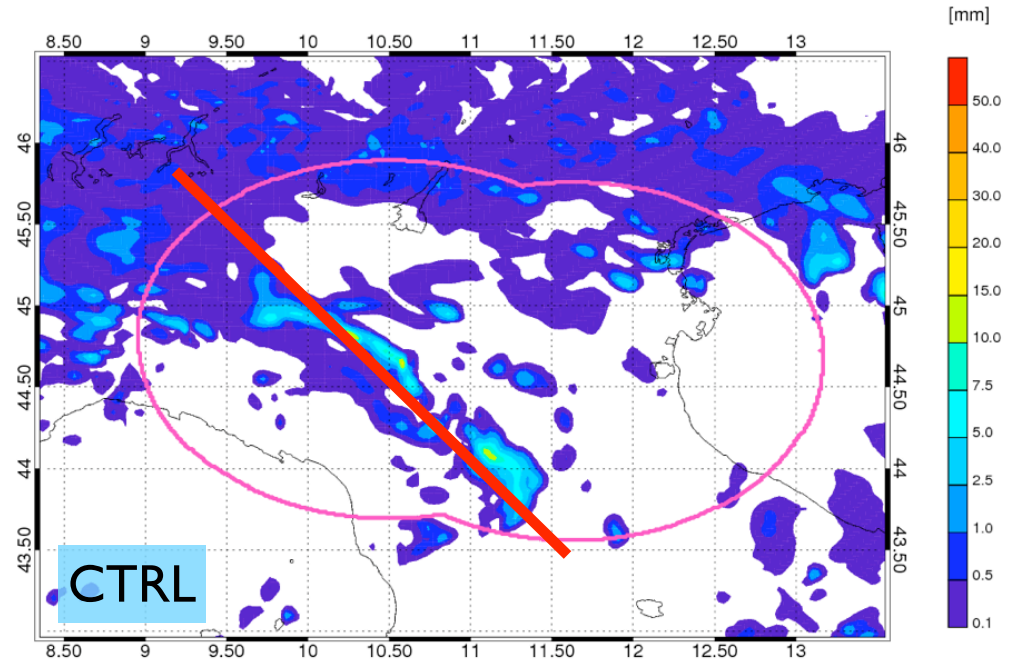
"Data thinning" impact

ASSIMILATION CYCLE: hourly precipitation at +12 hrs

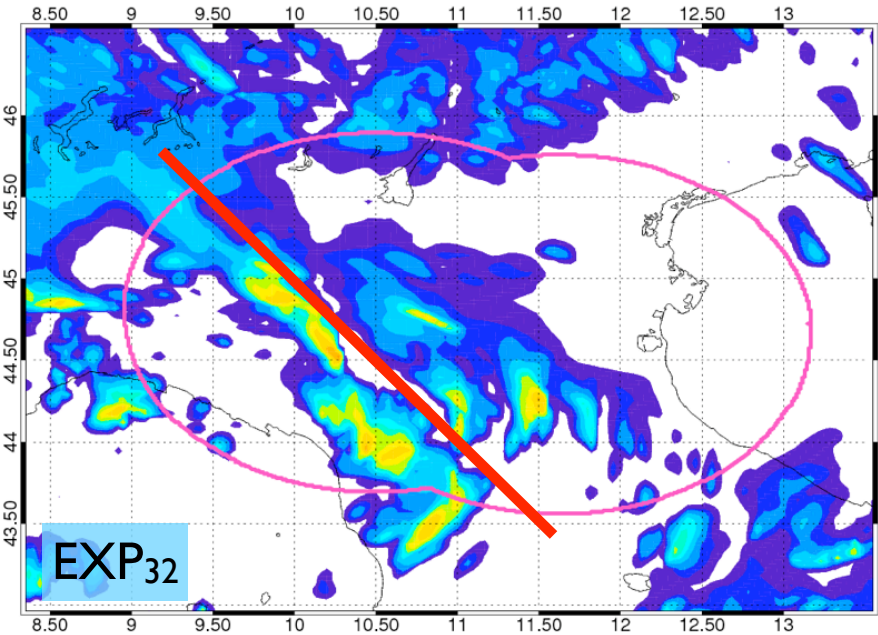
[mm]



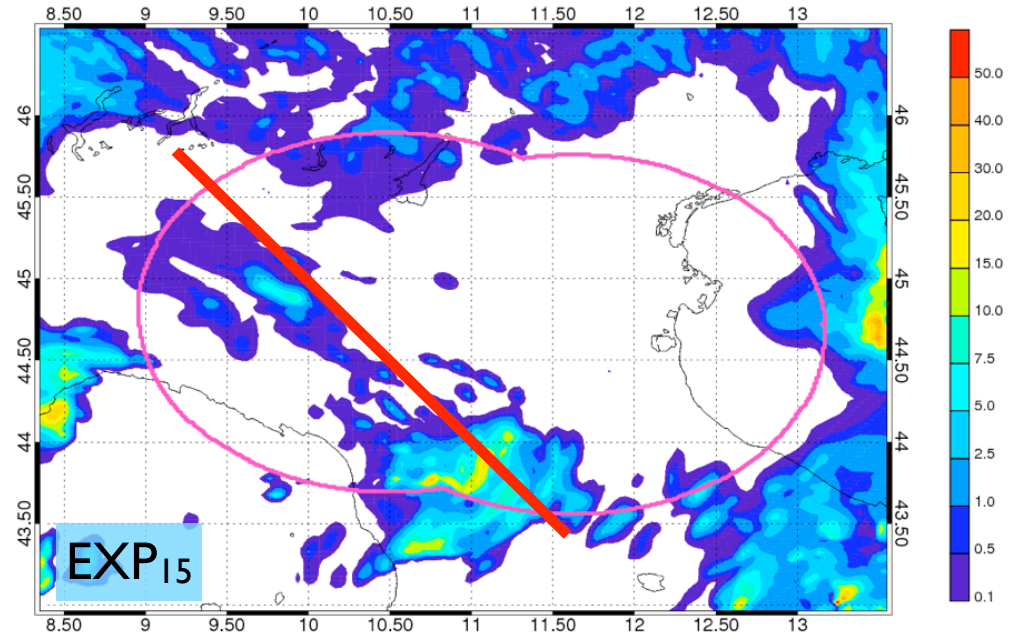
[mm]



[mm]



[mm]

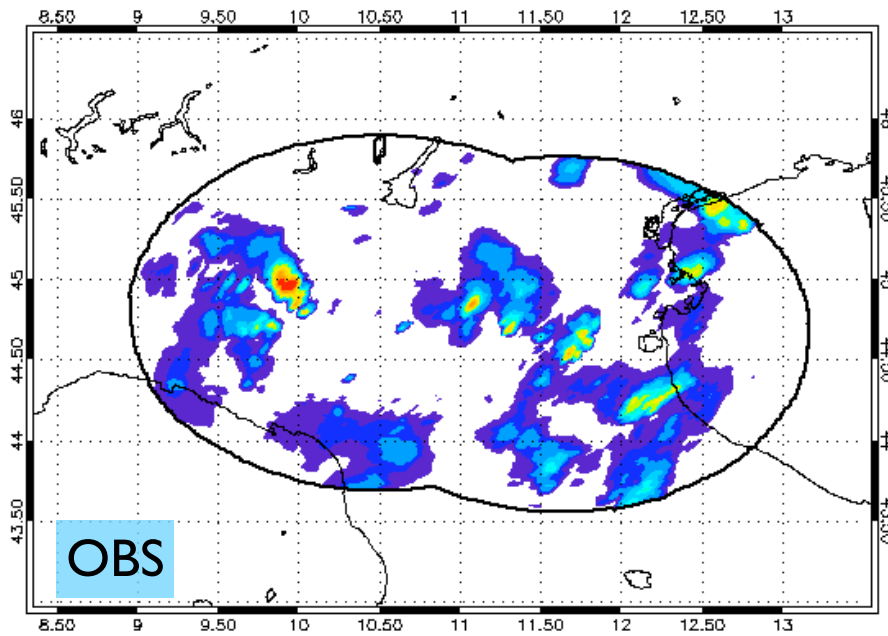


[mm]

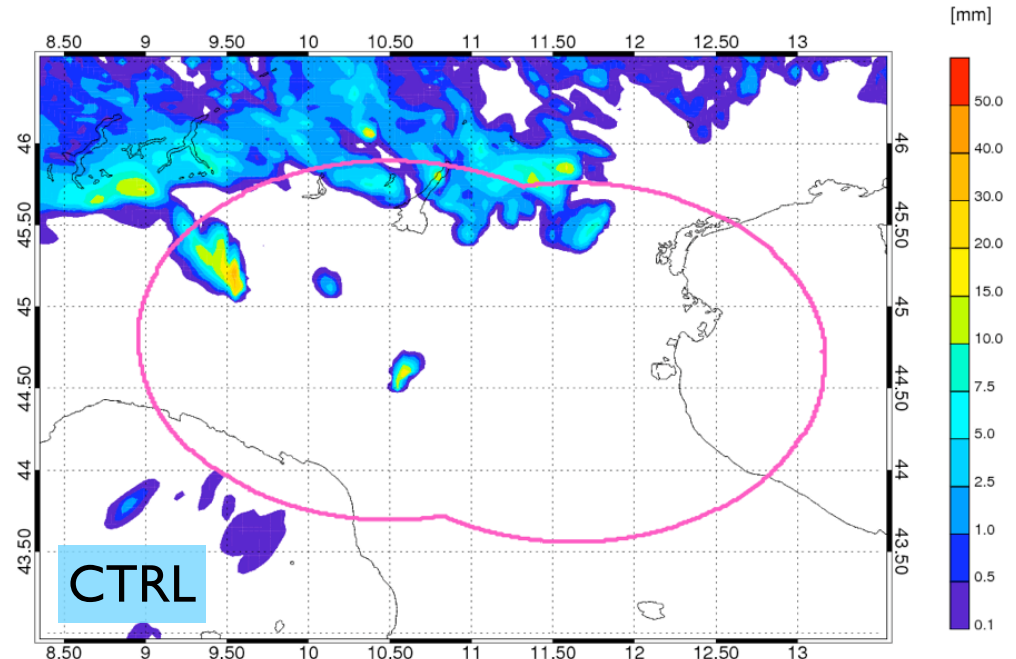
"Data thinning" impact

FORECAST CYCLE: hourly precipitation at +6 hrs

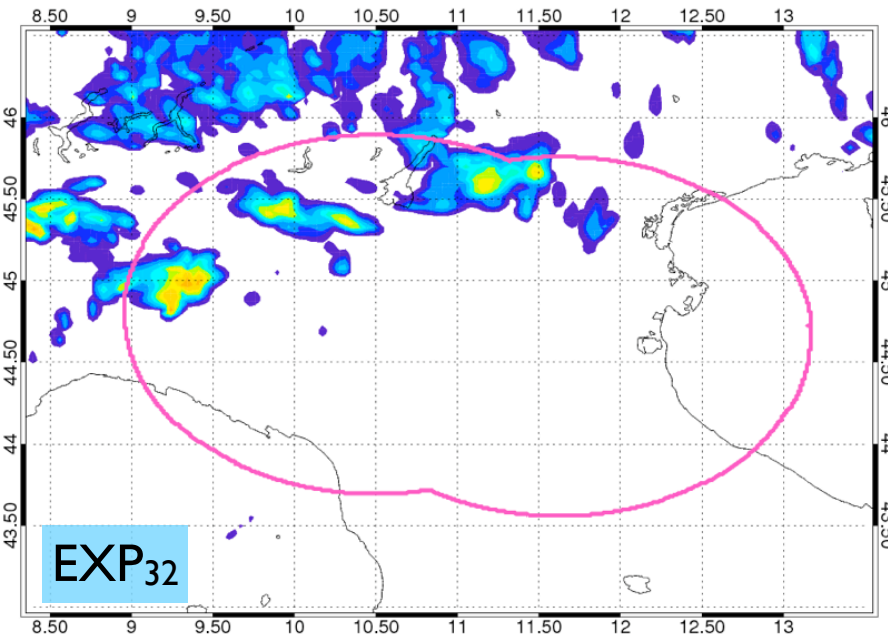
[mm]



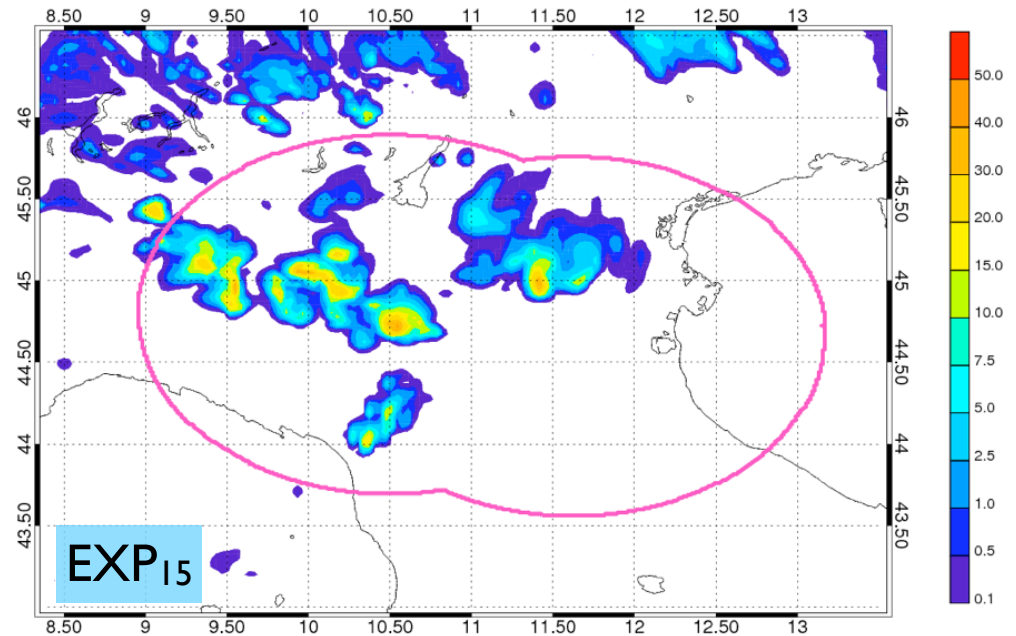
[mm]



[mm]



[mm]



[mm]

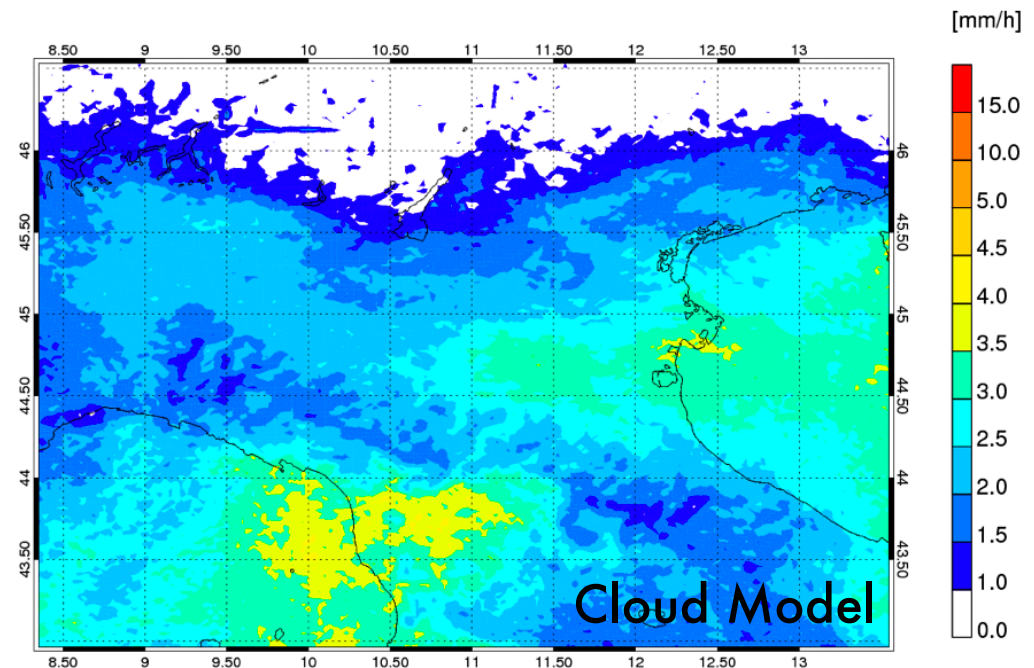
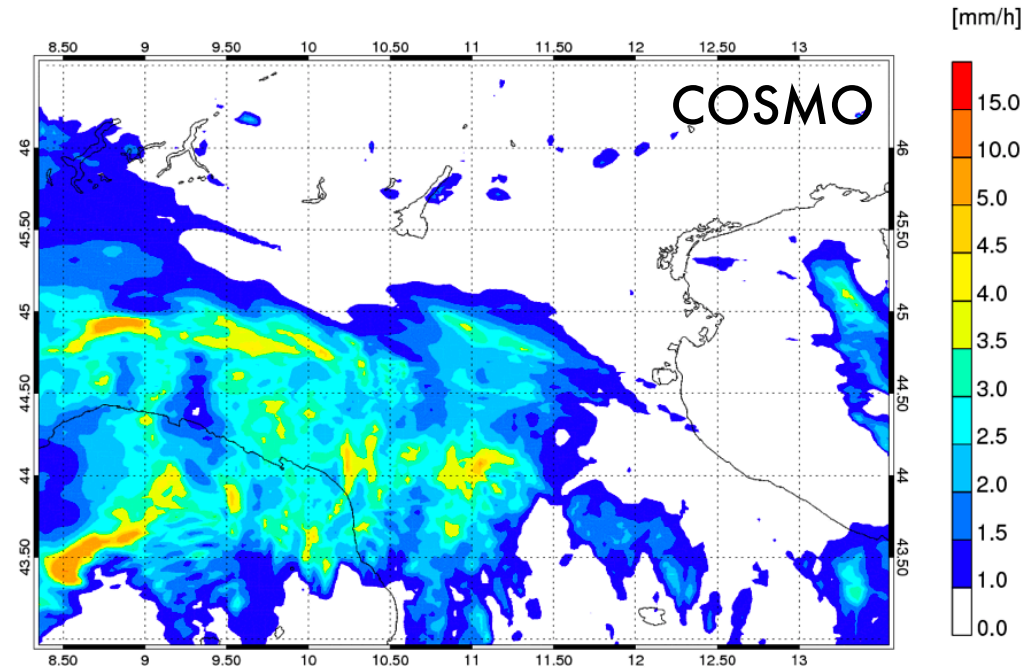
Model calibration (I)

The variational approach works in a statistically **optimal** way if observations and model errors are unbiased.

In our system cloud model has a different physics with respect to the actual one implemented into the COSMO model.

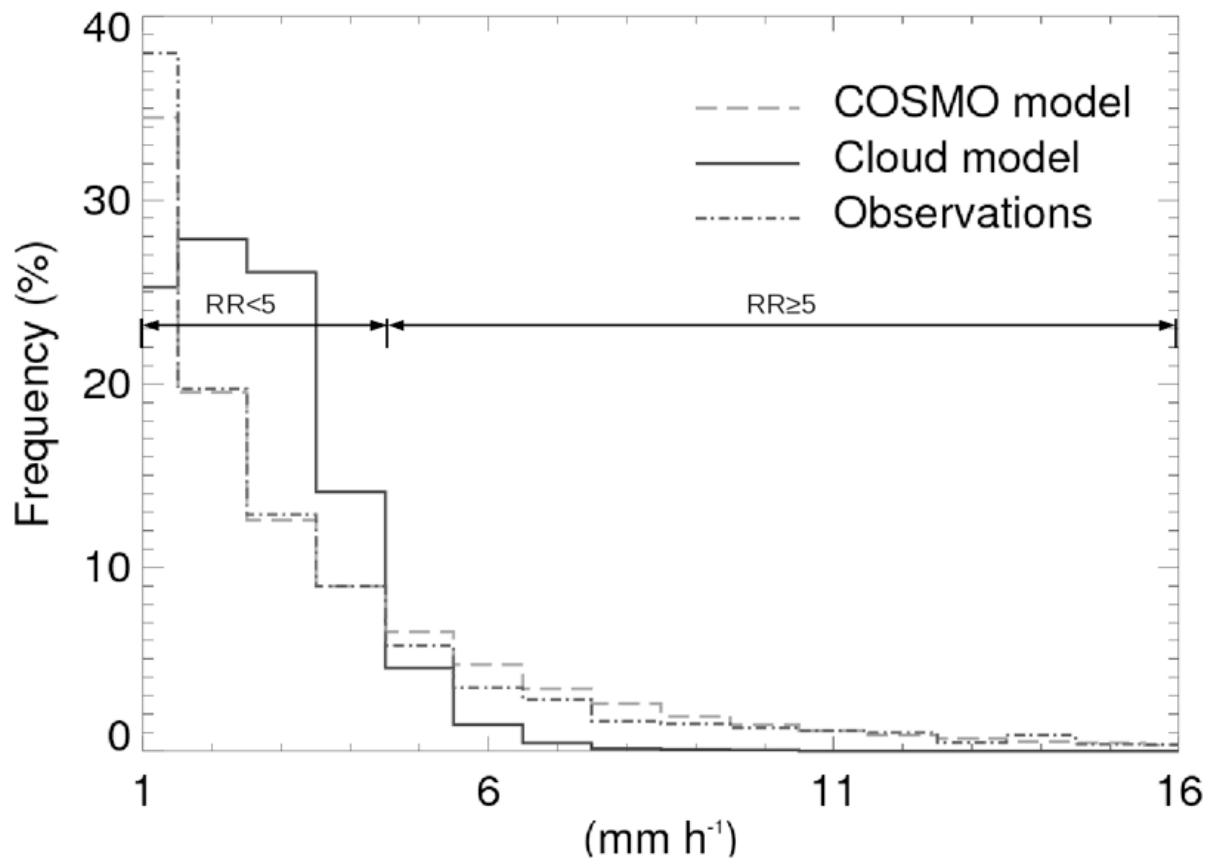
To quantify the difference between the two models the instantaneous surface rain rate has been analyzed.

Given a set of temperature and humidity profiles the mean properties of the cloud model generated precipitation field diverge from the ones which would be produced by the COSMO model. Precipitation is not only determined by the “physical” balance of the total water contained in a 1D column but it also depends on dynamical driven processes. The simplified cloud model cannot take these effects into account.



Model calibration (II)

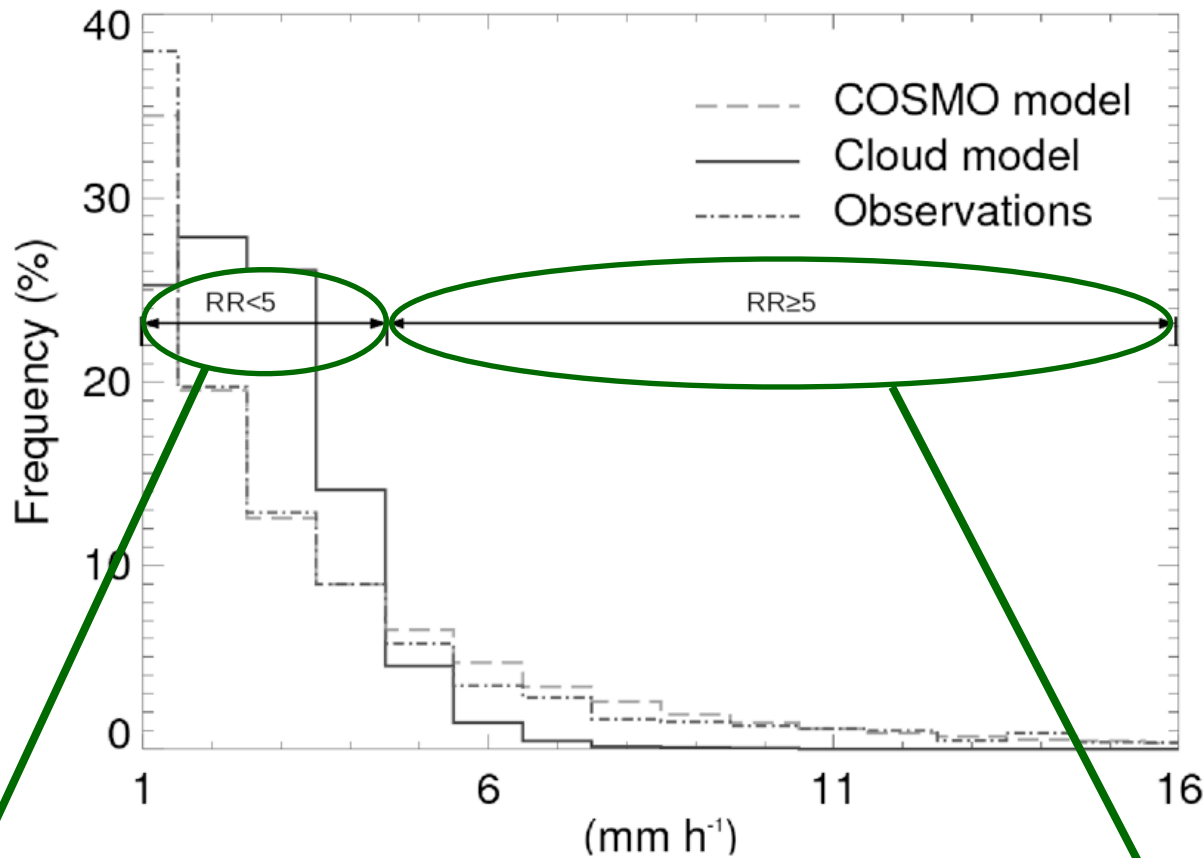
RR frequency histogram for cloud model, COSMO and radar observations.
Only precipitation data greater than 1 mm h⁻¹ are considered.



Bias correction on observations: $\overline{RR_{fg}} - \overline{RR_{obs}} = \overline{RR_{cosmo}} - \overline{RR_{obs}}$

Model calibration (II)

RR frequency histogram for cloud model, COSMO and radar observations.
Only precipitation data greater than 1 mm h⁻¹ are considered.



$$RR_{fg} > RR_{COSMO} \cong RR_{obs}$$

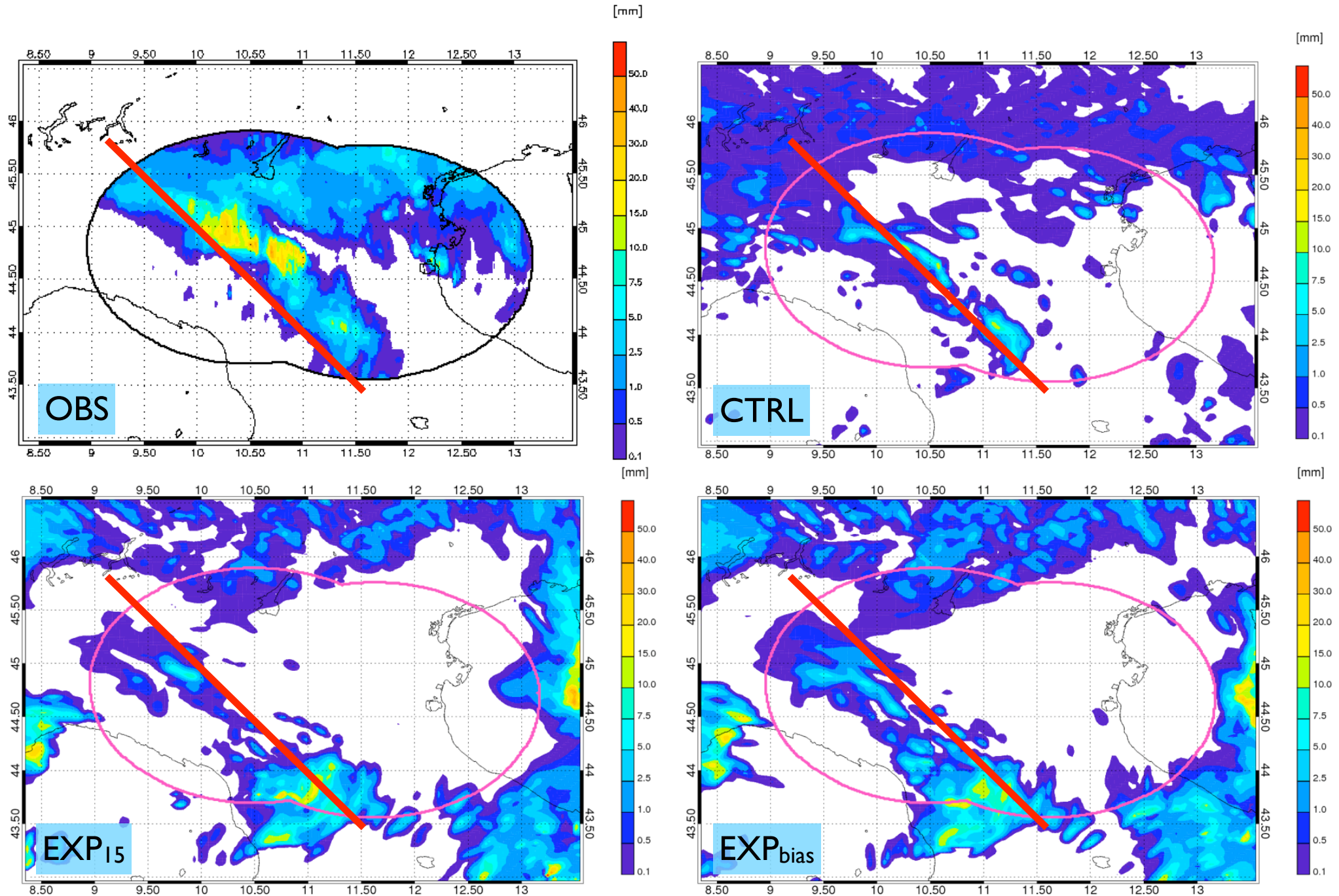
$$RR'_{obs} = RR_{obs} + 0.5 \text{ mm h}^{-1}$$

$$RR_{fg} < RR_{COSMO} \cong RR_{obs}$$

$$RR'_{obs} = RR_{obs} - 2 \text{ mm h}^{-1}$$

Calibration impact

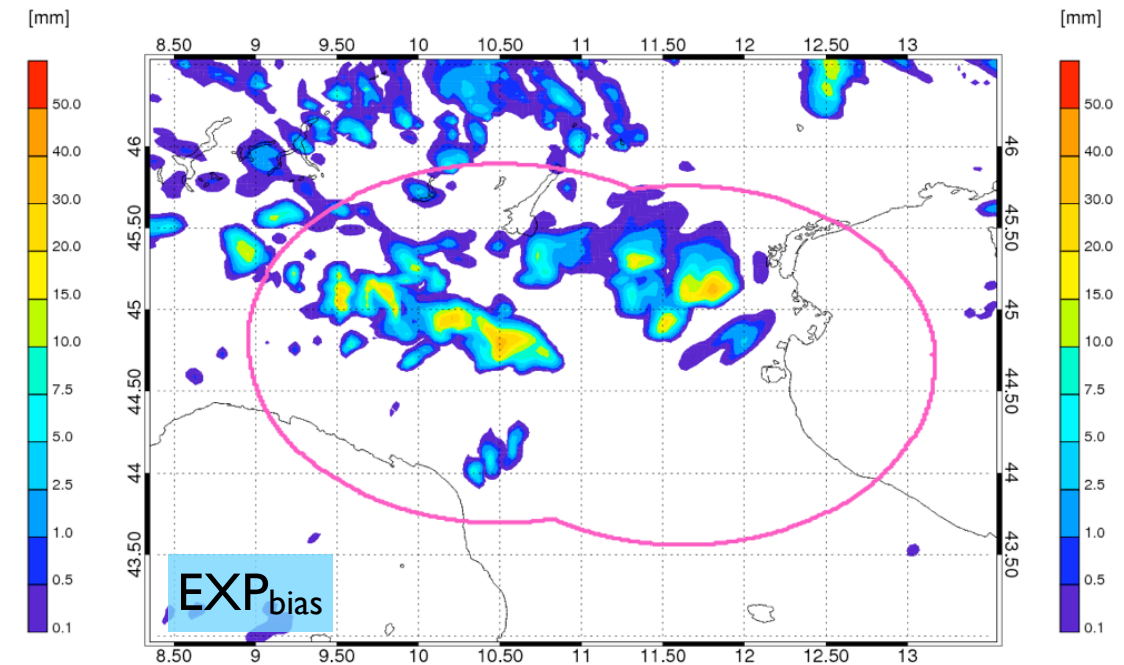
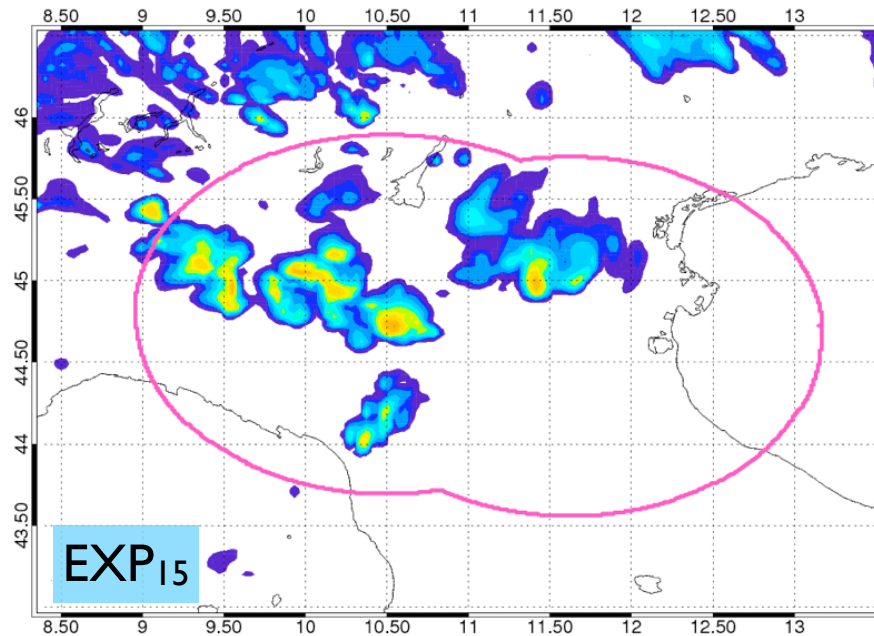
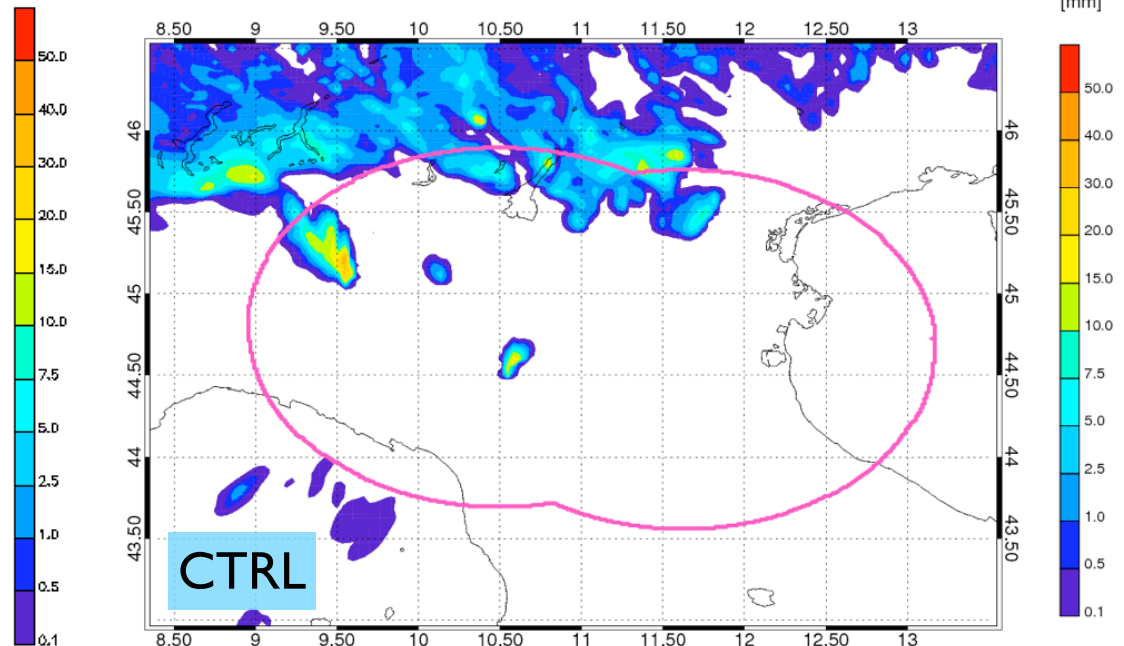
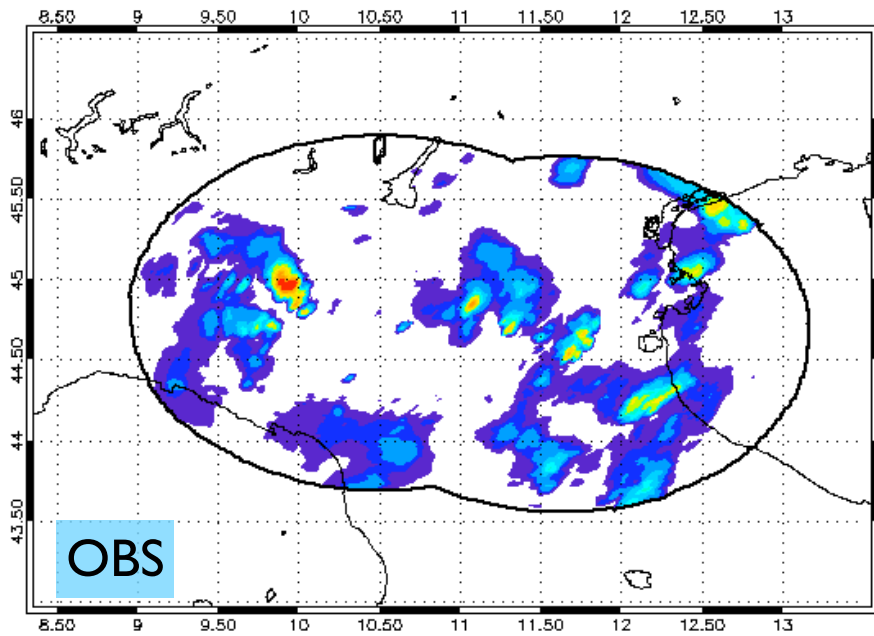
ASSIMILATION CYCLE: hourly precipitation at +12 hrs

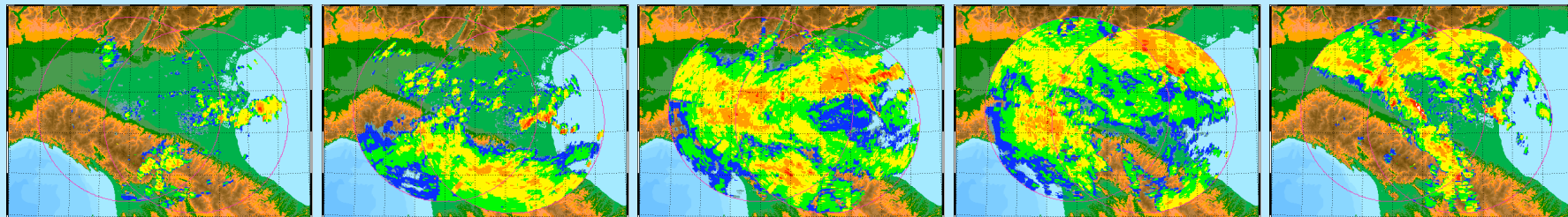


Calibration impact

FORECAST CYCLE: hourly precipitation at +6 hrs

[mm]





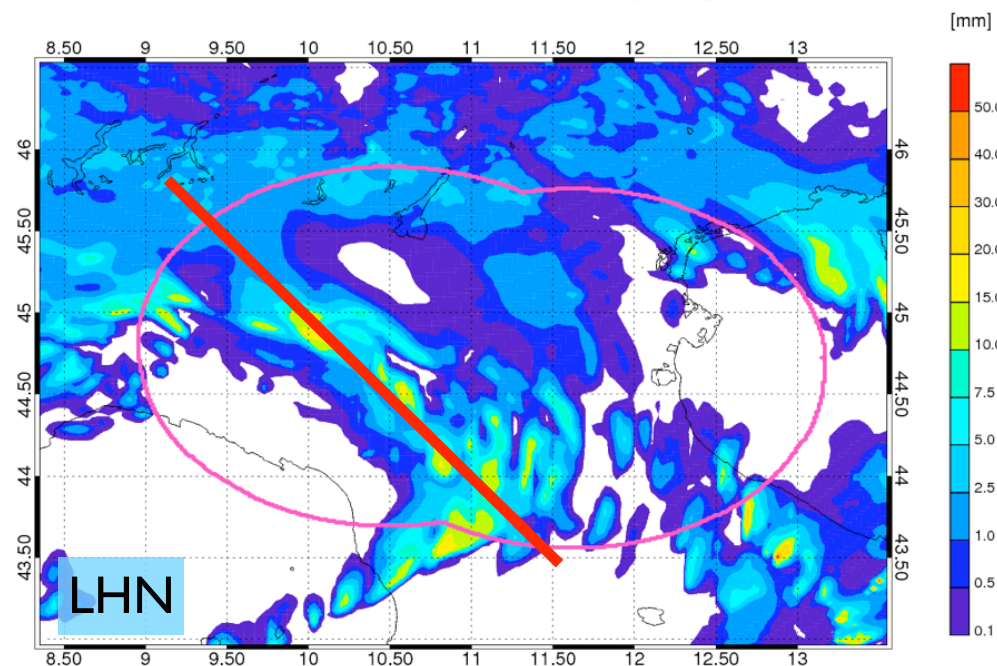
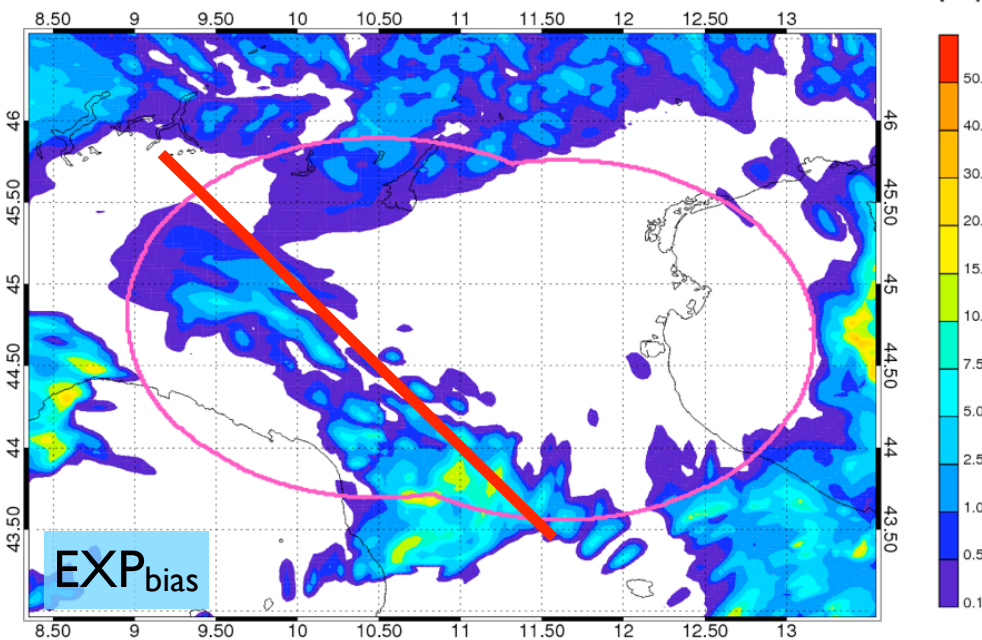
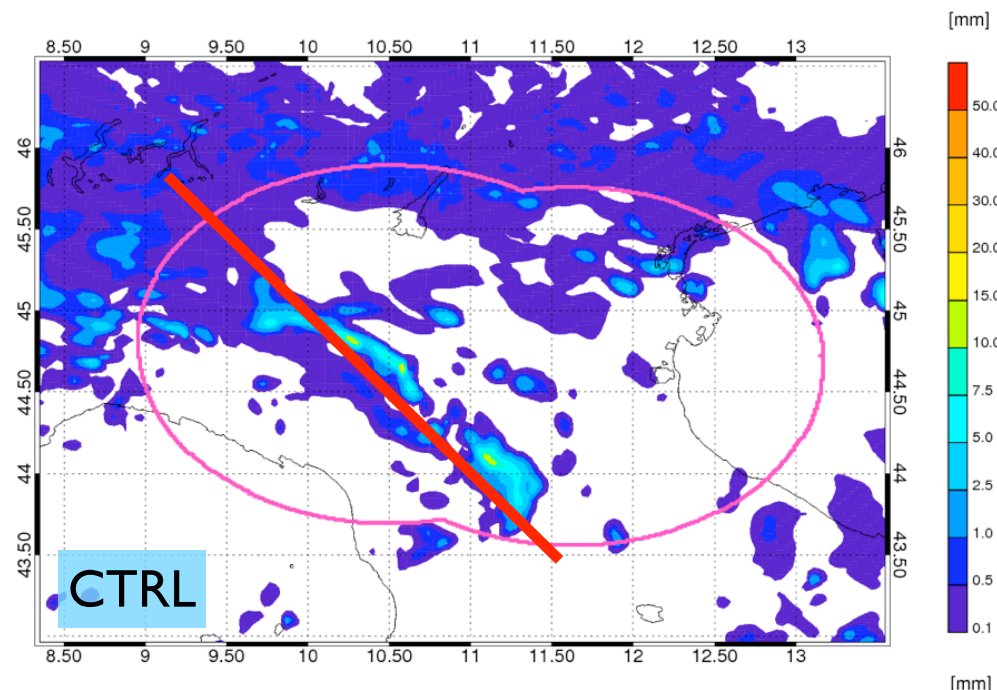
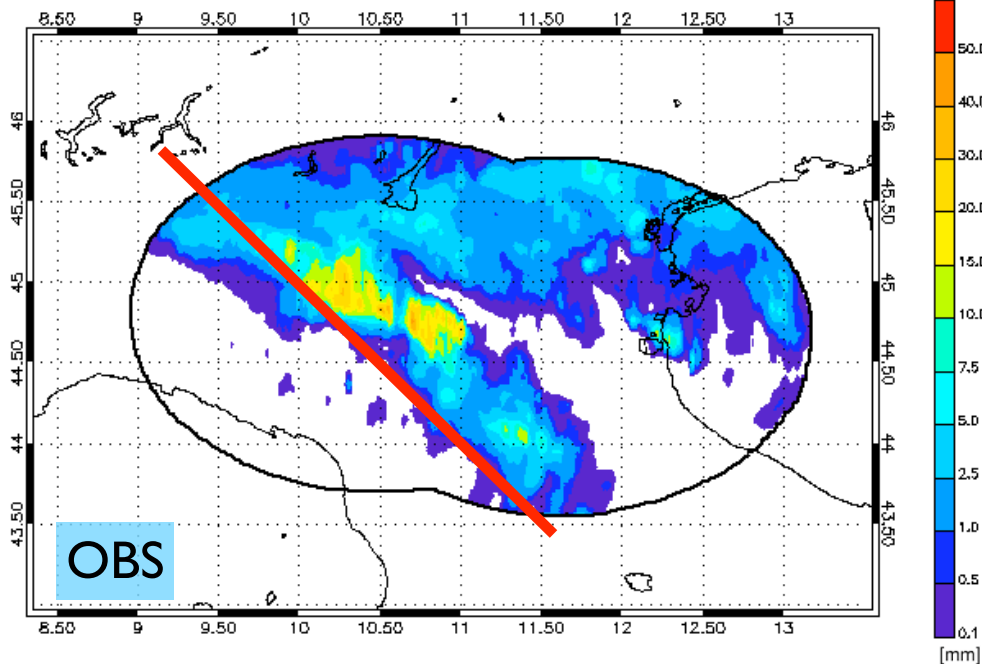
Radar surface rain rate



1D-Var vs LHN

ASSIMILATION CYCLE: hourly precipitation at +12 hrs

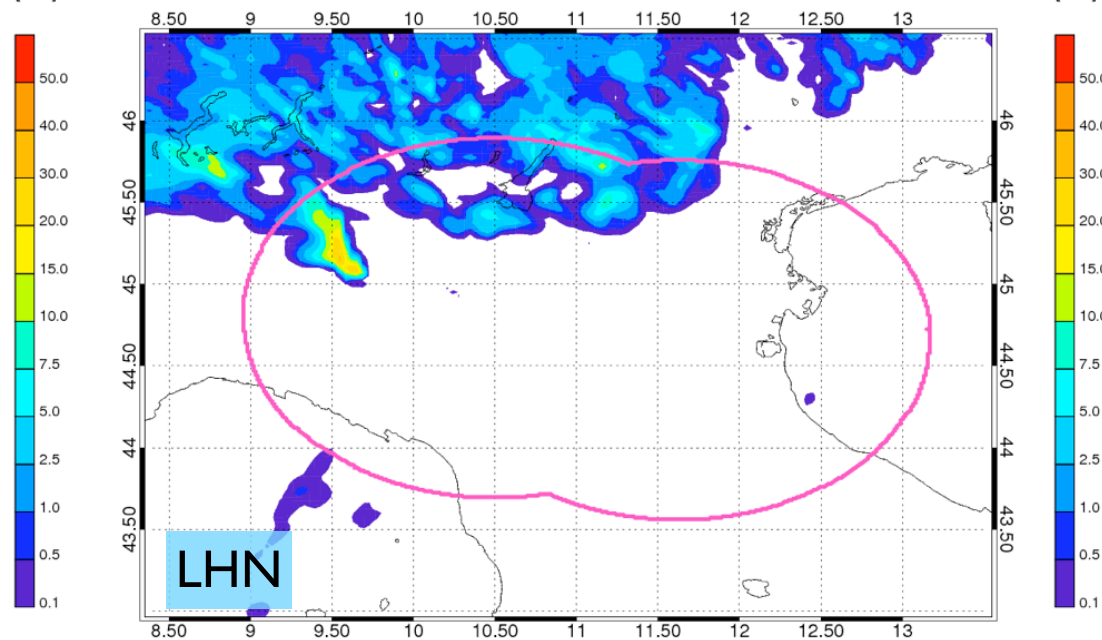
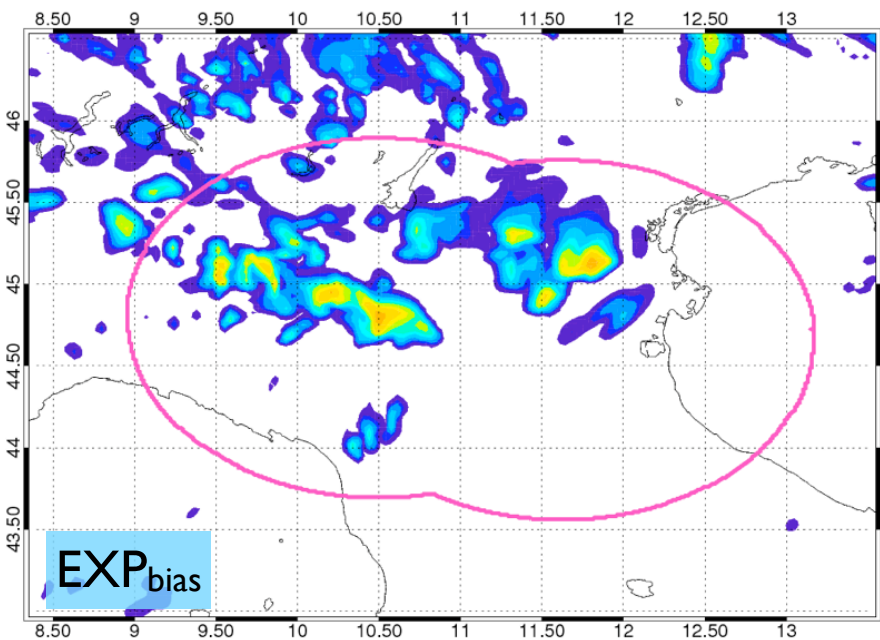
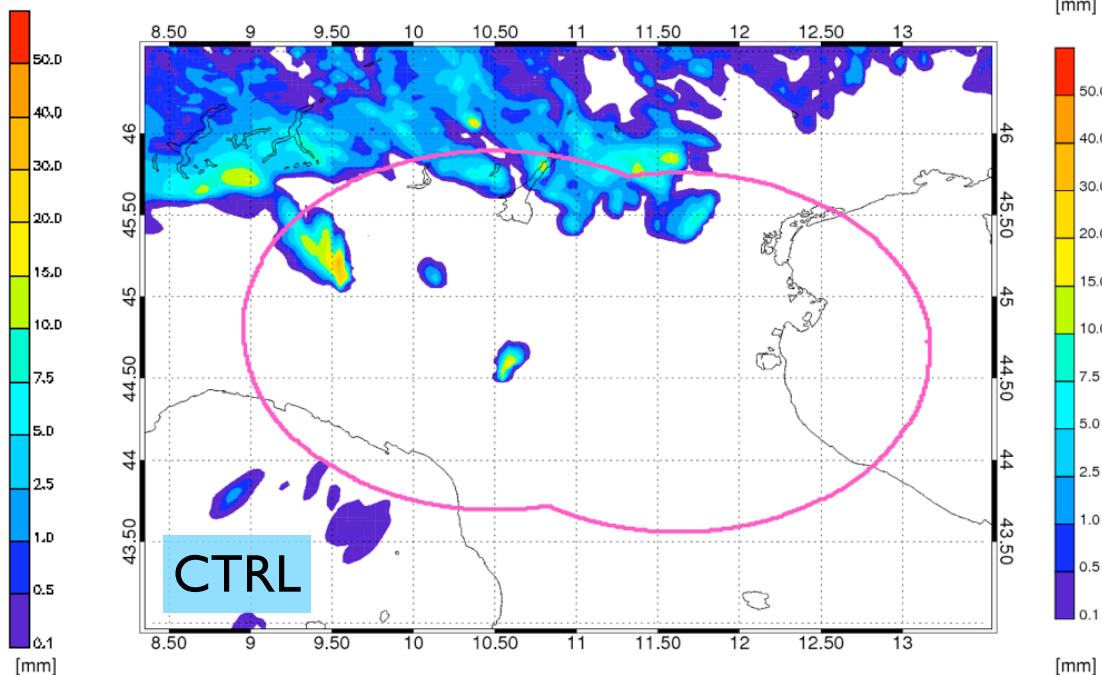
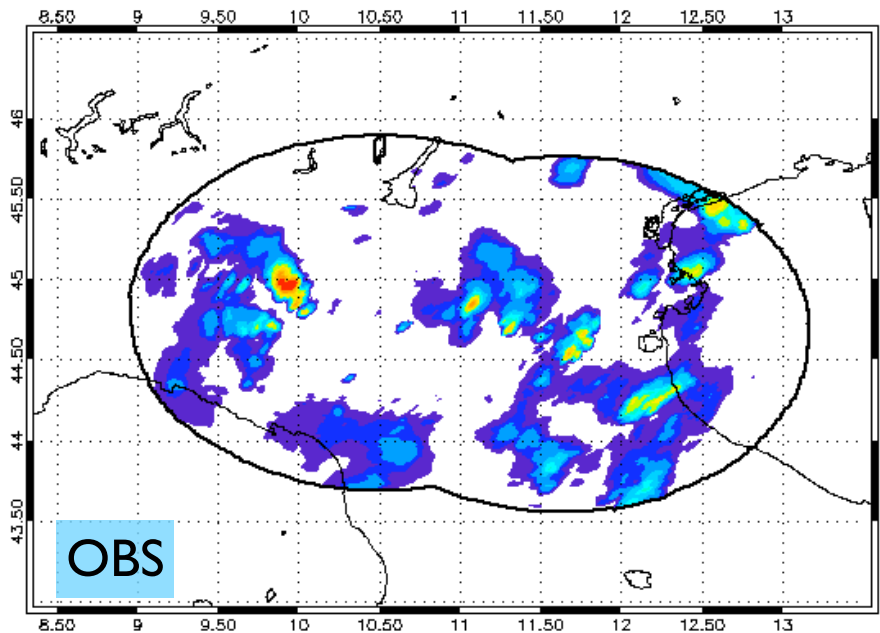
[mm]



1D-Var vs LHN

FORECAST CYCLE: hourly precipitation at +6 hrs

[mm]



Conclusions


- A 1D-var + nudging system has been developed in order to assimilate radar surface rain rate
- The system has been optimized:
 - making observations “data thinning”
 - estimating the observational error
 - calibrating the model
- Tests made show that the system shows promising results even if
 - results are strongly dependent on parameters used in nudging
 - in those areas where precipitation is observed, but not forecasted, the introduction of precipitating profiles do not trigger the model towards a precipitating condition
- A preliminary comparison with LHN has been made

Future plans

COSMO 4.0

 Switch to the newest COSMO release

Emilia Romagna radar composite

 Test methods using italian radar composite (provided by the National Department of Civil Protection)

Only one case study

 Validation of methods using a test bed

Qualitative verification of results

 Quantitative verification of forecasts using independent observations

1D-Var: application of a mean bias

 Definition of a case dependent bias

